

→ ATLANTIC FROM SPACE WORKSHOP

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Regional sea level, sea state change in the Atlantic from space

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Content

Within the Atlantic context, we address contribution and limitation of remote sensing to the issues:

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- 1) Sea level and sea state in coastal zone
- 2) Regional sea level budget
- 3) Ocean circulation

IWW:

- 1) Implication of Thema
- 2) What is done
- 3) What can be done

1) Implication : Climate Change, Coastal protection

- 2) Done: Coastal sea level by altimetry and tide gauges, VLM
- 3) What is still needed



Coastal sea level (SL)

Climate change \rightarrow Absolute SL trends from altimetry (background) & relative SL TGs <u>altimetry minus tide gauge trends</u> (VLM, stars) smaller than 1 mm/yr in mean and RMS differences of few centimetres.

Hotspots of accelerated sea level rise on the Atlantic coast of North America (Sallenger et al. 2012)

On the eastern Atlantic (European coast) 2-3 mm/yr



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Coastal sea level (SL)



Sentinel-3A ground-tracks

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Coastal sea level



Sentinel-3 altimetry SAR (I) and PLRM (r) in 2016-2017.

Noise (above) and NPoints (below) against along-track distance from coast. Noise is the absolute value difference between consecutive sea level heights (SSH) measurements.

Standard deviation of S3A SLA for two SAR and two RDSAR products

(Fenoglio et al., 2018)

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Coastal sea level

(Dinardo et al. 2017) **Altimetry** better agreement then model to in-situ

COASTAL ZONE	PLRM TALES			PLRM SINC2				BSH					
Index	stdd cm	slope	Corr	Number of Points									
HELG	4.23	0.96	0.96	11.7	0.93	0.86	3.3	1.00	0.97	18	85	0.75	67
KOSE	11.5	1.240	0.96	37.6	1.16	0.86	2.4	1.02	0.97	7.10	0.84	0.75	35
LHAW	9.1	1.01	0.96	16.7	1.00	0.86	7.8	0.99	0.97	21.2	0.75	0.75	58
SASS	6.5	1.03	0.89	13.6	0.86	0.78	4.3	0.99	0.95	6.10	0.98	0.85	28
WARN	11.1	1.047	0.87	21.8	1.076	0.74	4.3	1.02	0.97	9.60	0.87	0.87	22

(Dinardo et al., 2017)

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Table 1: Statistics of the SLAio in situ cross-comparison for SAR, PLRM TALES, PLRM SINC2 and BSH Model dataset in the coastal zone (0-10 km). The distance from the station is between 0 and 10 kilometres.

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Coastal sea level



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Vertical land motion along the Atlantic Coast of North America





subsidence of eastern seaboard of NA continues with constant rate with GIA as the main deriver. Exceptions are related to areas of recent excessive groundwater extraction and dam retention.

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Coastal Sea state

(Dinardo et al. 2017) **Model** better agreement than altimetry with in-situ

Measurement		DWD SV	PLRM SWH										
Index	mean cm	stdd cm	Corr	slope	mean cm	stdd cm	Corr	slope	mean cm	stdd cm	Corr	slope	Number of Points
HELG SUD DWR	-13.6	24.8	0.96	0.98	-14.7	62.6	0.73	0.6	3.4	38.7	0.91	1.13	32
HELG NORD DWR	-8.7	25.7	0.95	1.05	-23.6	49.7	0.78	0 .92	7.1	39.1	0.87	0.94	39
WES DWR	18.4	19.1	0.89	0.85	-15.3	27.5	0.98	1.27	-1.5	15.7	0.95	0.93	15

Table 3: Statistics of the SWH in situ cross-comparison for SAR, PLRM TALES, DWD datasets in the coastal zone (0-10 km).The distance from the station is between 0 and 10 kilometres.

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Coastal Sea state

Wiese et al. 2018



S3A, J2 and C2 SWH assessed against in situ & and spectral wave model (WAM) simulations forced with meteorological data to evaluate the sensitivity to wind different spatial/temporal resolutions

SWH from satellite > in situ by 7 cm to 26 cm. Near coast satellite data quality decreases; but within 10 km off the coast, Sentinel-3A performs better than the other satellites

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Extremes

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Figure 3. (left) Surge at the time of the overflight predicted by BSHcmod and HyFlux2 simulations with various wind forcing (DWD/BSH, ECMWF/JRC and DWD/JRC from left to right) and derived from in situ data (triangle). (right) Profiles at the SARAL/AltiKa overflight of wind speed, significant wave height, and surge height derived from altimeter observations (blue) and from models (red for DWD/BSH, light green for ECMWF/JRC, dark green for DWD/JRC, orange for ERA-INTERIM, and black for NOAA/GFS). Gray lines correspond to observations before (continuous line) and after Cyclone Xaver (black points).

