EXETER Centre for Geography and Environmental Science



Optical Remote Sensing for Marine applications

Dr Bob Brewin

Lecturer in Physical Geography,

University of Exeter,

ESA UNCLASSIFIED - For ESA Official Use Only

Penryn Campus, Cornwall, TR10 9FE, UK

(r.brewin@exeter.ac.uk)



Additional contact information can also be found on my website <u>https://geography.exeter.ac.uk/staff/in</u> <u>dex.php?web_id=Bob_Brewin</u>.

→ THE EUROPEAN SPACE AGENCY

A bit about myself: my path



2003-2010



BSc Surf Science & Technology MSc Applied Marine Science PhD Marine Remote Sensing Associate lecturer and demonstrator Marine Remote Sensing



2010-2019

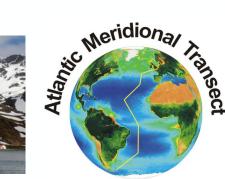


Research Scientist ESA post-doc fellow NCEO scientist



National Centre for Earth Observation

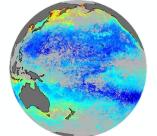




2019present



Lecturer in Physical Geography

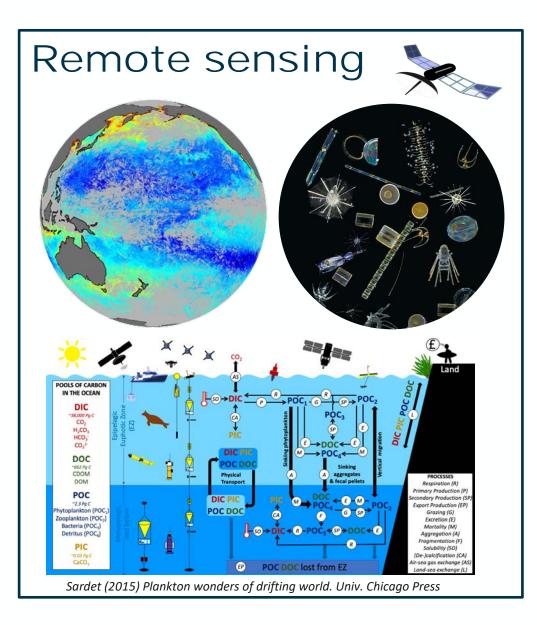


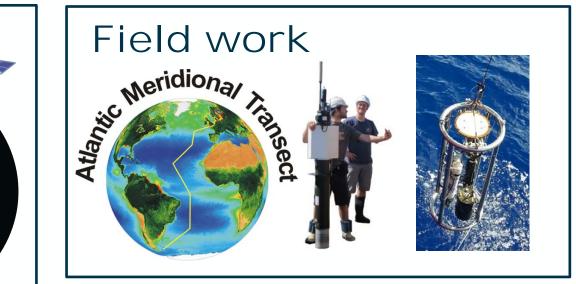




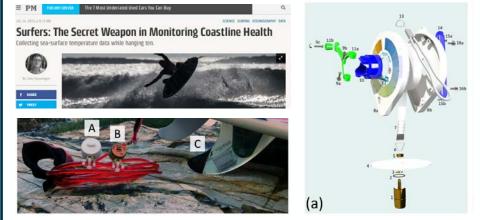
A bit about myself: research interests





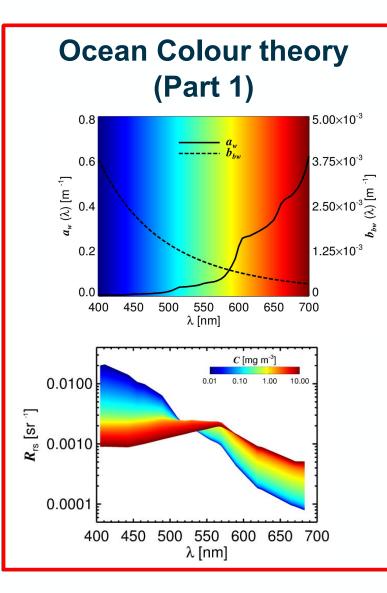


Citizen Science

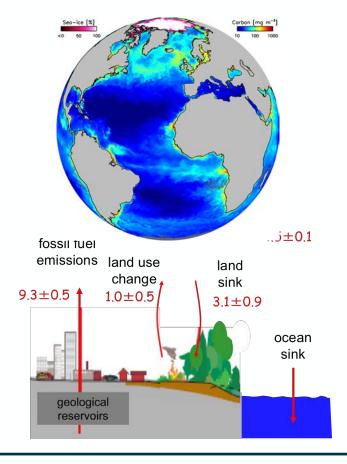


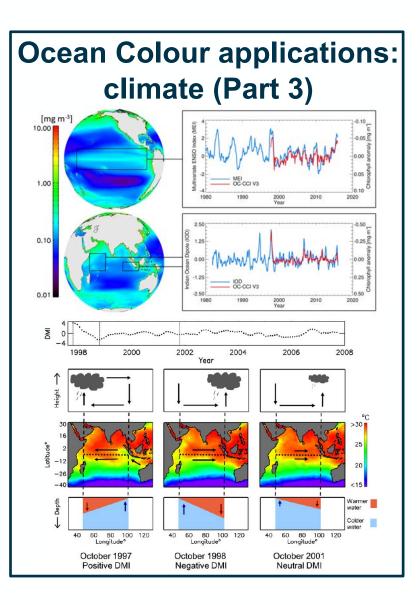
Lecture plan





Ocean Colour applications: carbon cycle (Part 2)





Part 1: Outline



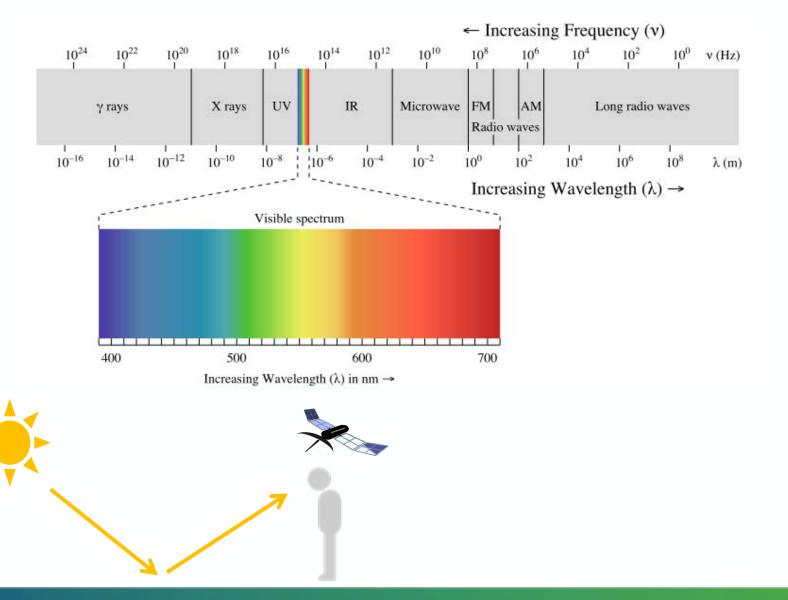
- 1) What is ocean colour?
- 2) History of ocean colour
- 3) Optical properties of water and its constituents
- 4) In situ measurements of ocean colour
- 5) Satellite remote-sensing of ocean colour

1) What is ocean colour?



Spectral variations in visible light reflected from water

Sunlight that enters the water and is absorbed, scattered and reflected by water and the constituents that make up the water. These processes vary with wavelength of light



Ocean Colour



Paintings by Albert Bierstadt, Emil Nolde, Ivan Aivazovsky

Early navigation: bottom topography, presence of ice bergs, river discharge, occurrence of fish (e.g., H. Hudson, First voyage, 1607; W. Scoresby, 1820)

OCEAN OPTICS FROM 1600 (HUDSON) TO 1930 (RAMAN) Shifting interpretation of natural water colouring

Marcel R. WERNAND & Winfried W.C. GIESKES

Wernand (2011) PhD Thesis <u>https://www.researchgate.net/publication/254886</u> <u>155 Poseidon%27s paintbox historical archives o</u> <u>f ocean colour in global-change perspective</u>





 Everyday observation: colour of water changes with time and place

- What is responsible for these changes?
- Can we quantify the responsible agents?
- Why should we care?

From CEOS Report 2018, Images Courtesy CSIRO

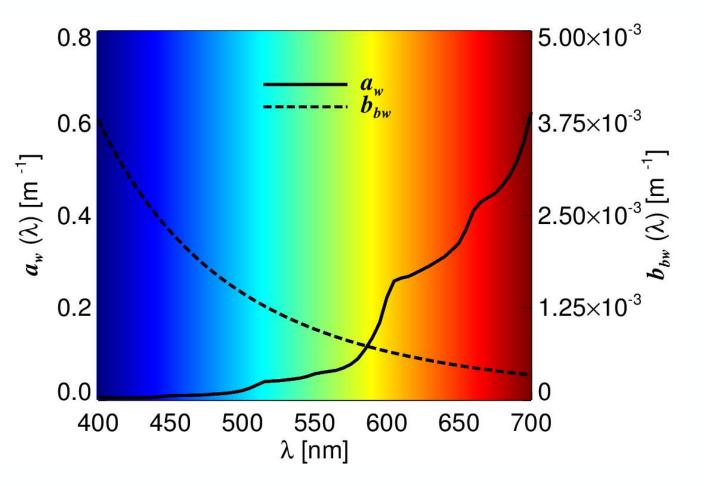
Slide kindly prepared by Shubha Sathyendranath



A sufficient deep layer of pure water exhibits by molecular scattering a deep blue colour more saturated than skylight and of comparable intensity. The colour is primarily due to diffraction, the absorption only making it of a fuller hue. The theories hitherto advanced that the dark blue of the deep sea is reflected sky-light or that it is due to suspended matter are discussed and shown to be erroneous

C. Raman, 1922

$$R(\lambda) = \frac{E_u(\lambda)}{E_d(\lambda)} \approx r \frac{b_b(\lambda)}{a(\lambda)}$$



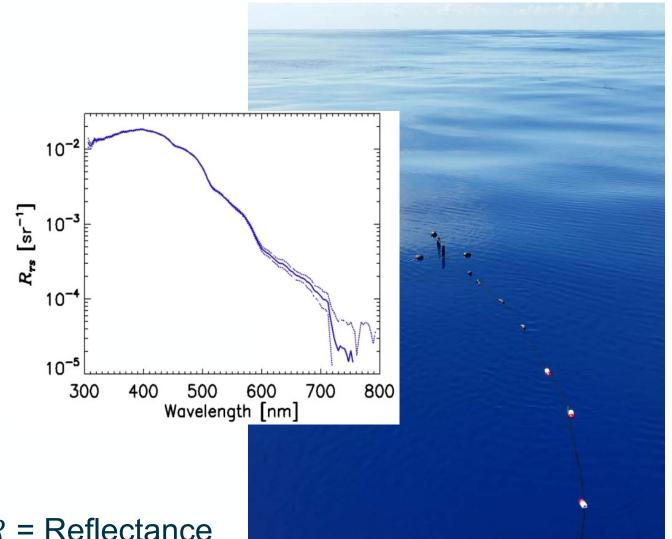
R = ReflectanceE = Irradiance (u = upwelling, d = downwelling) $a(\lambda) = \text{absorption}$ $b_b(\lambda) = \text{backscattering}$

A sufficient deep layer of pure water exhibits by molecular scattering a deep blue colour more saturated than skylight and of comparable intensity. The colour is primarily due to diffraction, the absorption only making it of a fuller hue. The theories hitherto advanced that the dark blue of the deep sea is reflected sky-light or that it is due to suspended matter are discussed and shown to be erroneous

C. Raman, 1922

$$R(\lambda) = \frac{E_u(\lambda)}{E_d(\lambda)} \approx r \frac{b_b(\lambda)}{a(\lambda)}$$

R = ReflectanceE = Irradiance (u = upwelling, d = downwelling) $a(\lambda)$ = absorption $b_b(\lambda)$ = backscattering





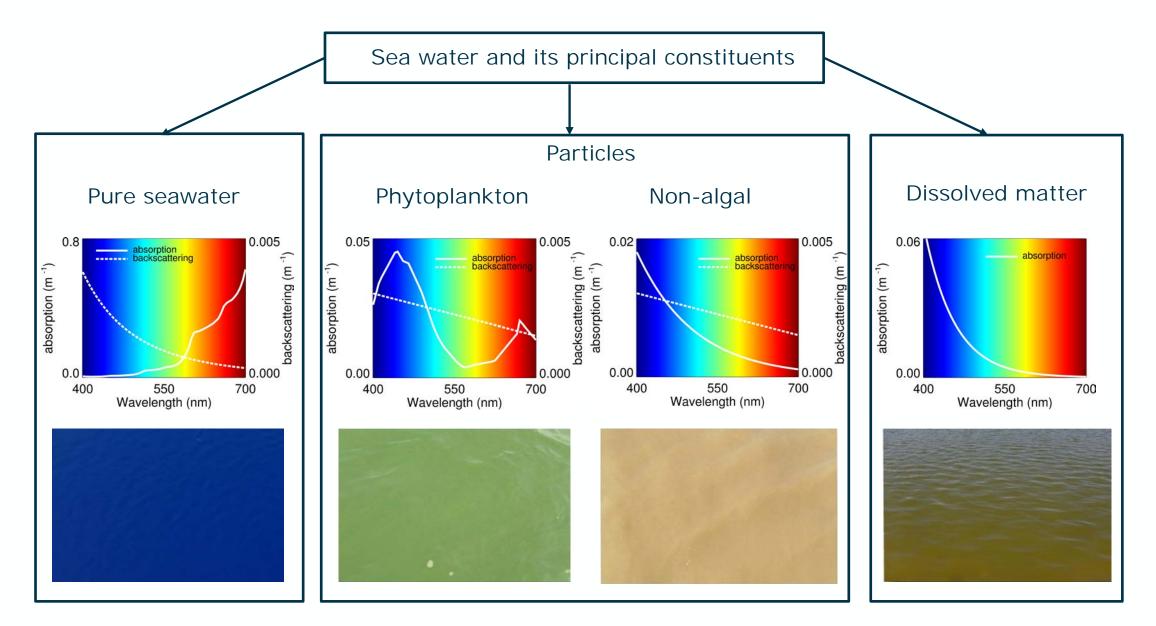
A sufficient deep layer of pure water exhibits by molecular scattering a deep blue colour more saturated than skylight and of comparable intensity. The colour is primarily due to diffraction, the absorption only making it of a fuller hue. The theories hitherto advanced that the dark blue of the deep sea is reflected sky-light or that it is due to suspended matter are discussed and shown to be erroneous

Taken from Wernand (2011, see link to reference on earlier slide)

C. Raman, 1922

$$R(\lambda) = \frac{E_u(\lambda)}{E_d(\lambda)} \approx r \frac{b_b(\lambda)}{a(\lambda)}$$

R = ReflectanceE = Irradiance (u = upwelling, d = downwelling) $a(\lambda)$ = absorption $b_b(\lambda)$ = backscattering



Apparent optical properties (AOPs) Dependant on the inherent optical properties of the water, and the directional distribution of the light field in the sea (e.g., water leaving radiance, reflectance and diffuse attenuation coefficient of seawater)

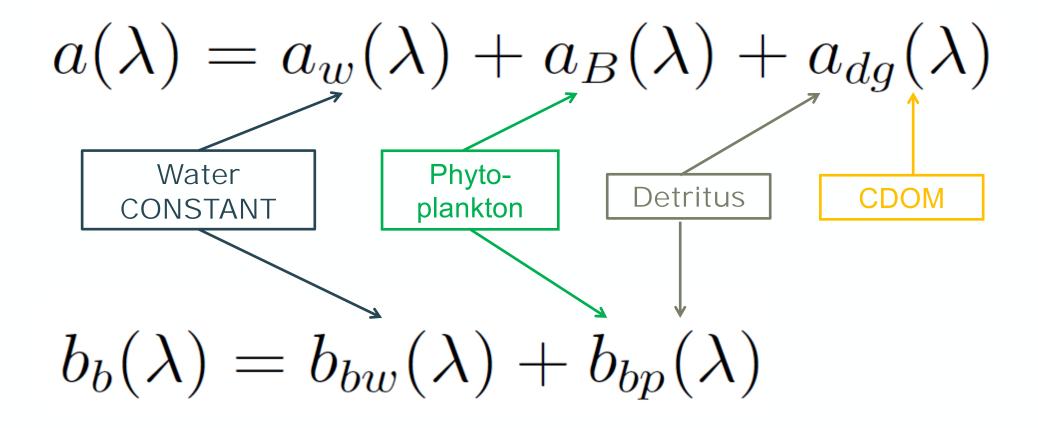
Inherent optical properties (IOPs) The optical properties of the water and its constituents independent of the directional distribution of the light field in the sea (e.g., absorption, backscattering, beam attenuation).

Inherent Optical Properties

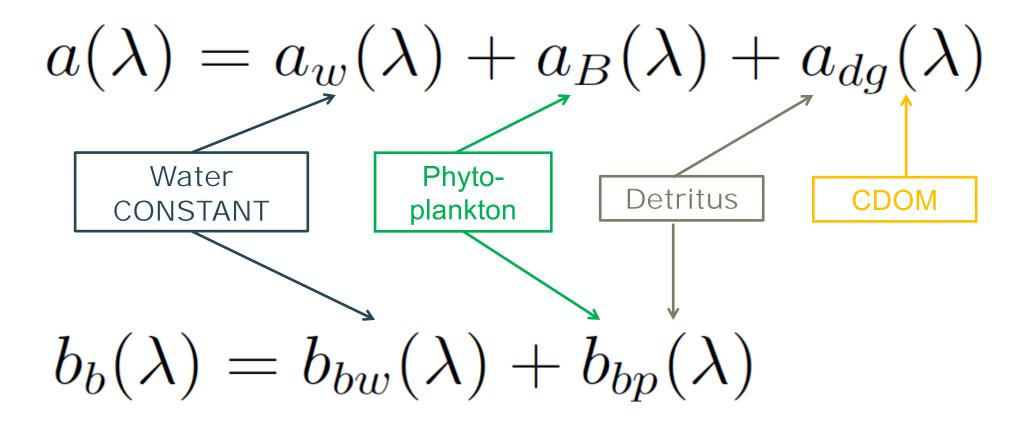
Specific inherent optical properties: the individual scattering and absorption components per unit concentration $(a_B^*(\lambda) = a_B(\lambda)/B)$

Bulk inherent optical properties: the individual scattering and absorption components per unit concentration multiplied by concentration ($a_B(\lambda) = a_B^*(\lambda)B$). Will vary with variations in specific IOPs and concentration.





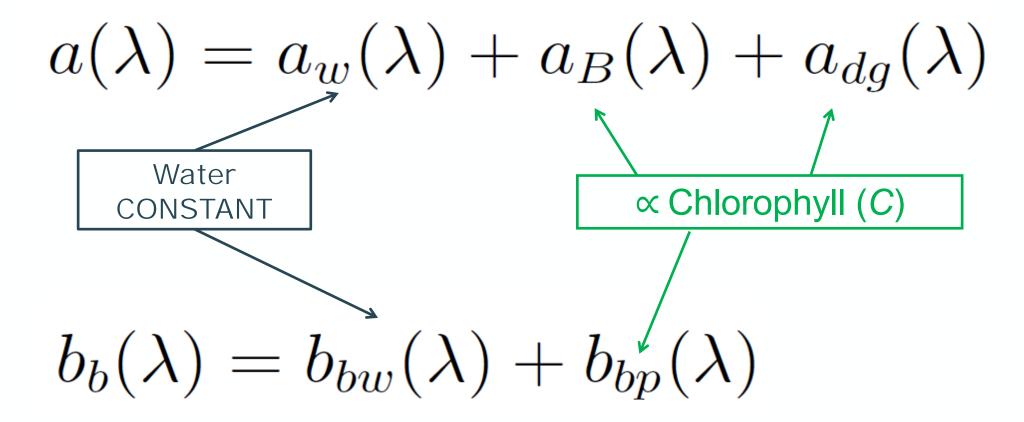




Case-2: Phytoplankton biomass does not covary with detritus and CDOM Case-1: Phytoplankton biomass covaries with detritus and CDOM. IOPs can be tied with reasonable confidence to the chlorophyll concentration (*C*)

Morel and Prieur (1977) Limnol. Oceanogr.





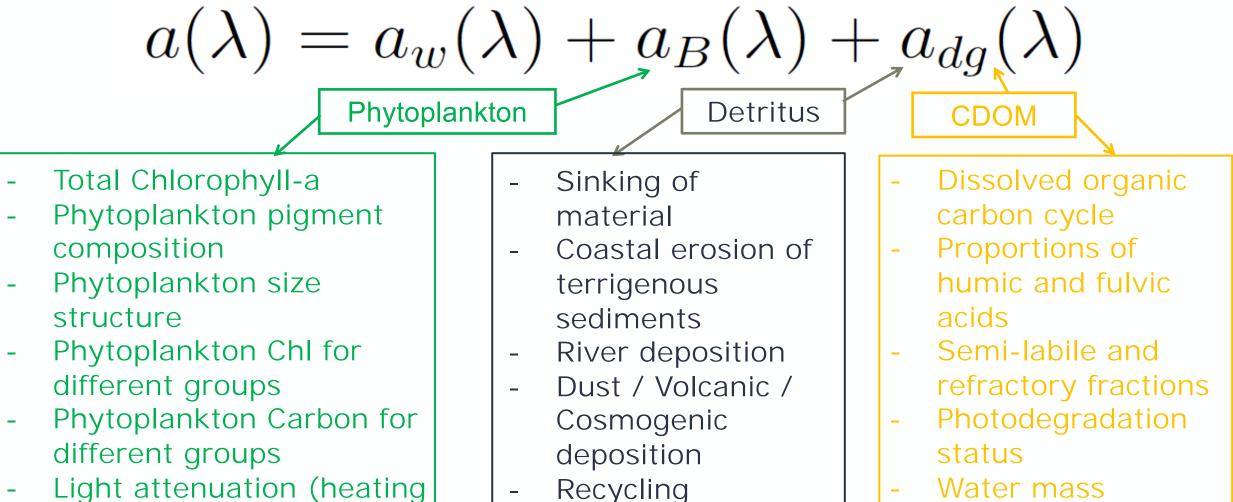
Case-2: Phytoplankton biomass does not covary with detritus and CDOM Case-1: Phytoplankton biomass covaries with detritus and CDOM. IOPs can be tied with reasonable confidence to the chlorophyll (*C*)

Morel and Prieur (1977) Limnol. Oceanogr.



change

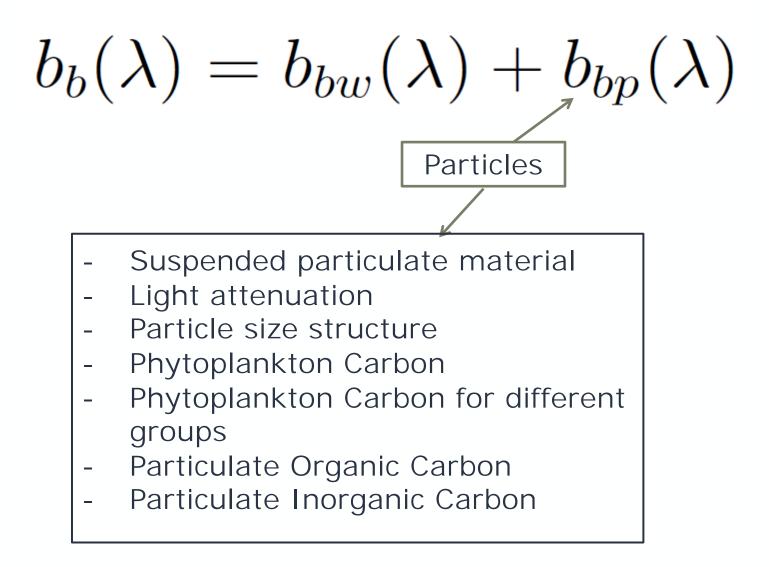
Light attenuation



 Light attenuation (heating of ocean and primary production)

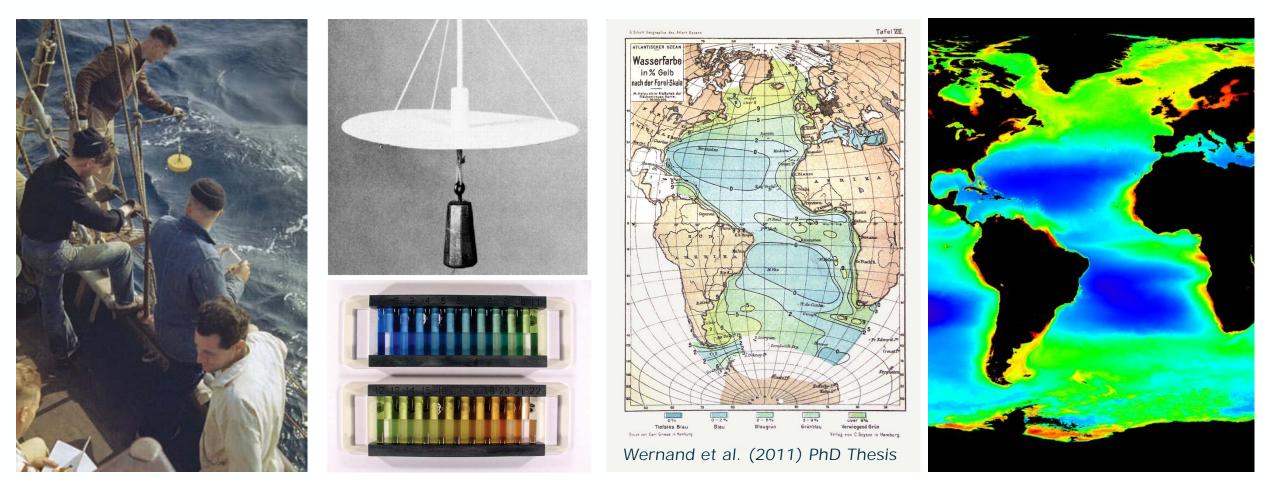
- Light attenuation





4) In situ measurements of ocean colour



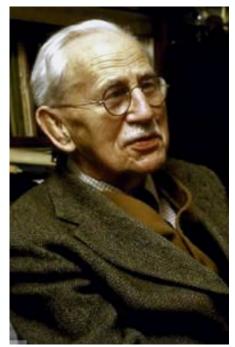




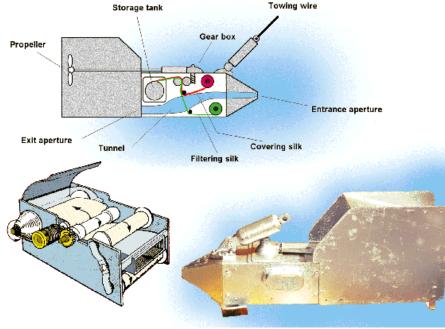
Siegel & Franz (2010) <u>https://doi.org/10.1038/466569a</u> Seafarers et al. (2017) <u>https://doi.org/10.1371/journal.pone.0186092</u> Brewin et al (2019) <u>https://doi.org/10.3390/s19040936</u> Garaba et al (2015) <u>https://doi.org/10.3390/ijerph121215044</u> Ceccaroni et al. (2020) <u>https://doi.org/10.1371/journal.pone.0230084</u>

4) In situ measurements of ocean colour



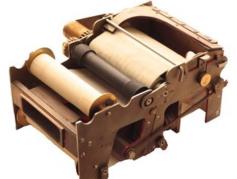


Sir Alister Clavering Hardy (1896- 1985)

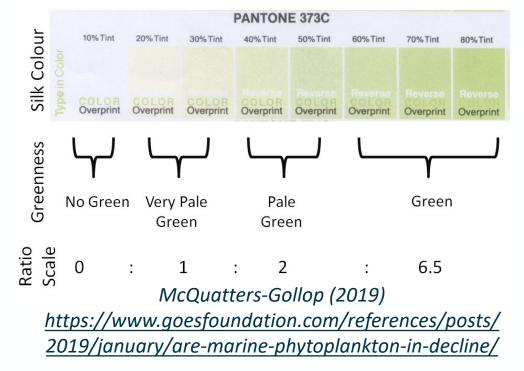


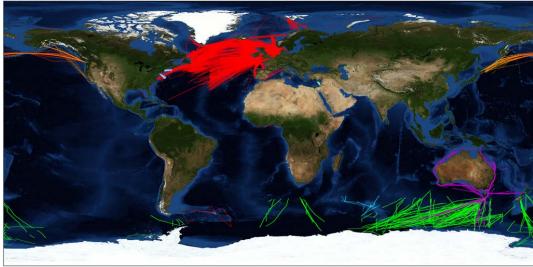
Richardson et al. (2006) https://doi.org/10.1016/j.pocean.2005.09.011





https://www.cprsur vey.org/services/th e-continuousplankton-recorder/

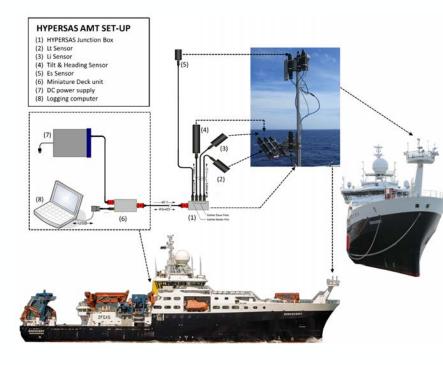


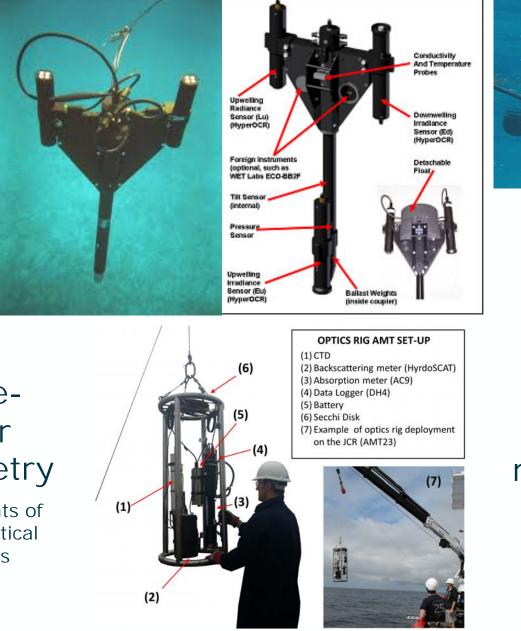


4) In situ measurements of ocean colour











In water radiometry

Measurements of apparent optical properties

In water measurements of inherent optical properties



Abovewater radiometry

Measurements of apparent optical properties

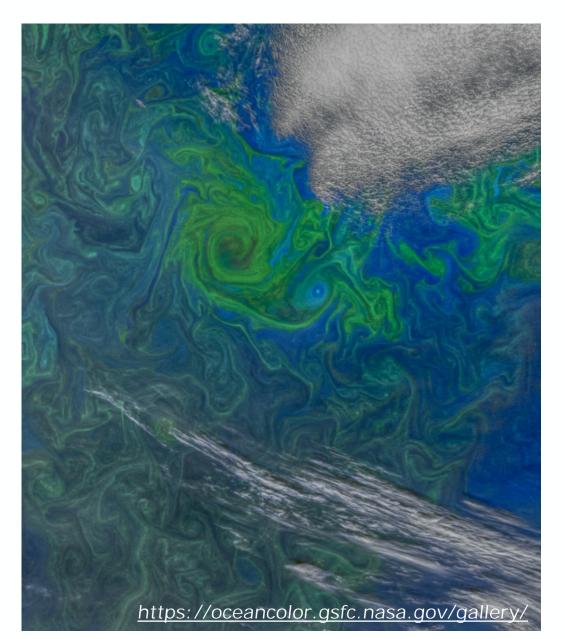




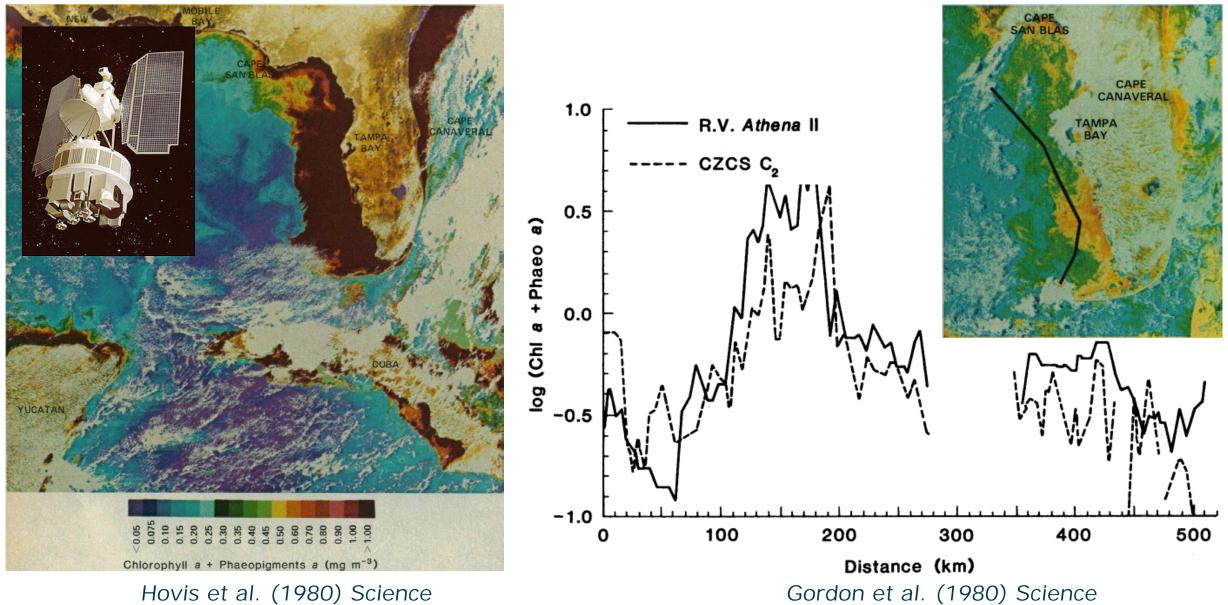
Considered among the biggest achievements in biological oceanography in the 20th century (Barber & Hilting 2000)

Revealed:

- Oceanography's chronic problem of undersampling;
- (2) Dominance of mesoscale physical processes in determining the spatial distribution of phytoplankton;
- (3) Effect of topography on biomass;
- (4) Complexity of the seasonal progression of phytoplankton blooms
- (5) Magnitude of interannual variability.