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Sea level and sea state in coastal zone

- 1) Implication : Climate Change, Coastal protection
- 2) Done: Coastal sea level by altimetry and tide gauges, VLM
- 3) What remains to do:
 - Reduce coastal gap < 2-3 Km
 - improve SLA accuracy and precision (error estimation) continuity coastal/estuaries/in-land waters

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- Enhanced Conventional
- Delay Doppler (DDA) altimetry/SAR
- fully focused SAR
- "swath-altimetry" (SWOT mission)
- Synthetic Aperture Radar imaging

Sea level and sea state in coastal zone

1) What remains to do (cont.):

For coastal altimetry we have to

separate SLA sea surface variability and altimeter noise-> we need
better models of the temporal correlation of sea surface variability (wrt
new technologies with improved temporal sampling GNSS-IR, SWOT) and
new methods to account for this (10d SSH physically decorrelated, SWOT 1-2d not).

For VLM we have to

connect GNSS VLM with altimetry

separate local and episodic/nonlinear VLM from VLM along entire coastlines and trends we need better **models/data/methods** (SAR, sea state, coastal tide models)

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Regional sea level budget

- 1) Implication : Climate Change & Coastal protection
- 2) Done: direct comparison of GRACE & altimetry & model simulations

3) What is still needed



Regional sea level budget from Global Inversion Method

- Idea of the global fingerprint inversion (Rietbroek et al., 2016)
 - Forward modeling of gravitationally-elastic rotationally consistent sea level patterns → fingerprints
 - Consistent treatment of reference frames
 - Time-variable amplitudes are fitted to timeinvariant fingerprints



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UBonn/IGG/APMG approach: Global inversion

GRACE (mass driven sea level) versus radar altimetry (total geocentric sea level) \rightarrow combine and resolve for sea level budget components, 1.74 from AVISO

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Budget for Greenland coast

UBonn/IGG/APMG approach: Global inversion + Ocean Modelling

Aim:

- Resolve consistently for mass changes over land and steric changes in the ocean
- Improve understanding of changing dynamics of Greenland icesheet (Greenland ice sheet Ocean Interaction, GROCE)

Hagen

Glacier

Zachariæ Isstrøm

• Feedbacks of Greenland glacier melting in the NA and Arctic Focus

Analyzes

- mass balances of Greenland (GRACE gravimetric data)
- sea level budgets of NA (radar altimeter data)

We study relevant changes in temperature and salinity by oceanographic modelling

Focus: 79N glacier & North East Greenland Ice Stream (NEGIS).

GROCE



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First modelling results: Influence of Greenland freshwater flux 4/4 CSA

• global ocean model FESOM (AWI); forced with JRA-55; temperature [°C] and salinity [psu]



Further plans: Compare with GRACE and SWARM data and satellite altimetry

- Increase model resolution in the Arctic and around the 79N glacier

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1) Implications: Climate Change & Coastal protection

- 2) Done: direct comparison of GRACE & altimetry & model simulations
- 3) What is still needed :

Improved data (sea ice altimetry and GRACE, climatic data) assimilation in model simulations

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Dynamic Topography

- Implications: Climate Change & Coastal protection & Vertical Datum Unification & Navigation
- 2) Done: various methodologies tested (ocean & geodetic approach) for low resolution surface, here time variable approach
- 3) What is still needed :



Estimation of the MDT



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Estimation of the DT

Time variability can be modelled make the FE coeffcients continuous functions in time $\sum_{i=1}^{n} \sum_{j=1}^{n} \sum_{i=1}^{n} \sum_$

$$\zeta(\theta,\lambda,t) = \sum_{k \in K} a_k(t) b_k(\theta,\lambda) = \sum_{k \in K} \left[\sum_{i \in I} a_{ki} B_{ki}(t) \right] b_k(\theta,\lambda) \qquad \text{(Müller, 2014, Müller etal)}$$

Example result: Dynamic Topography for 5/2005, linear finite elements in space, B-Splines in time



DT from project PARASURV (DFG)

PArametric determination of the dynamic ocean topography from geoid, altimetric sea surface heights and SAR derived **RA**dial **SUR**face Velocities

Goals:

- Develop and implement a C¹smooth FE space to model MDT
- Observation equations link (reduced) current observations to MDT (radial SAR doppler derived, SKIM, ...)
- Implementation in HPC environment for real data analysis (if high quality data sets exist)



Dynamic Topography

- Implications: Climate Change & Coastal protection & Vertical Datum Unification & Navigation
- Done: various methodologies tested (ocean & geodetic approach) for low resolution surface, here time variable approach
- 3) What is still needed : high resolution surface in space and time improved input data (altimetry for coast, currents, winds for validation and integration improved methods for identifying eddies (fast, automated, e.g. machine learning)