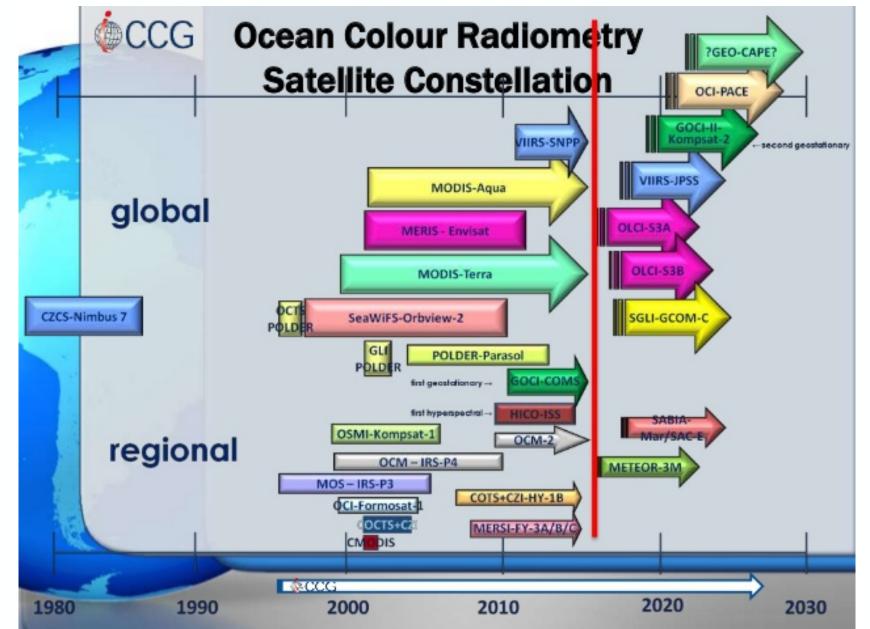




http://doi.org/10.1126/science.210.4465.60

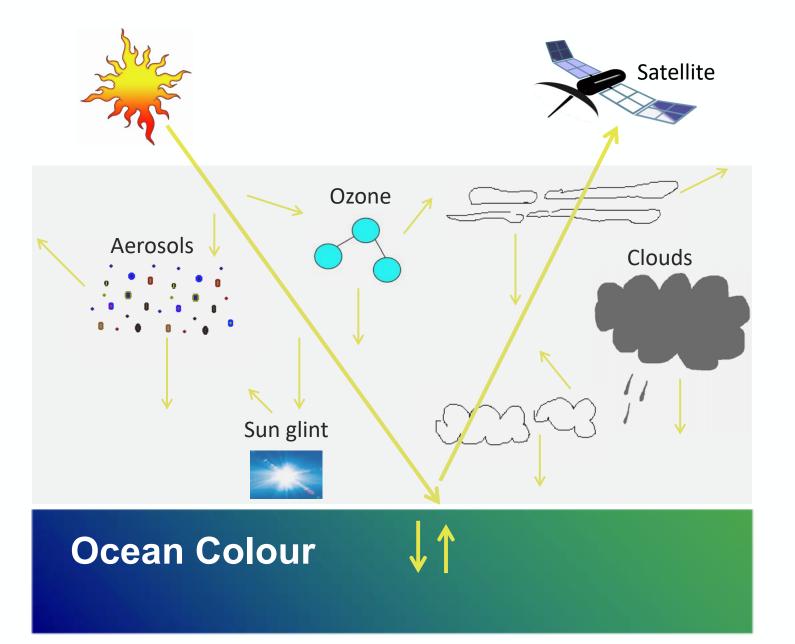
http://doi.org/10.1126/science.210.4465.63





http://chlo4msfd.azti.es/products-and-services-for-satellite-chlorophyll-a-data/



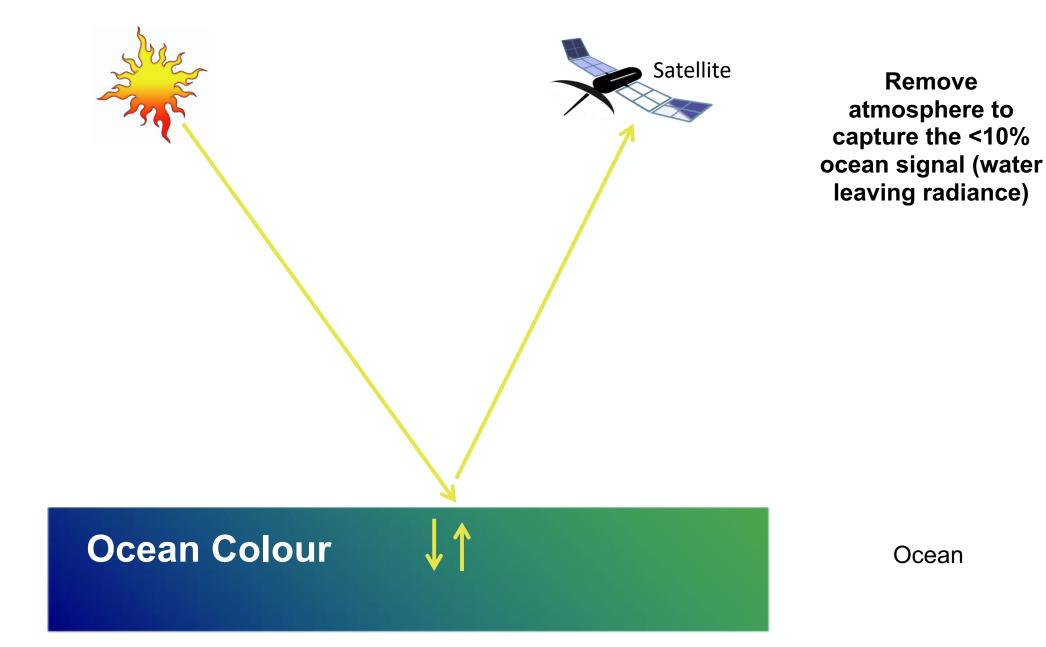


Atmosphere responsible for >90% of the satellite signal, the ocean only <10%

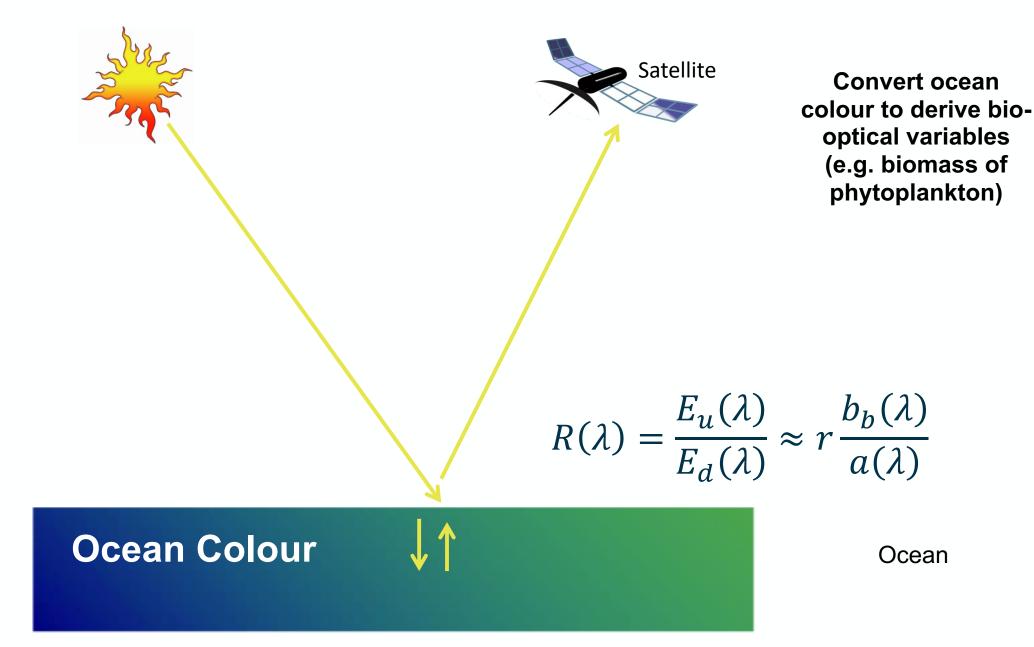
Atmosphere



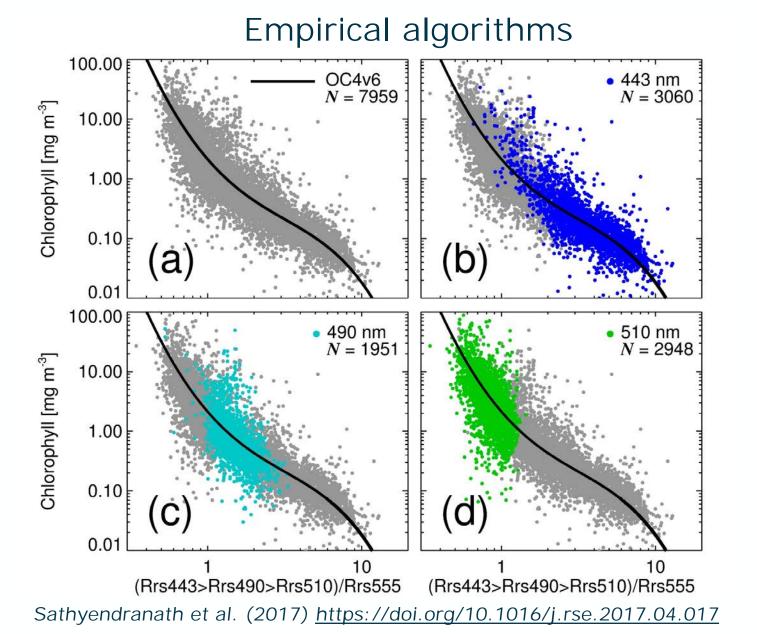










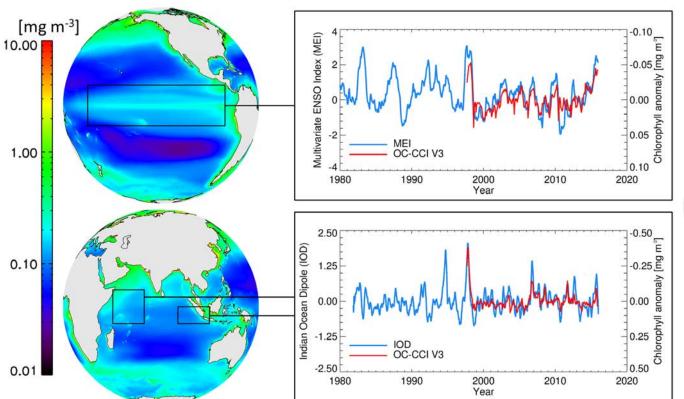


Semi-analytical algorithms

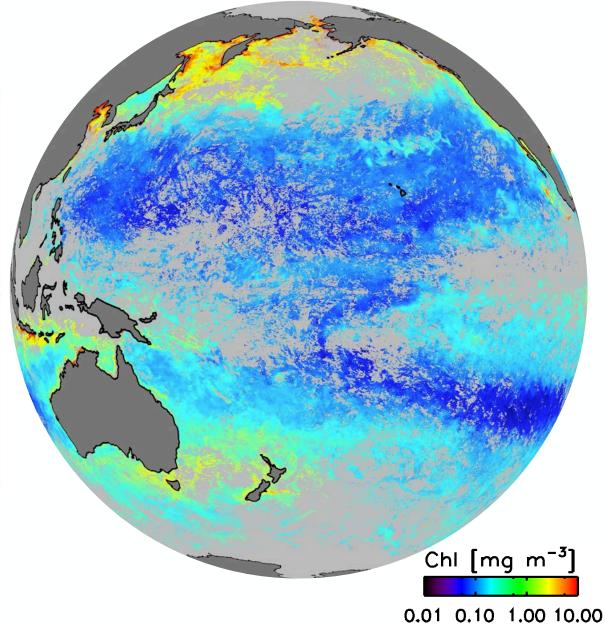
$$R(\lambda) = \frac{E_u(\lambda)}{E_d(\lambda)} \approx r \frac{b_b(\lambda)}{a(\lambda)}$$
$$a(\lambda) = a_w(\lambda) + a_B(\lambda) + a_g(\lambda) + a_d(\lambda)$$
$$b_b(\lambda) = b_{bw}(\lambda) + b_{bp}(\lambda)$$
$$a_B(\lambda) = a_B^*(\lambda) + B$$
$$b_{bp}(\lambda) = b_{bp}^*(\lambda) + SPM$$
$$a_g(\lambda) = a_g^*(\lambda) + CDOM$$

IOCCG (2006) <u>https://ioccg.org/wp-</u> <u>content/uploads/2015/10/ioccg-report-05.pdf</u>





von Schuckmann et al. (2016) https://doi.org/10.1080/1755876X.2016.1273446



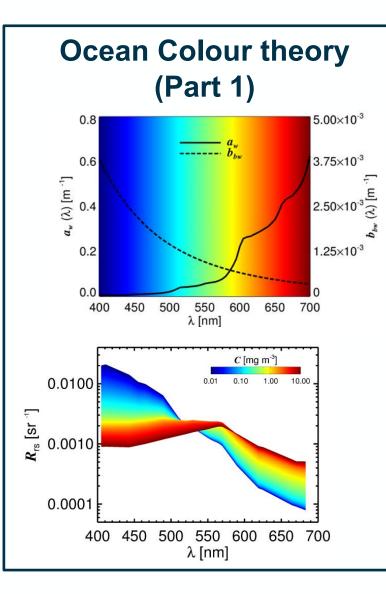
Part 1: Outline



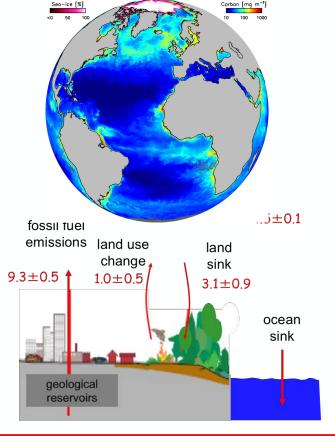
- 1) What is ocean colour?
- 2) History of ocean colour
- 3) Optical properties of water and its constituents
- 4) In situ measurements of ocean colour
- 5) Satellite remote-sensing of ocean colour

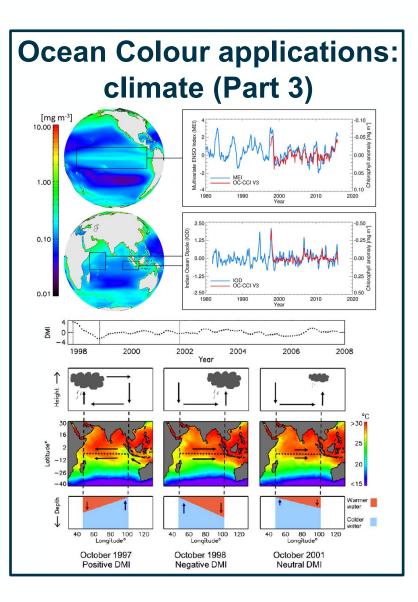
Lecture plan





Ocean Colour applications: carbon cycle (Part 2)





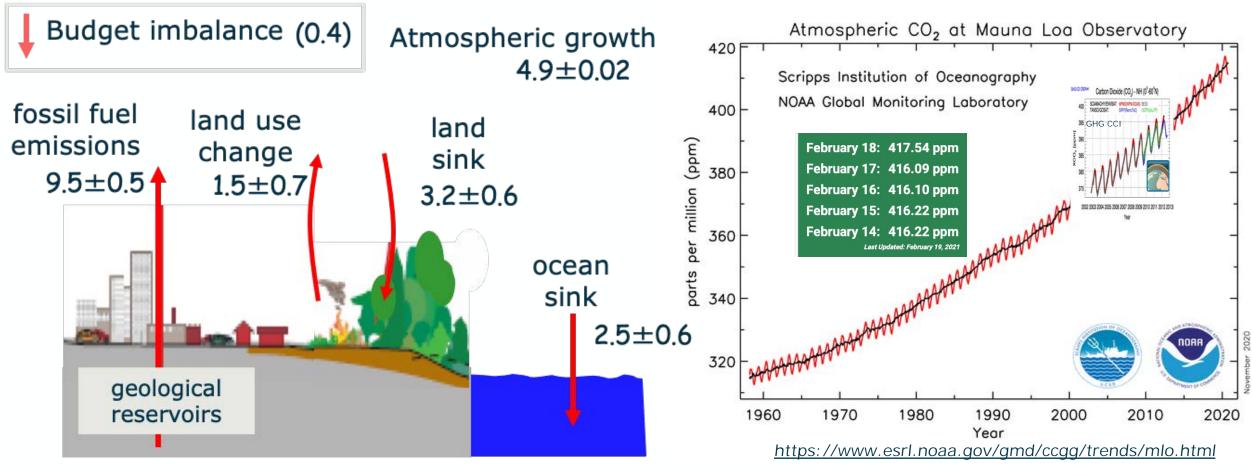
Part 2: Outline



- 1) Earth's Carbon Cycle
- 2) Ocean Carbon Cycle
- 3) Ocean carbon satellite products
- 4) Integration with robotic platforms
- 5) Integration with models

1) Earth's Carbon Cycle

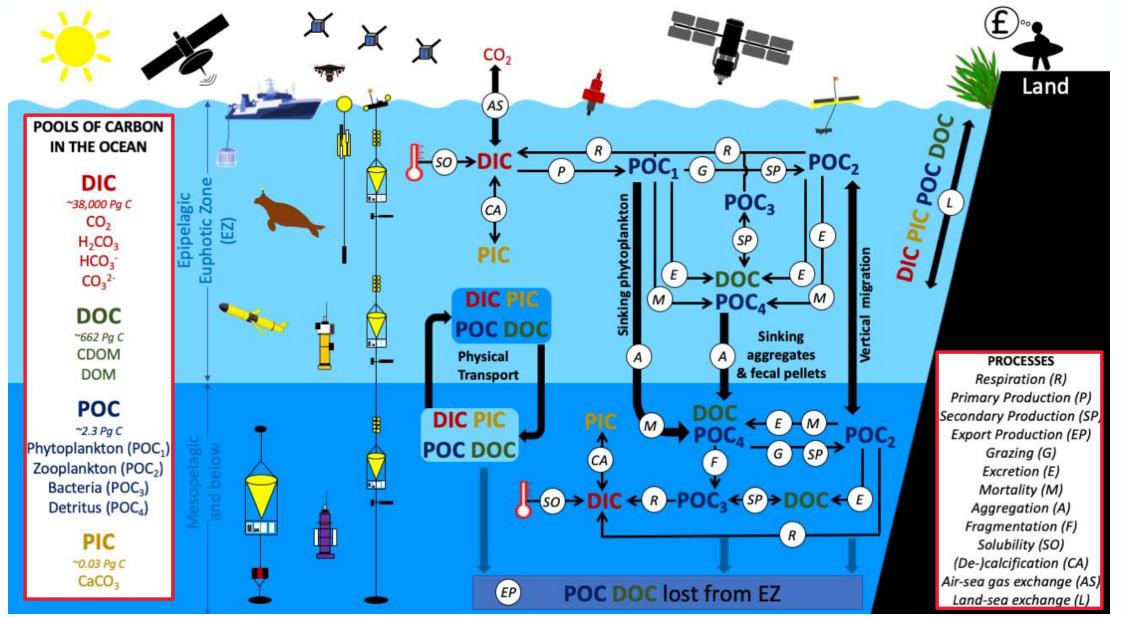




Friedlingstein et al. (2019) Global Carbon Budget (2019) Earth Syst. Sci. Data, 11, 1783–1838, <u>https://doi.org/10.5194/essd-11-1783-2019</u>

2) Ocean Carbon Cycle

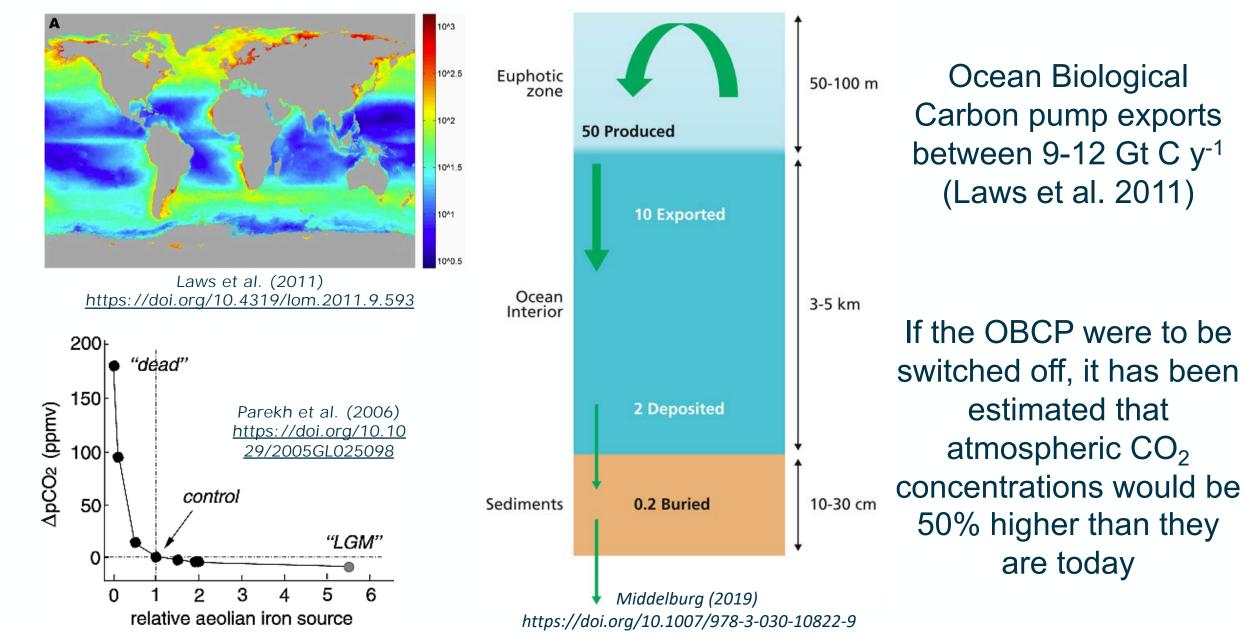




Brewin et al. (Under Review)

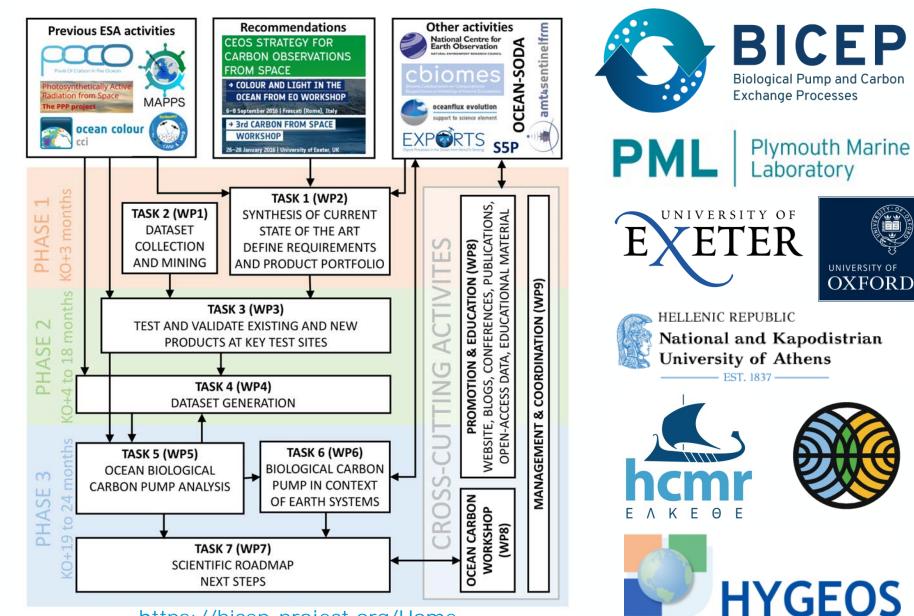
2) Ocean Carbon Cycle





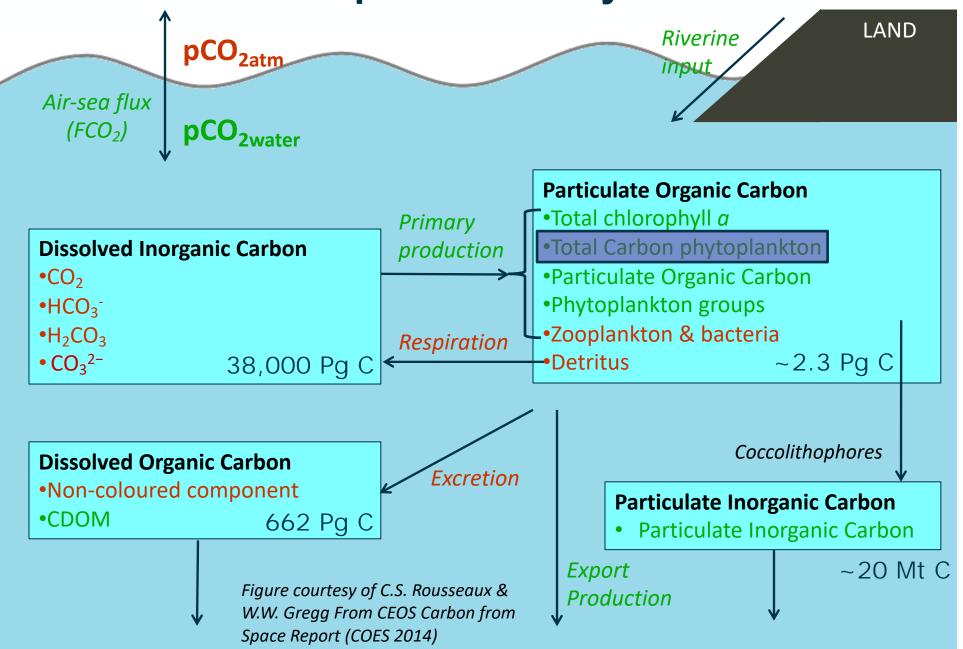
2) Ocean Carbon Cycle



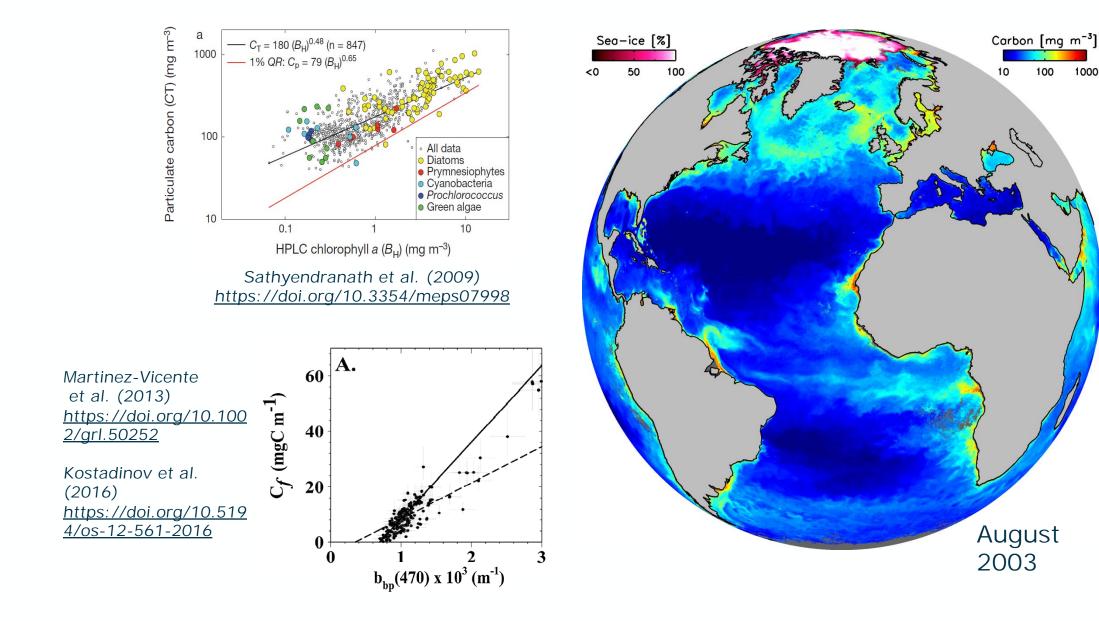


https://bicep-project.org/Home

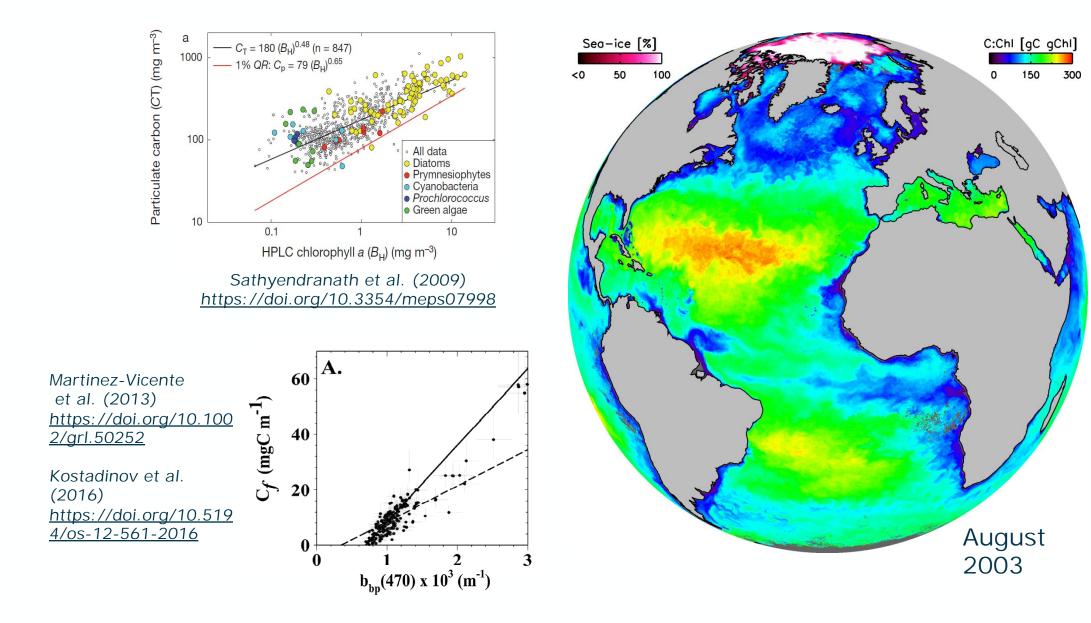






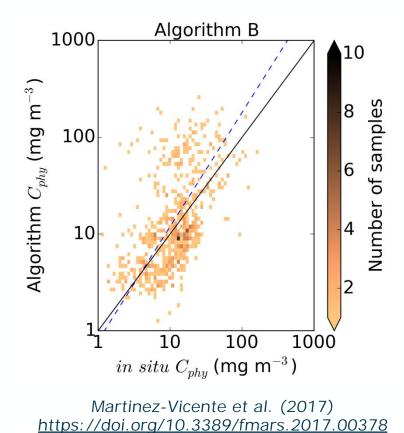


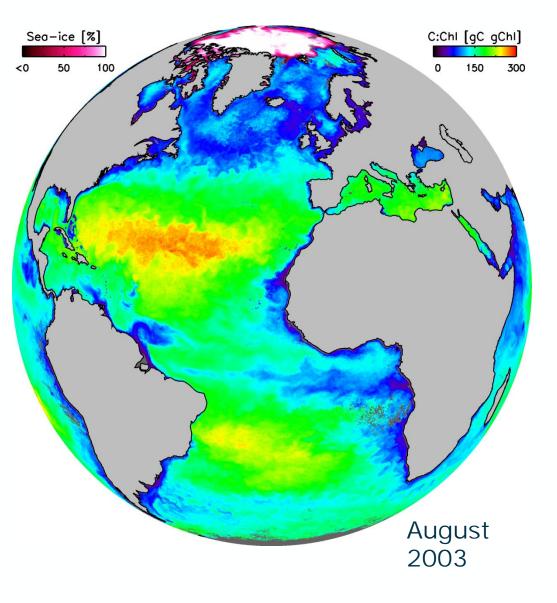






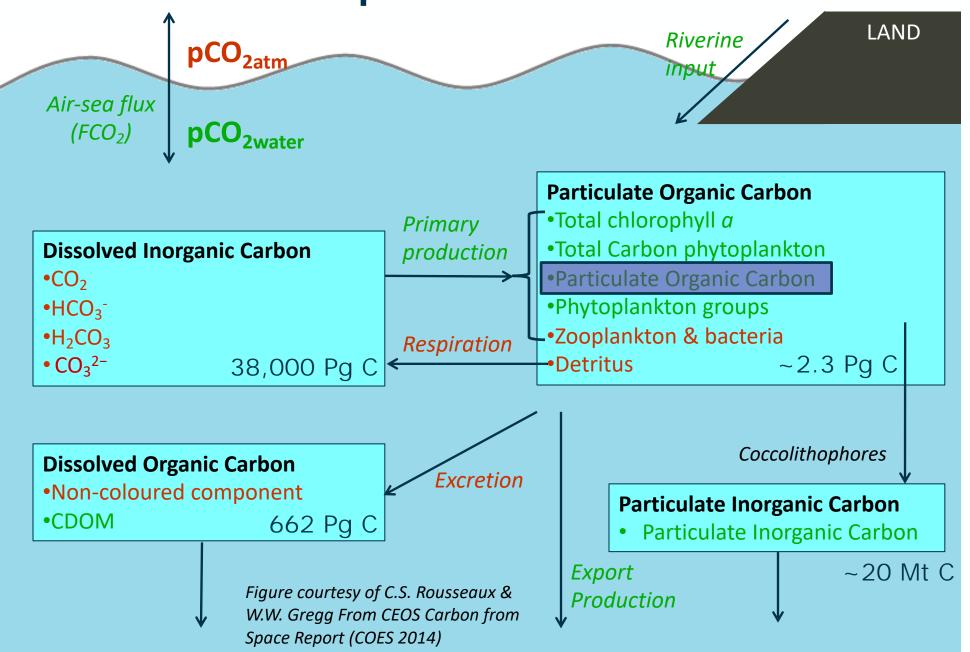
Validation



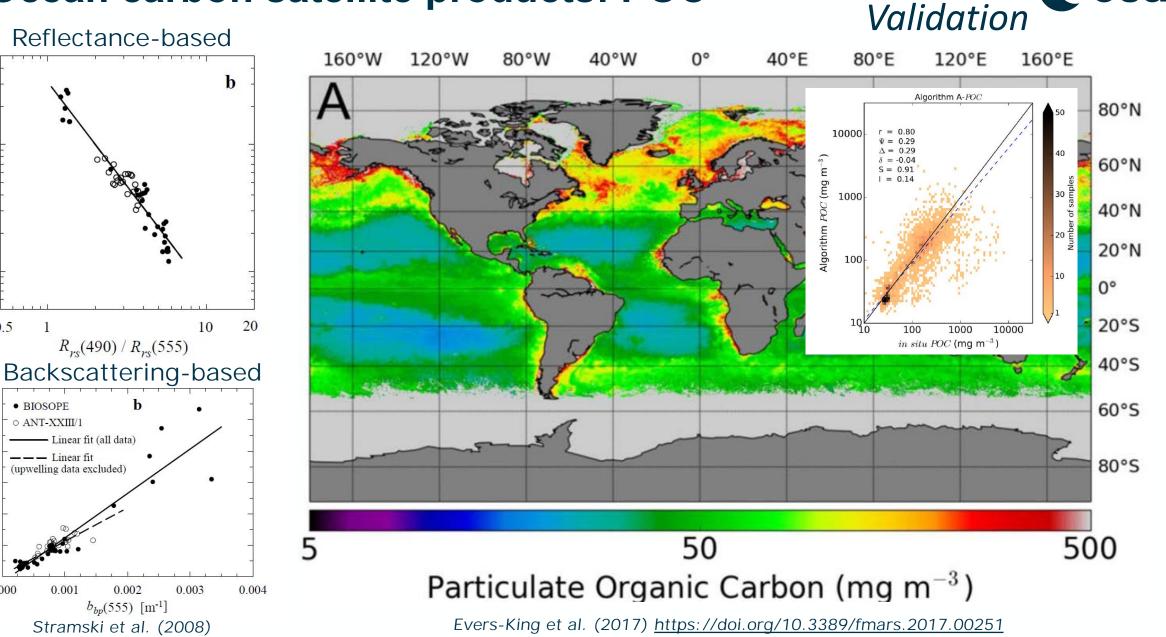


3) Ocean carbon satellite products: POC





3) Ocean carbon satellite products: POC



eesa

https://doi.org/10.5194/bg-5-171-2008

500

100

10

5

300

250

200

150

100

50

0.000

POC [mg m⁻³]