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Conclusions (Bullets)

- Relative & absolute sea level: Satellite observation are an unique tool for estimating absolute sea level far from the coast and at coasts when tide gauge are missing or incomplete. Together with GPS give sea level change relative to coast, together with tide gauge give Vertical Land Motion. Needed: improved and continuous data processing, no coastal gap
- Regional Sea level budget: Altimetry and gravimetry essential for identifying ATL mass balance, allowing separation of steric and mass components, steric effects at distance (ATL). Assimilation of satellite observations in ocean model simulation/reanalysis and land hydrology to be fostered. Needed: improved and continuous data. Hydrological, oceanic and solid Earth changes need to be studied together contribution and limitation of remote sensing.
- **Dynamic Topography:** high resolution dynamic topography needs high resolution input data. Multi-mission altimetry + wide-swatch altimetry needed to detect high features in open ocean and at the coast (e.g. gyres). Additional input : ocean currents by SAR image are extremely important but not yet widely available, quality need to be improved

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Recommendation (on future investments on the Atlantic region)

- New data, technologies (from space) : continuity of altimetric and gravimetric missions.
 - Sea level (GNSS reflectometry, 2D wide-swath SWOT) coastal gap and spatial resolution
 - Sea state (from SAR and CFOSAT etc.)
 - Surface Currents (from SAR, SKIM mission)
- New data (in-situ) : Co-located TGs & GPS are needed, easy access, central efforts
- New data (model) : improved meteorological data (e.g. . COSMO-REA reanalysis)
- State-of-the art ocean models (downscaling to coast, e.g. NEMO (regional), UG estuarine and coastal SCHISM model - with fine resolution (unstructured grid models)
- Integrating waves, circulation, atmosphere and HD model (e.g GCOAST system)
- Improved in-situ/remote sensing data assimilation
- New methods (spatial correlation, pattern identification, machine learning)
- Hydrological, oceanic and solid Earth changes need to be **studied together** to better understand contribution and limitation of remote sensing

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Extra slides





Use Finite Elements on a triangulation: cover the ocean, extrapolate to the coast, flexible geometry, approximate the unknown function space



Linear Elements:

- Function values in the nodes
- 3 degrees of freedom
- Composed surface is C⁰smooth

ARGYRIS Element:

- Function values, 1st & 2nd derivatives in the nodes
- Normal derivatives in midpoints of edges
- 21 degrees of freedom
- Composed surface C¹smooth

The surface is continuously defined in the region of interest A full covariance matrix can be provided

Different observation types/functionals can be easily integrated (e.g. surface currents)

Finite element space determines smoothness conditions (C⁰ or C¹)

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DT: Account for Uncertainty of Observation 5/6

Example: iterative refinement of along track error model e.g by covariance functions or stochastic processes along the track for altimetric SSH observations

Setup of an iterative procedure:

- 1. Input data and information
 - predefined finite element setup (triangulation and base function type)
 - predefined spherical harmonic setup (maximal resolution)
 - ▶ NEQs of gravity field model: N_{gf}, n_{gf},
 - OEQs of ssh observation of a mission *i*: \mathbf{A}_{ssh_i} , ℓ_{ssh_i} , and initial covariance model $\mathbf{Q}_{ssh_i}^{(0)}$ • initial weights for all observation groups (w_i and w_{af})
- 2. Compute NEQs for ssh observations, $\mathbf{N}_{ssh_i} = \mathbf{A}_{ssh_i}^{\mathcal{T}} \mathbf{Q}^{(0)-1}_{ssh_i} \mathbf{A}_{ssh_i}$, $\mathbf{n}_{ssh_i} = \mathbf{A}_{ssh_i}^{\mathcal{T}} \mathbf{Q}^{(0)-1}_{ssh_i} \mathbf{A}_{ssh_i}$
- 3. Combine NEQs (different parameter spaces!)



- 4. Rigorous parameter estimation (huge dimensional) $\mathbf{\tilde{x}} = \mathbf{N}^{-1}\mathbf{n}$
- 5. Estimate weights for all observation groups (w_i and w_{gf}), continue with 3.
- 6. Compute SSH residuals $\mathbf{v}_{ssh_i} = \mathbf{A}_{ssh_i} \tilde{\mathbf{x}} \boldsymbol{\ell}_{ssh_i}$
- 7. Analyse \mathbf{v}_{ssh_i} and adjust a stochastic model to \mathbf{v}_{ssh_i} , update $\mathbf{Q}_{ssh_i}^{(1)}$ and continue with 2.



=> more realistic error model => better combination of different data sets

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