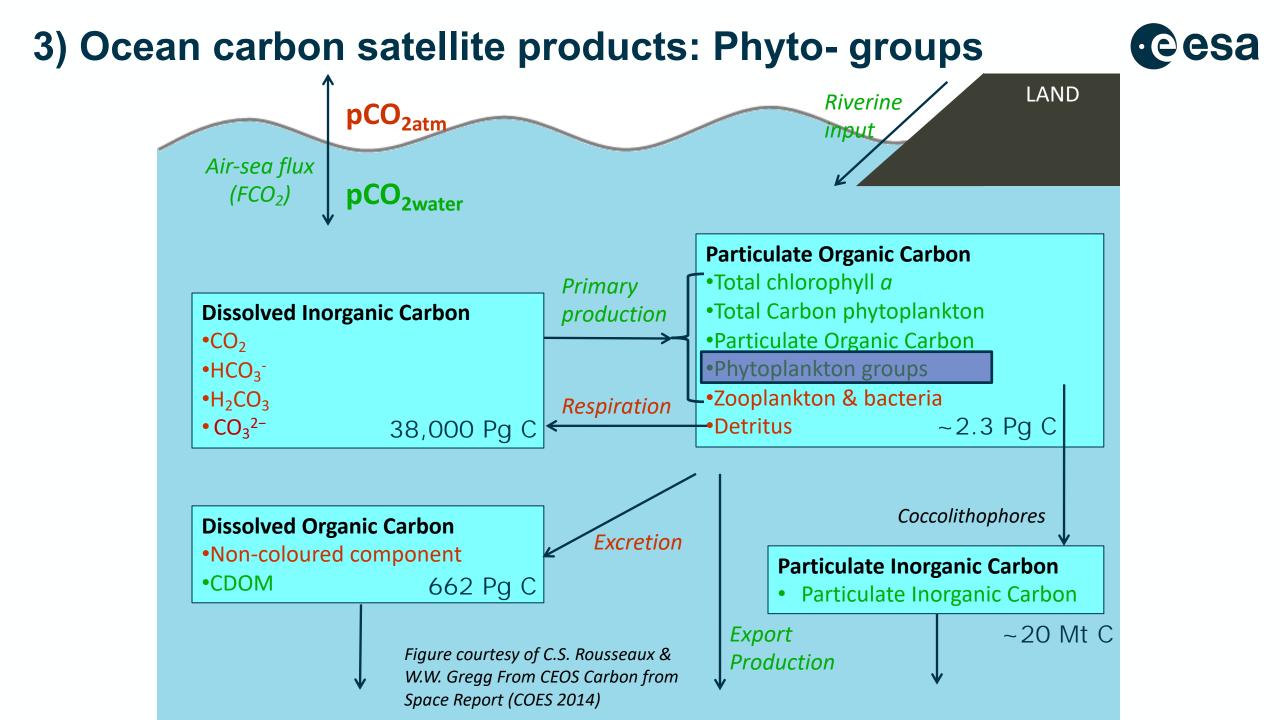
0.5

BIOSOPE

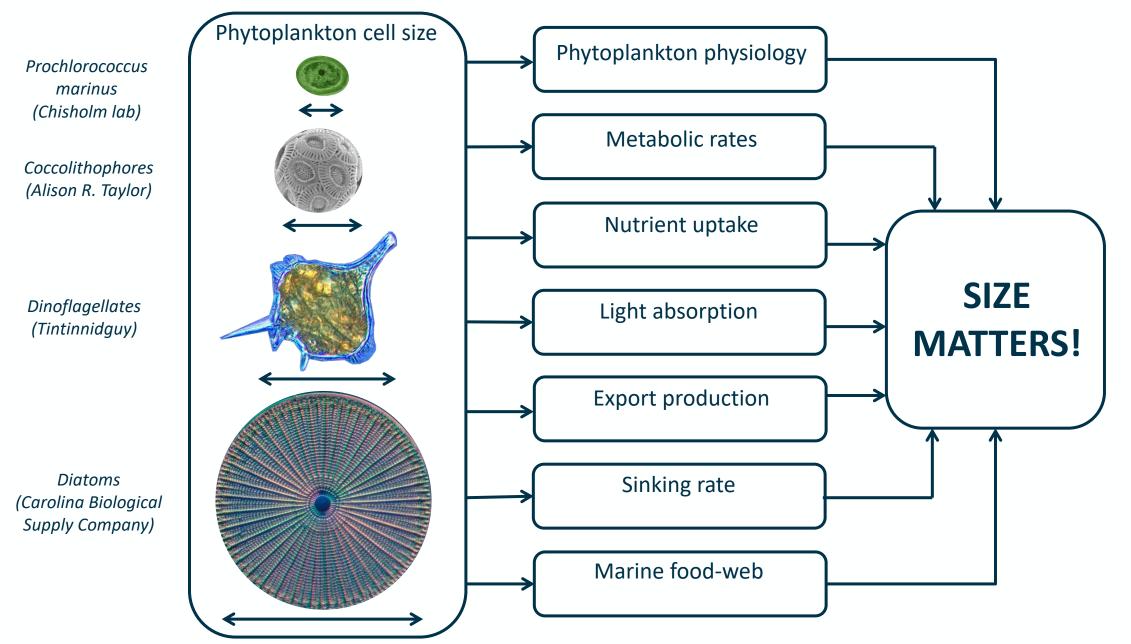
O ANT-XXIII/1

0.001

POC [mg m⁻³]

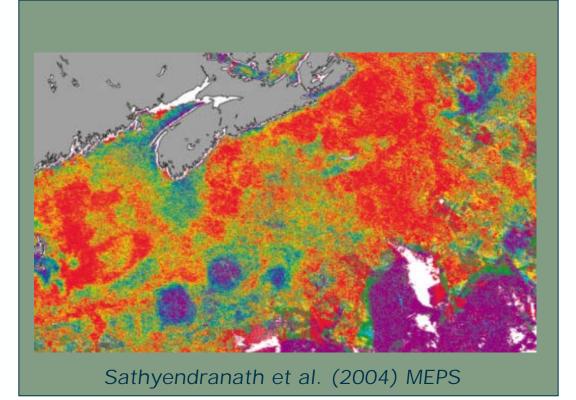




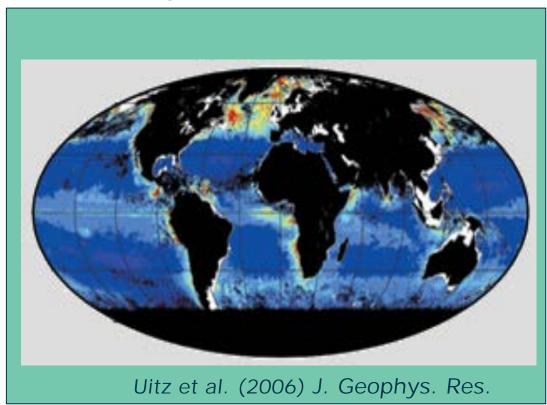




Detection-based

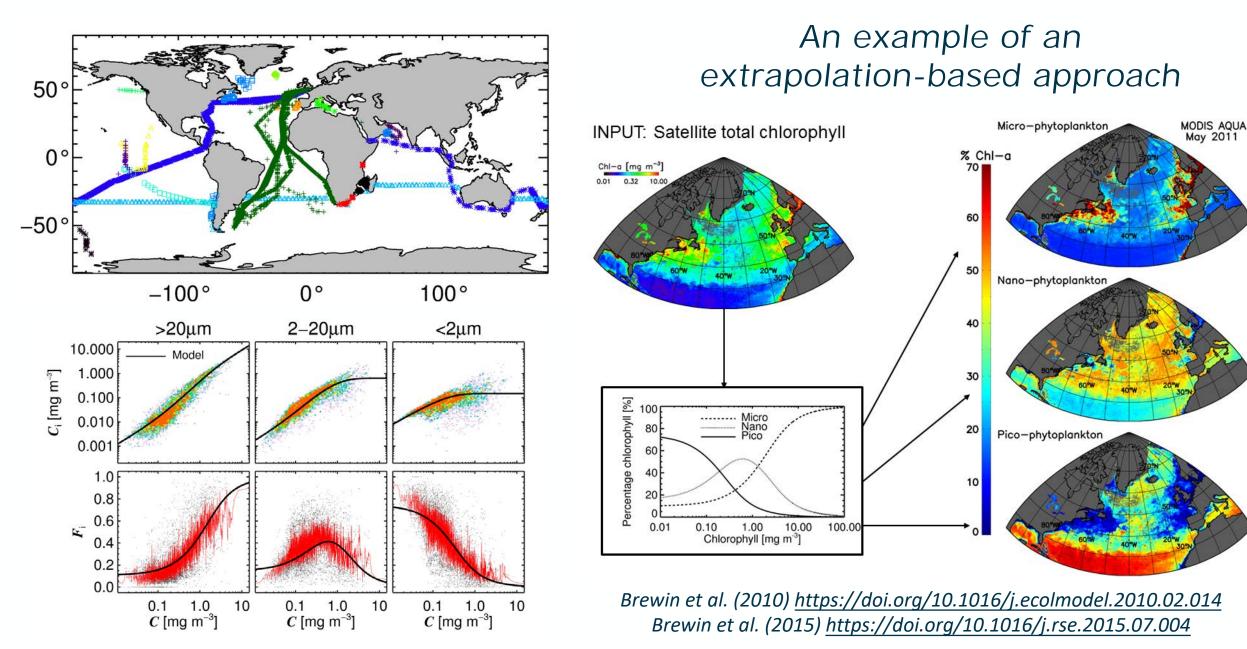


Extrapolation-based

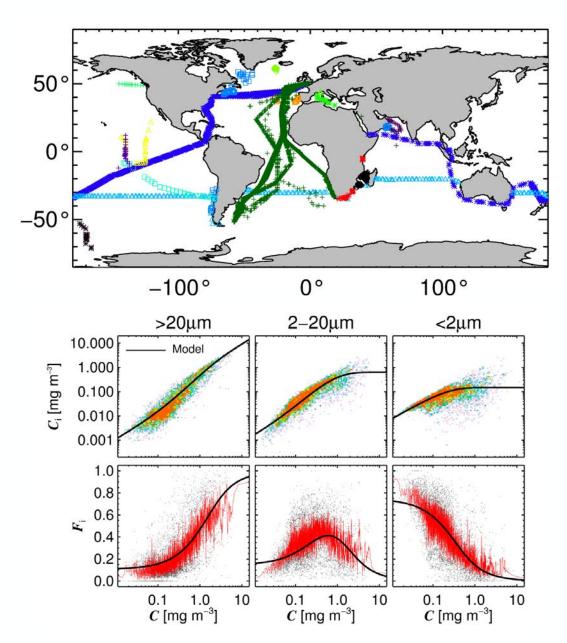


IOCCG (2014) Report Number 15 <u>http://ioccg.org/wp-content/uploads/2018/09/ioccg_report_15_2014.pdf</u> Mouw et al. (2017) <u>https://doi.org/10.3389/fmars.2017.00041</u>



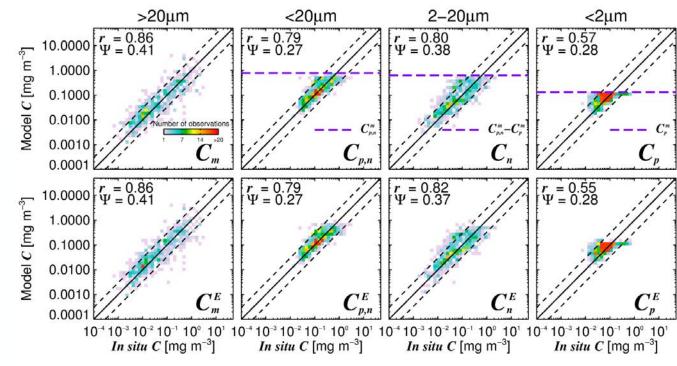






An example of an extrapolation-based approach

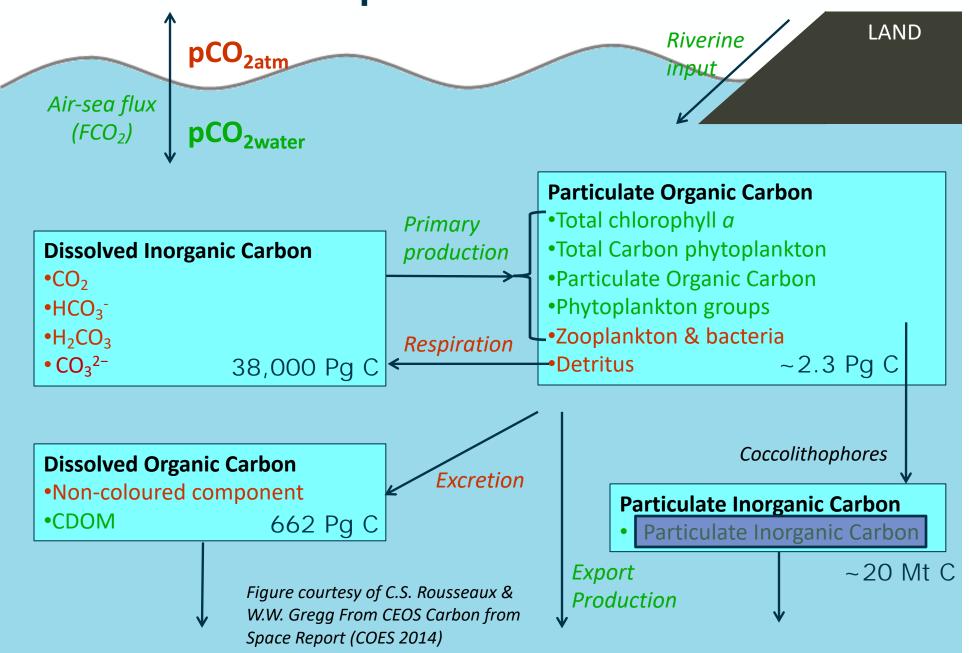
Validation



Brewin et al. (2010) <u>https://doi.org/10.1016/j.ecolmodel.2010.02.014</u> Brewin et al. (2015) <u>https://doi.org/10.1016/j.rse.2015.07.004</u>

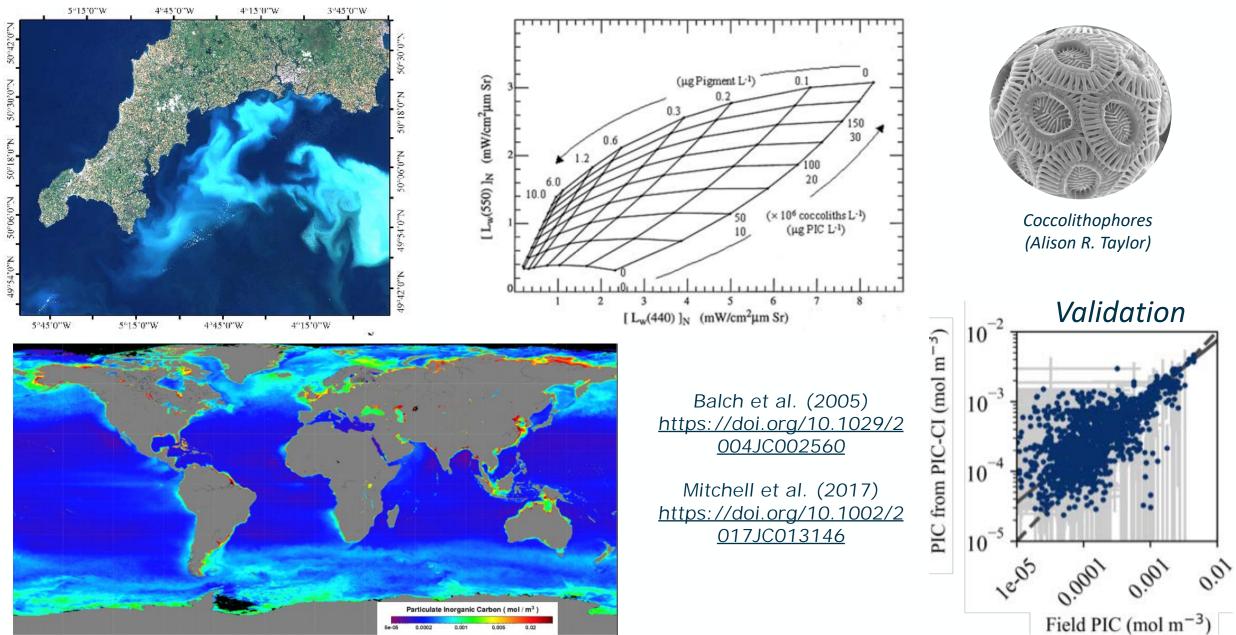
3) Ocean carbon satellite products: PIC

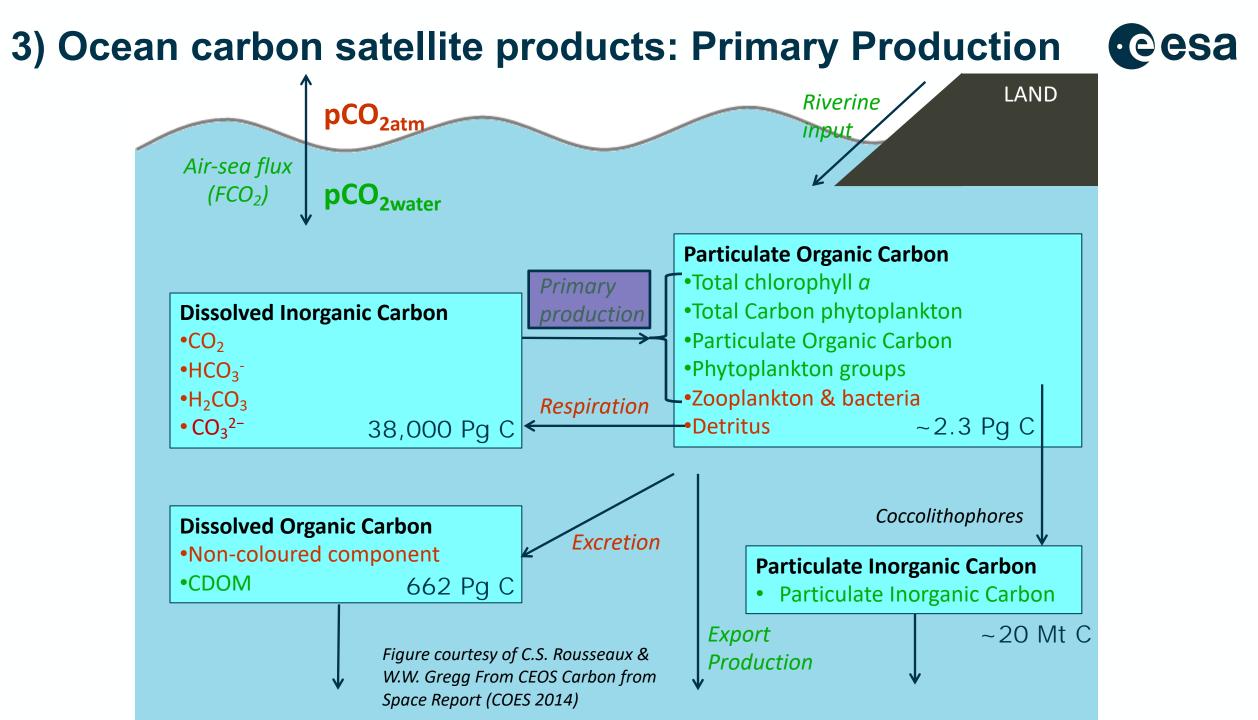




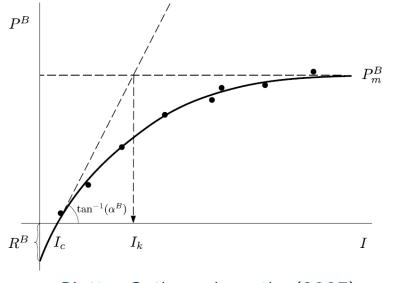
3) Ocean carbon satellite products: PIC



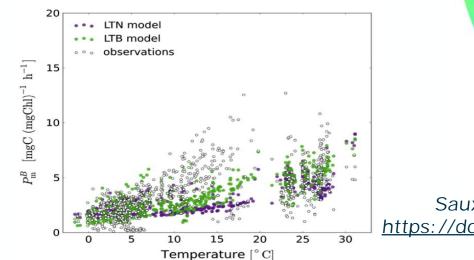


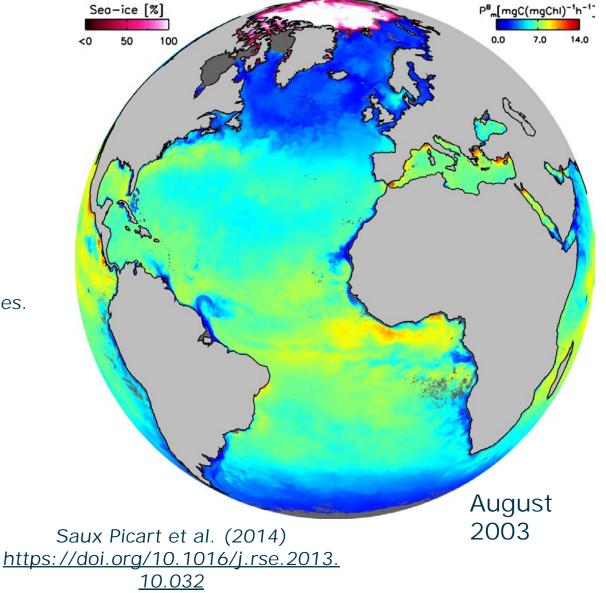


3) Ocean carbon satellite products: Primary Production @esa

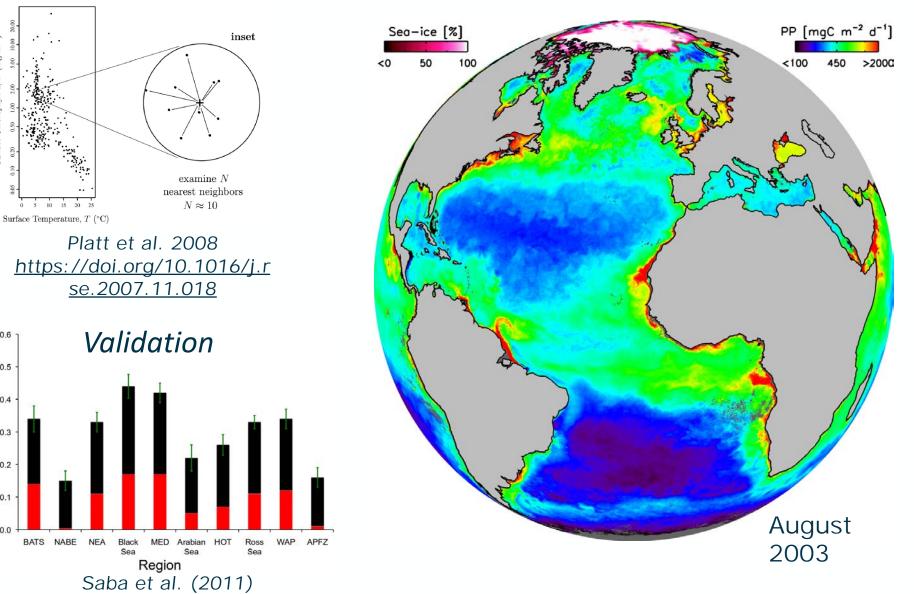


Platt & Sathyendranath (2007) Modelling Primary production article Series.





3) Ocean carbon satellite products: Primary Production cesa



https://doi.org/10.5194/bg-8-489-2011

In Situ data

 $(\mathrm{mg}\ \mathrm{m}^{-3})$

0.6

0.5

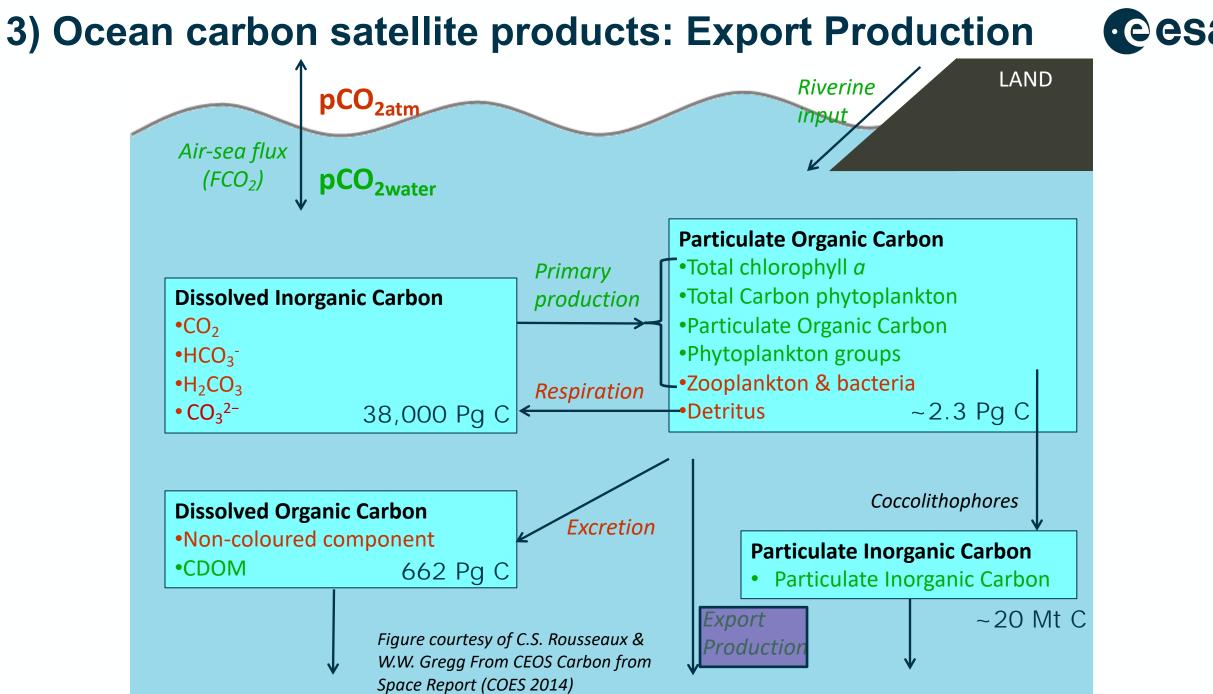
0.4

0.2

0.1

0.0

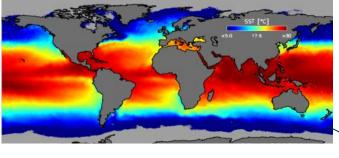
RMSD 80.3



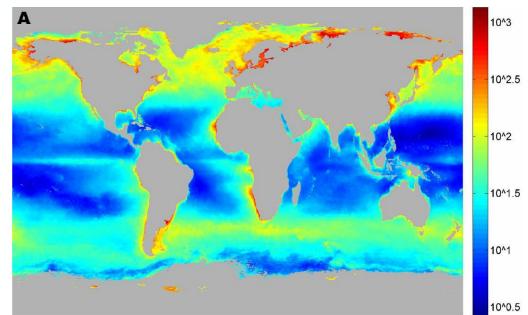


3) Ocean carbon satellite products: Export Production

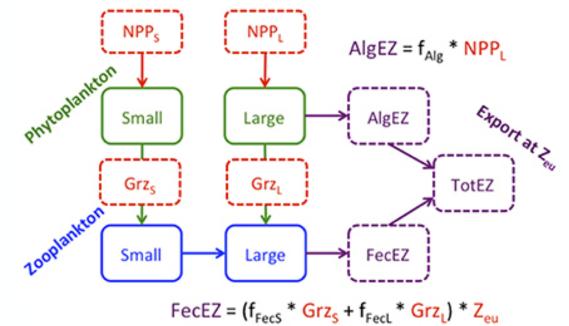


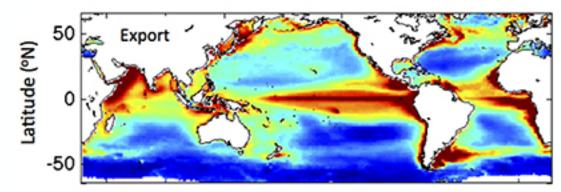


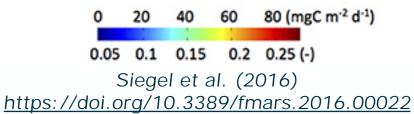
 $ef = 0.04756 \left(0.78 - \frac{0.43T}{30} \right) tp^{0.307}$



Laws et al. (2011) https://doi.org/10.4319/lom.2011.9.593

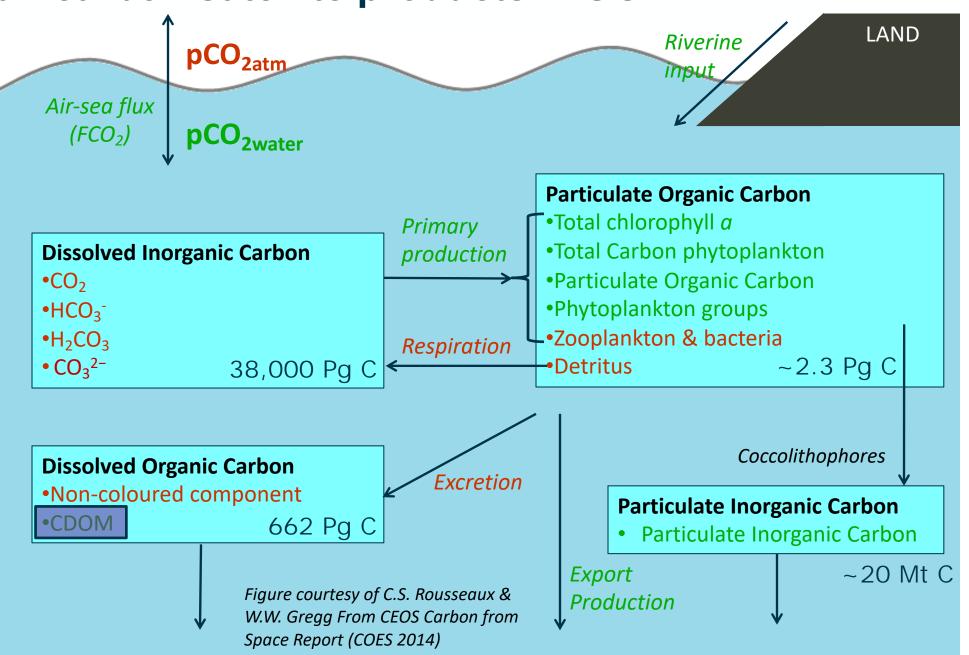






3) Ocean carbon satellite products: DOC





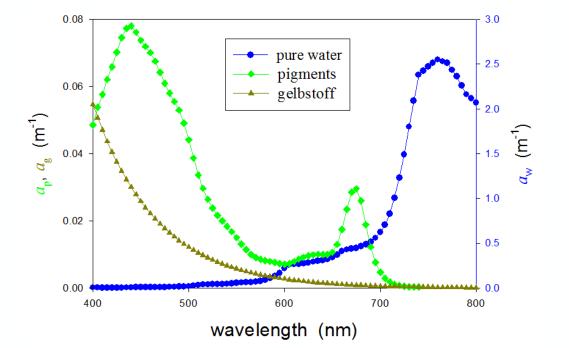
3) Ocean carbon satellite products: DOC

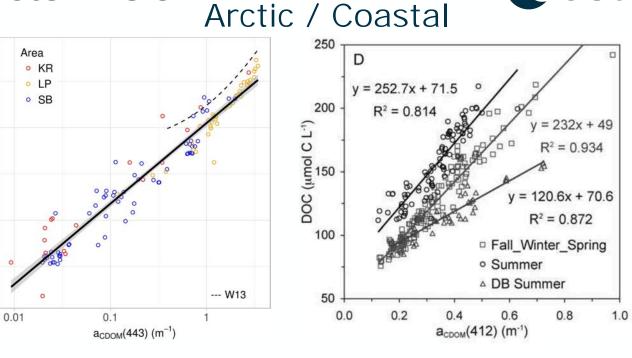
10^{2.5}

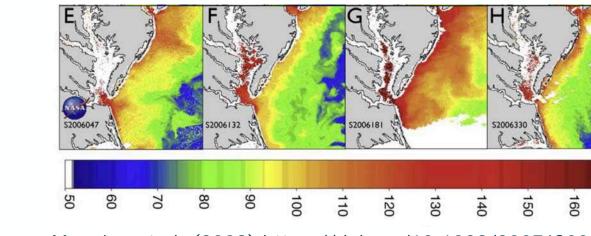
 10^{2}

DOC (µM)

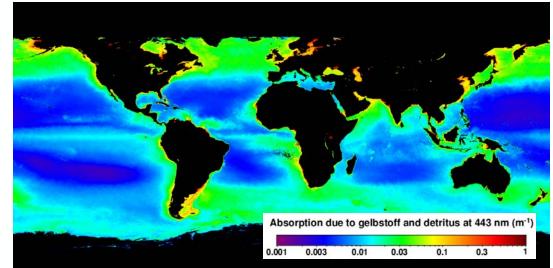






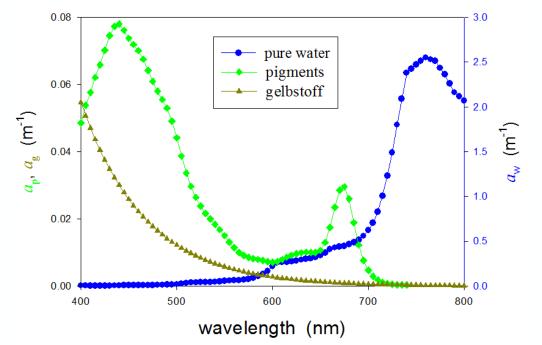


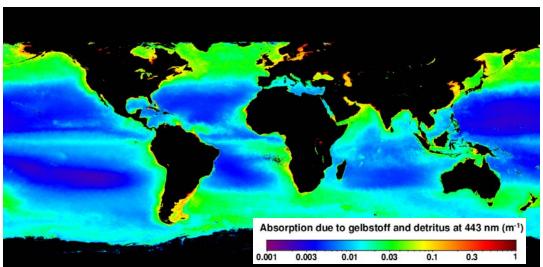
Mannino et al. (2008) <u>https://doi.org/10.1029/2007JC004493</u> Matsuoka et al. (2017) <u>https://doi.org/10.1016/j.rse.2017.08.009</u>

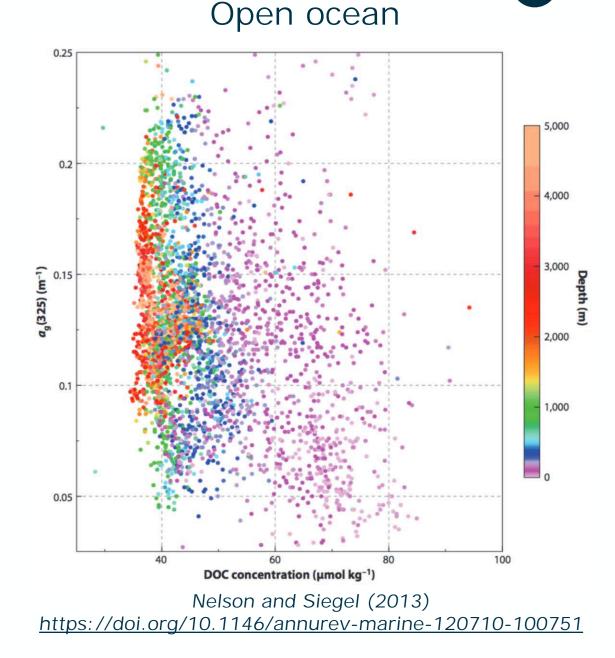


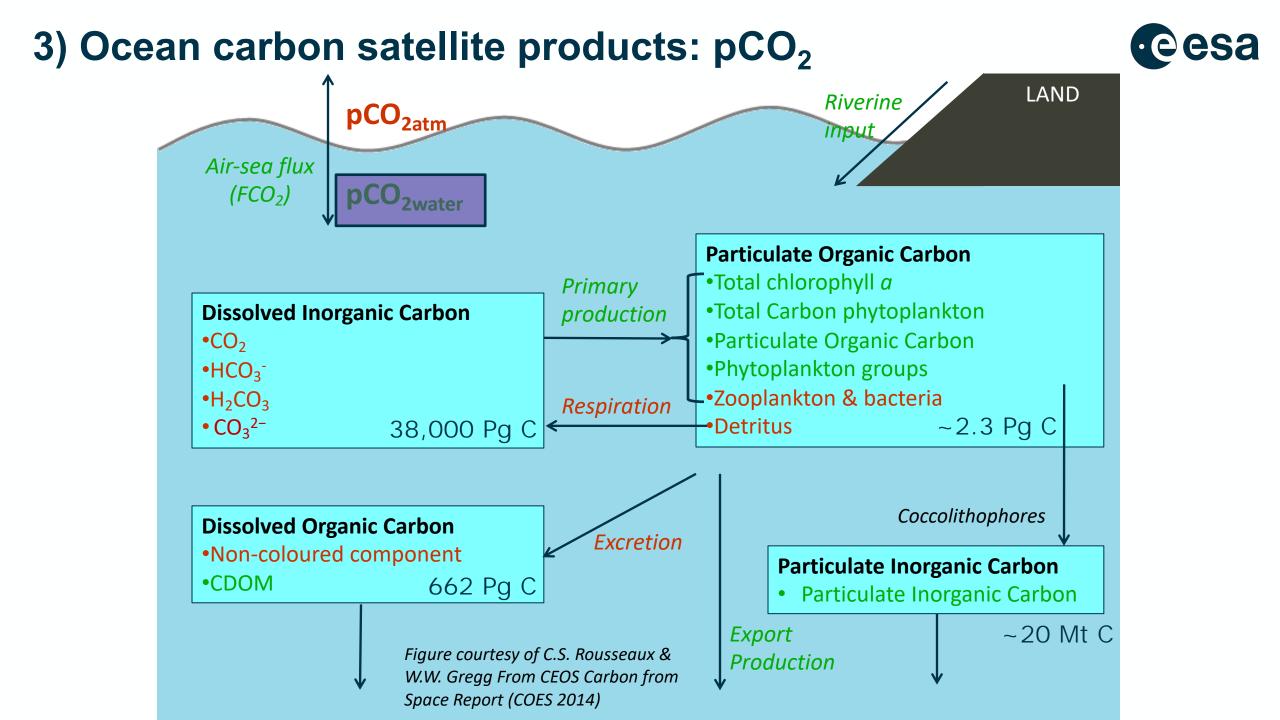
3) Ocean carbon satellite products: DOC







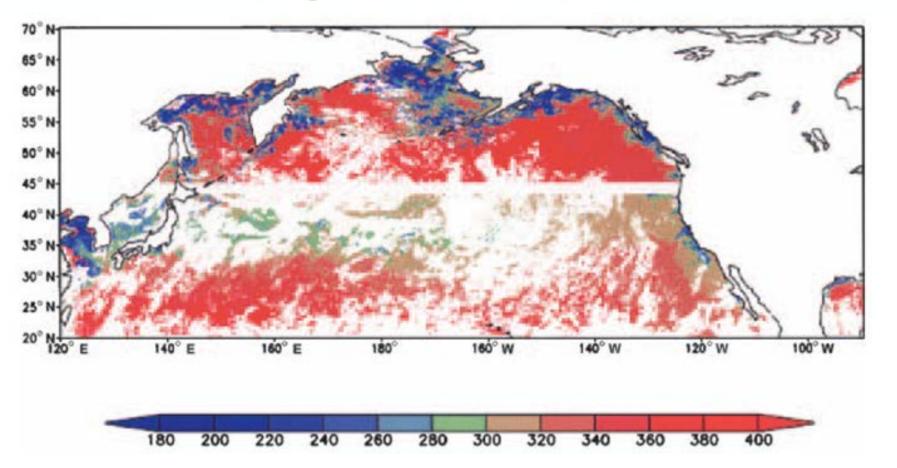




3) Ocean carbon satellite products: pCO₂



 $pCO_2 = AT + BT^2 + CChla + DChla^2 + E$



Ono et al. (2004) <u>https://doi.org/1</u> <u>0.1080/0143116</u> <u>0310001657515</u>

Figure 4. The basin-scale pCO₂ field of the North Pacific in May 1997 obtained from the ADEOS/OCTS monthly-average SST and monthly-maximum Chla fields by using the multiple regression equations.

4) Integration with robotic platforms

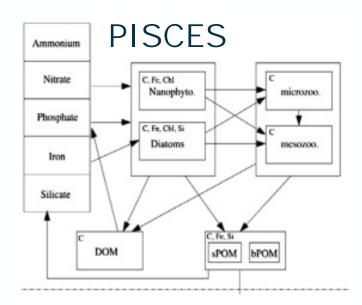


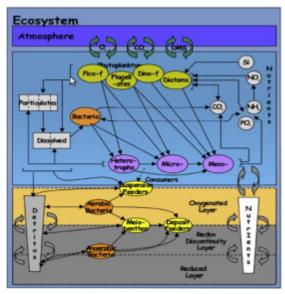


Mixed-layer pump Dall'Olmo et al. (2016) <u>https://doi.org/10.1038/ngeo2818</u>

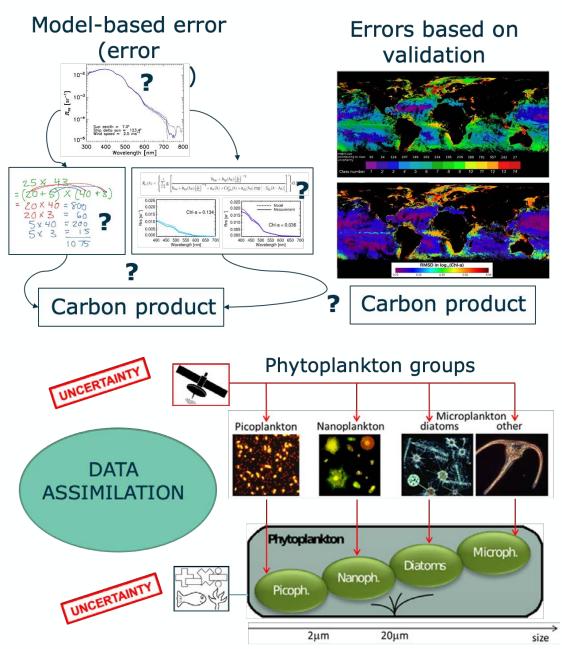
5) Integration with models







ERSEM



Improve understanding of pools and fluxes not seen from space

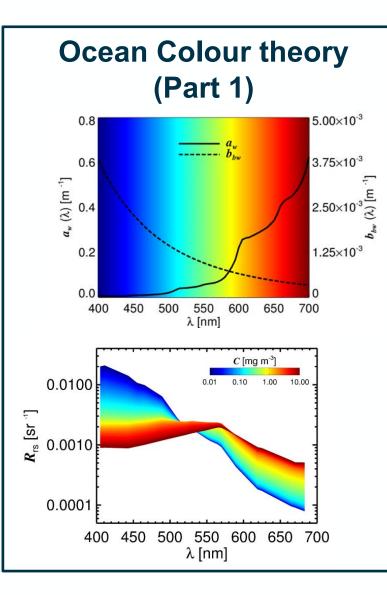
Part 2: Outline



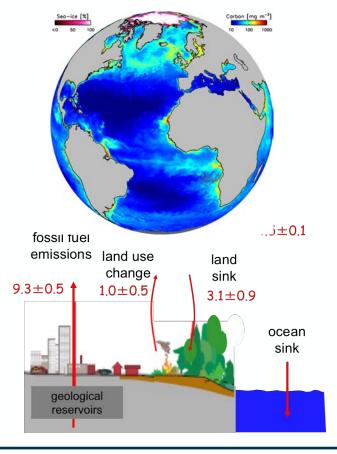
- 1) Earth's Carbon Cycle
- 2) Ocean Carbon Cycle
- 3) Ocean carbon satellite products
- 4) Integration with robotic platforms
- 5) Integration with models

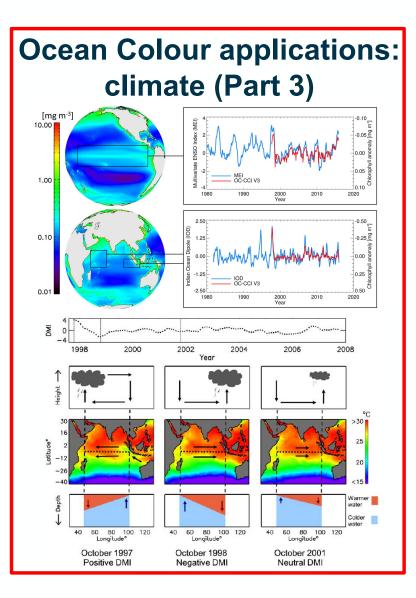
Lecture plan





Ocean Colour applications: carbon cycle (Part 2)





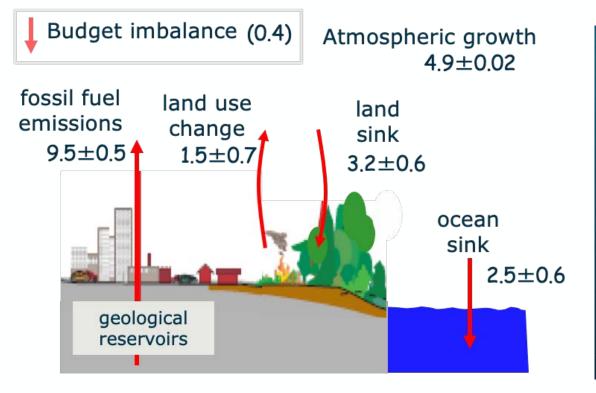
Part 3: Outline



- 1) The Greenhouse effect
- 2) Essential Climate Variables
- 3) Ocean colour datasets requirements for climate
- 4) Responses of the marine ecosystem to climate change
 - Total phytoplankton biomass
 - Phytoplankton community structure
 - Phytoplankton phenology
 - Other responses
- 5) Other applications

1) The Greenhouse Effect





Friedlingstein et al. (2019) Global Carbon Budget (2019) Earth Syst. Sci. Data, 11, 1783–1838, <u>https://doi.org/10.5194/essd-11-1783-2019</u>



Credit: NASA/JPL-Caltech https://climatekids.nasa.gov/climate-change-meaning/

2) Essential Climate Variables



Established in 1992, the Global Climate Observing System (GCOS) identified 50 Essential Climate Variables (ECVs) to support the work of the intergovernmental expert group on the evolution of climate, the UNFCCC and IPCC.

An ECV is a physical, chemical or biological variable, or a group of linked variables, that critically contributes to the characterisation of Earth's climate.

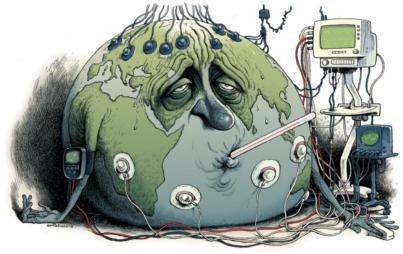




United Nations Framework Convention on Climate Change

2) Essential Climate Variables

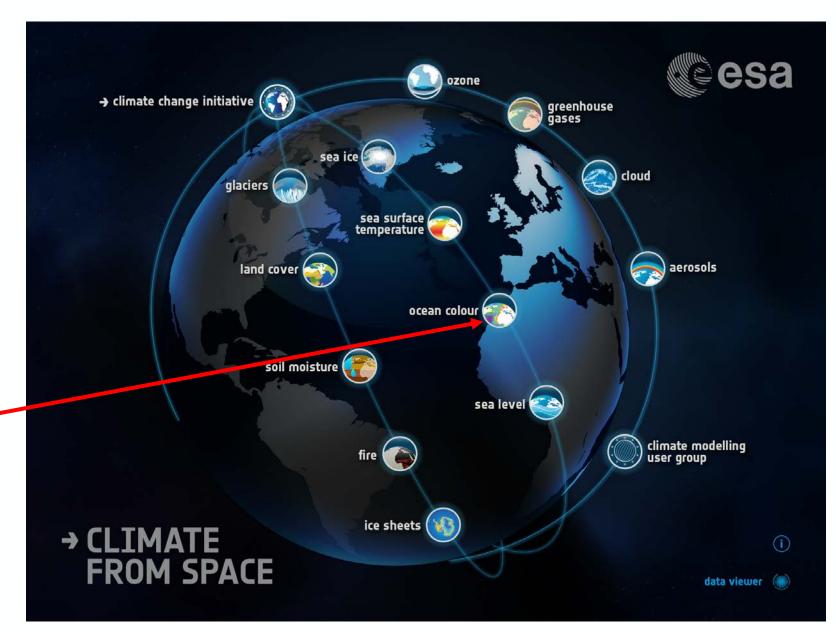




Victor & Kennel (2014) Nature

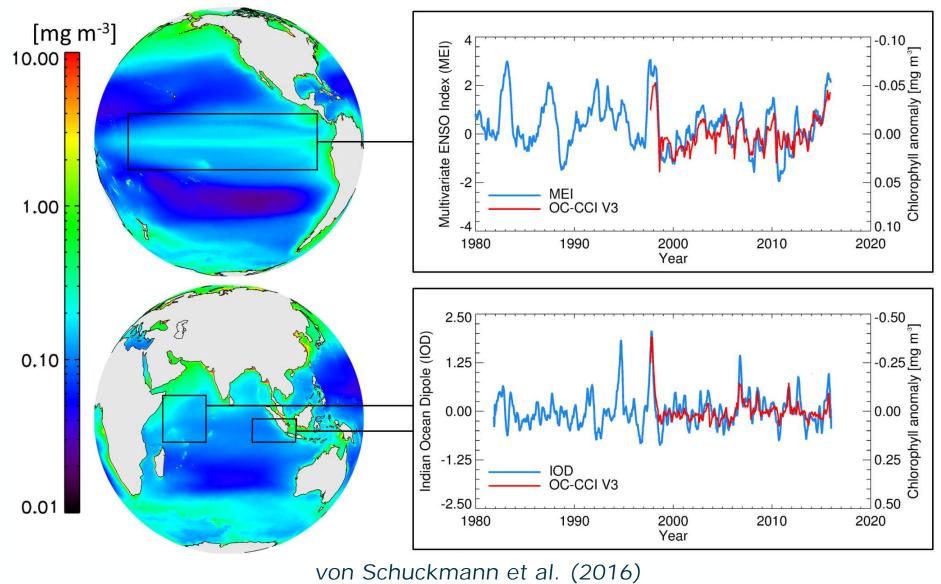


Water leaving radiance Chlorophyll-a concentration



2) Essential Climate Variables





https://doi.org/10.1080/1755876X.2016.1273446

3) Ocean colour datasets requirements for climate



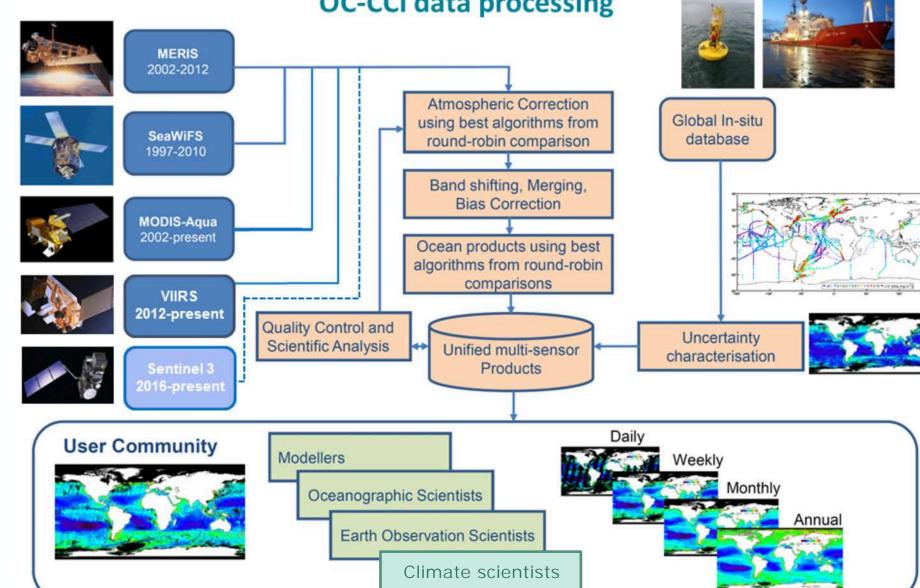
Goal: create the most complete and consistent, error- characterized time series of multi-sensor global satellite ECV products for climate research and modeling meeting GCOS requirements

GCOS-200 October 2016						
Variable/ Parameter	Horizontal Resolution	Vertical Resolution	Temporal Resolution	Accuracy	Stability	The Global Observing System for Climate: Implementation Needs
Water Leaving Radiance	4km	N/A	daily	5% (blue to green)	0.5%	
Chlorophyll-a concentration	4km	N/A	weekly averages	30%	3%	GCOS-200 (GOO5-214)

https://www.ncdc.noaa.gov/gosic/gcos-essential-climate-variable-ecv-data-access-matrix/gcos-ocean-biogeochemistry-ecv-ocean-color

3) Ocean colour datasets requirements for climate





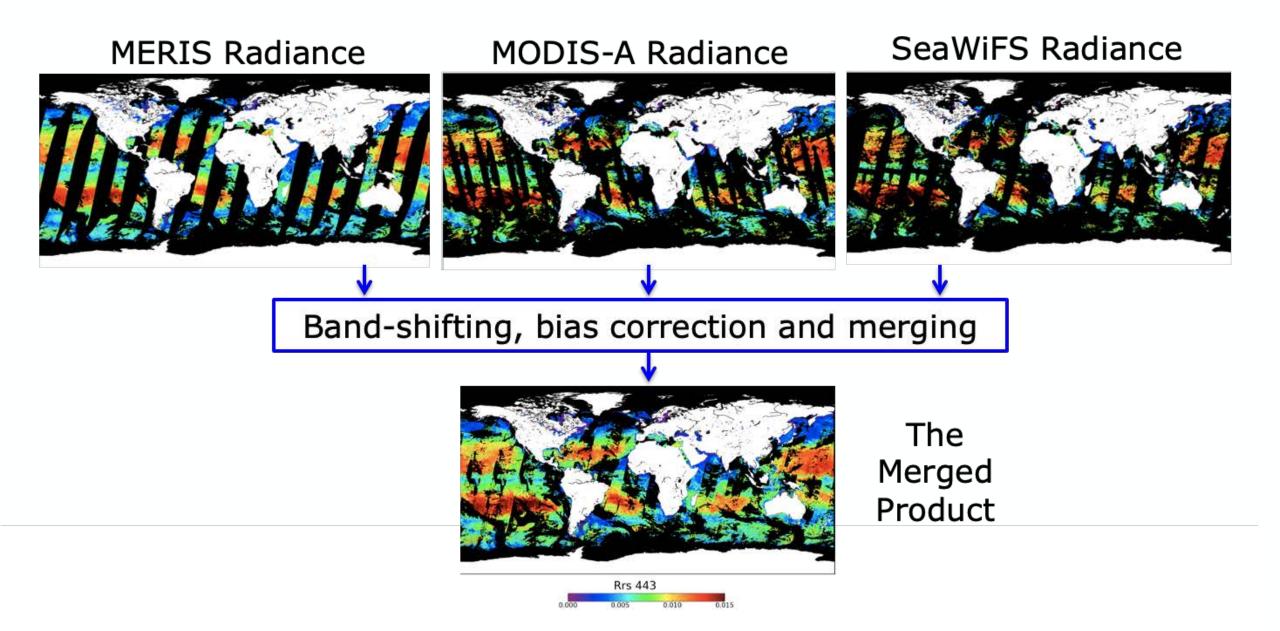
OC-CCI data processing



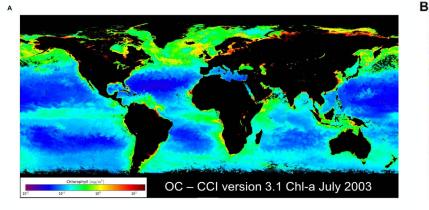
Sathyendranath et al. (2019) https://doi.org/10.3 390/s19194285

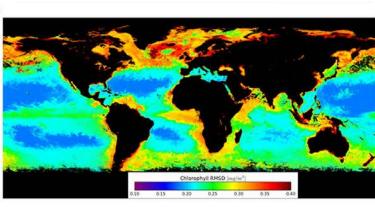
http://www.esaoceancolour-cci.org



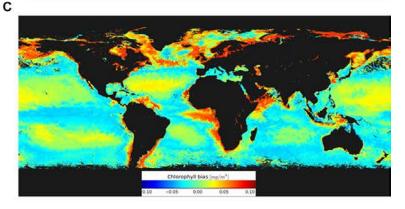




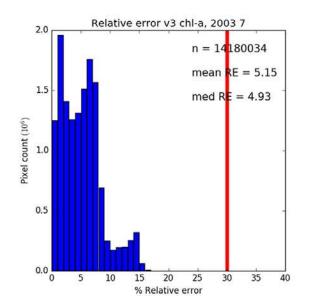




D



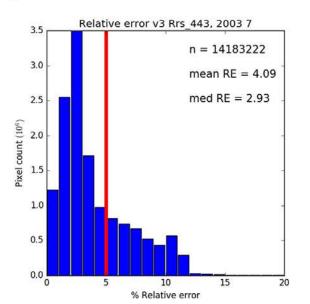
Е



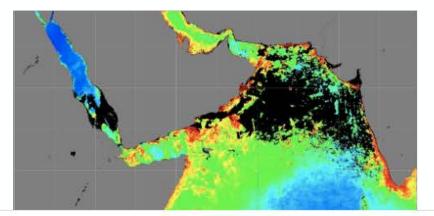
Uncertainty according to a method that uses optical water classification

Moore et al. 2009 https://doi.org/10.1016/j.rse.2009.07.016 Jackson et al. 2017 https://doi.org/10.1016/j.rse.2017.03.036)

Meeting current GCOS requirements for majority of ocean

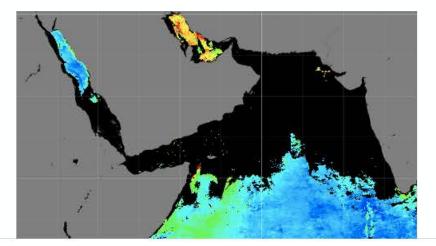




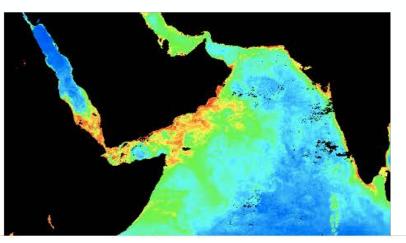


MODIS July Climatology from NASA



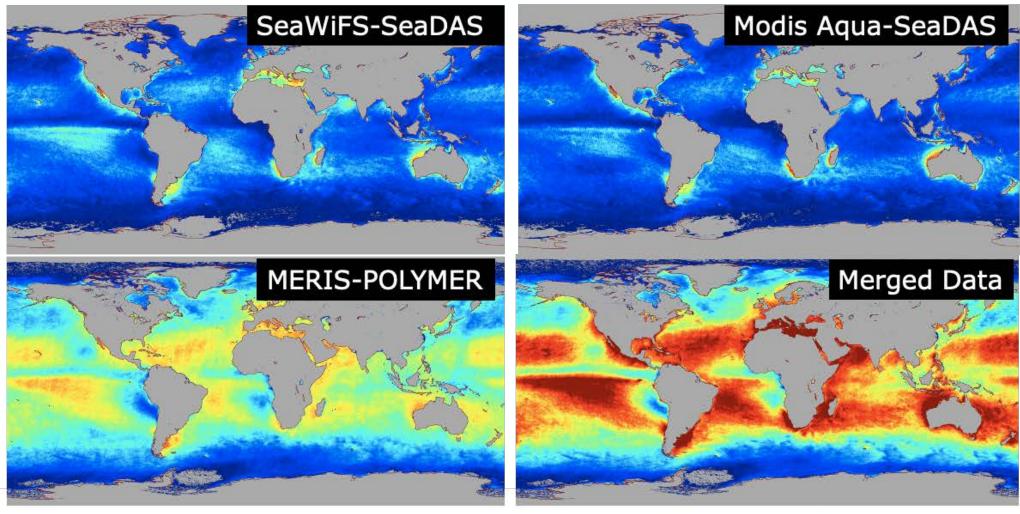


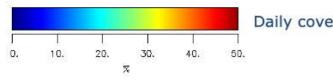
CZCS July Climatology from NASA



OC-CCI July 2003

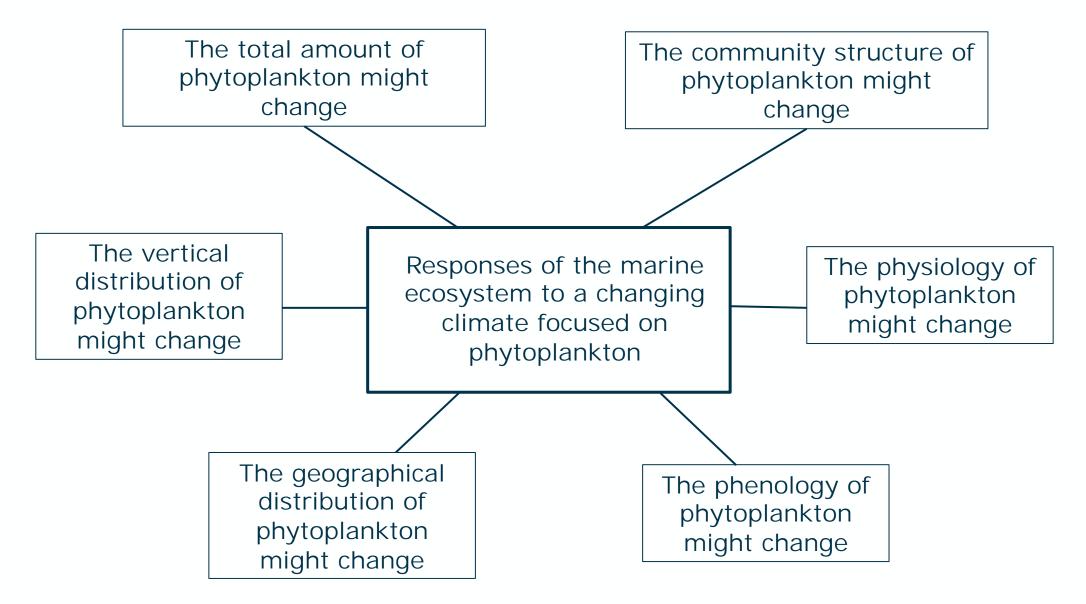






Daily coverage in %

4) Responses of the marine ecosystem to climate change **@esa**



Sathyendranath et al. (2017) <u>https://doi.org/10.1016/j.rse.2017.04.017</u>

4) Responses: Total phytoplankton biomass may change Cesa

Vol 466 29 July 2010 doi:10.1038/nature09268

ARTICLES

nature

Global phytoplankton decline over the past century

Daniel G. Boyce¹, Marlon R. Lewis² & Boris Worm¹

Access provided by ESA-ESTEC European Space Agency

Brief Communications Arising | Published: 13 April 2011

nature

Does blending of chlorophyll data bias temporal trend?

David L. Mackas 🛤

Nature 472, E4-E5 (14 April 2011) Download Citation ±

Abstract

ARISING FROM D. G. Boyce, M. R. Lewis & B. Worm <u>Nature 466, 591–596</u> (2010)10.1038/nature09268; Boyce et al. reply

LIMNOLOGY and OCEANOGRAPHY: METHODS



Integrating global chlorophyll data from 1890 to 2010

Daniel G. Boyce*, Marlon Lewis, and Boris Worm

Abstract

n	a	h	D	1	e	
	nationa					

Access provided by ESA-ESTEC European Space Agency

Brief Communications Arising | Published: 13 April 2011

A measured look at ocean chlorophyll trends

Ryan R. Rykaczewski 🏧 & John P. Dunne

Nature 472, E5-E6 (14 April 2011) Download Citation ±

Abstract

ARISING FROM D. G. Boyce, M. R. Lewis & B. Worm Nature 466, 591–596 (2010)10.1038/nature09268; Boyce et al. reply



-0.2 0.0 0.2 0.4 Rate of Chl change (mg m⁻³ yr⁻¹)

Instant I bernational journal of science

-04

Access provided by ESA-ESTEC European Space Agency

Brief Communications Arising | Published: 13 April 2011

Is there a decline in marine phytoplankton?

Abigail McQuatters-Gollop ⁶⁰, Philip C. Reid, Martin Edwards, Peter H. Burkill, Claudia Castellani, Sonia Batten, Winfried Gieskes, Doug Beare, Robert R. Bidigare, Erica Head, Rod Johnson, Mati Kahru, J. Anthony Koslow & Angelica Pena

Nature 472, E6-E7 (14 April 2011) Download Citation ±

Abstract

ARISING FROM D. G. Boyce, M. R. Lewis & B. Worm Nature **466**, 591-596 (2010)10.1038/nature09268; Boyce et al. reply

Contents lists available at ScienceDirect

Progress in Oceanography

journal homepage: www.elsevier.com/locate/pocean

Estimating global chlorophyll changes over the past century Daniel G. Boyce ^{a,b,c,*}, Michael Dowd ^d, Marlon R. Lewis ^e, Boris Worm ^a

OCEANOGRAPHY

Century of phytoplankton change

David A. Siegel and Bryan A. Franz

Phytoplankton biomass is a crucial measure of the health of ocean ecosystems. An impressive synthesis of the relevant data, stretching back to more than 100 years ago, provides a connection with climate change.

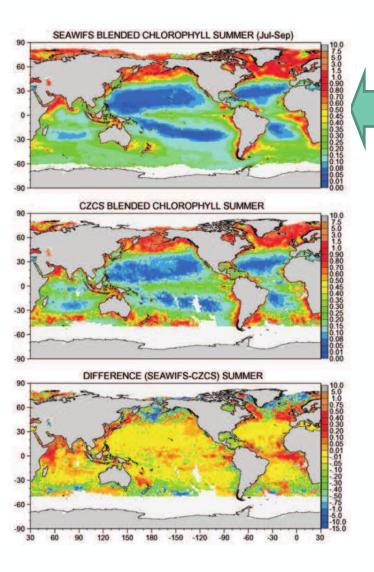
In 1865, Father Pietro Angelo Secchi was asked to map the clarity of the Mediterranean Sea for the Papal navy. He invented the simplest of oceanographic instruments: a 20-centimetre-wide white disk that is lowered until the observer loses sight of it, and for nearly 100 years determinations of Secchi depth were a routine part of oceanographic

observations^{1,2} (Fig. 1, overleaf). Secchi-depth determinations assess light penetration in the upper ocean, and can be related to phytoplankton abundance. Along with measurements of the upper-ocean concentration of chlorophyll, which is found in all phytoplankton, Secchi-disk depths provide the only data available for assessing changes in the global

© 2010 Macmillan Publishers Limited. All rights reserved

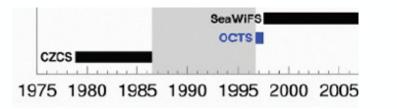
"The analyses of Boyce et al. document the historical record. Looking into the future, however, satellite measurements will be the main source of data for assessing change in pelagic ecosystems."

4) Responses: Total phytoplankton biomass may change Cesa

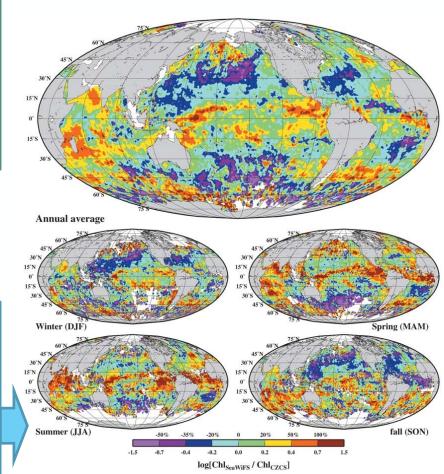


Global spatial distributions and seasonal variability of ocean chlorophyll were similar, but global means decreased over the two observational segments.

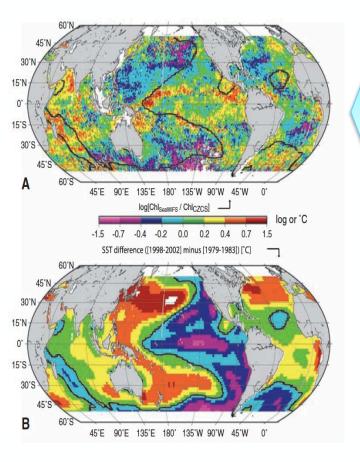
Gregg & Conkright (2002) https://doi.org/10.1029/2002GL014689



The analysis of decadal changes from the CZCS to the SeaWiFS era shows an overall increase of the world ocean average chlorophyll concentration by about 22%, mainly due to large increases in the intertropical areas Antoine et al .(2005) https://doi.org/10.1029/2004JC002620

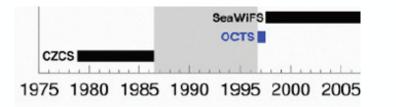


4) Responses: Total phytoplankton biomass may change esa



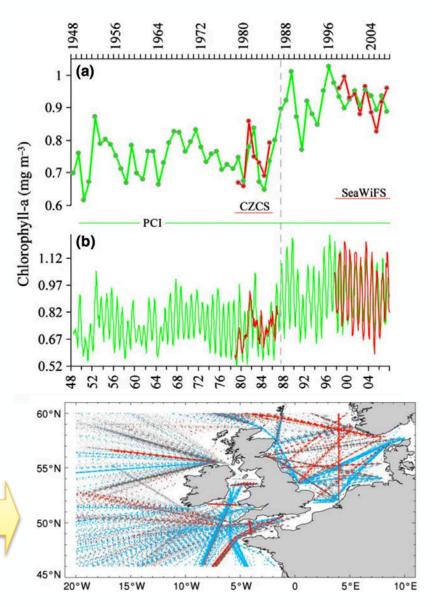
Multidecadal changes in global phytoplankton abundances are related to basin-scale oscillations of the physical ocean.

Martinez et al . (2009) https://doi.org/10.1126/science.1177012

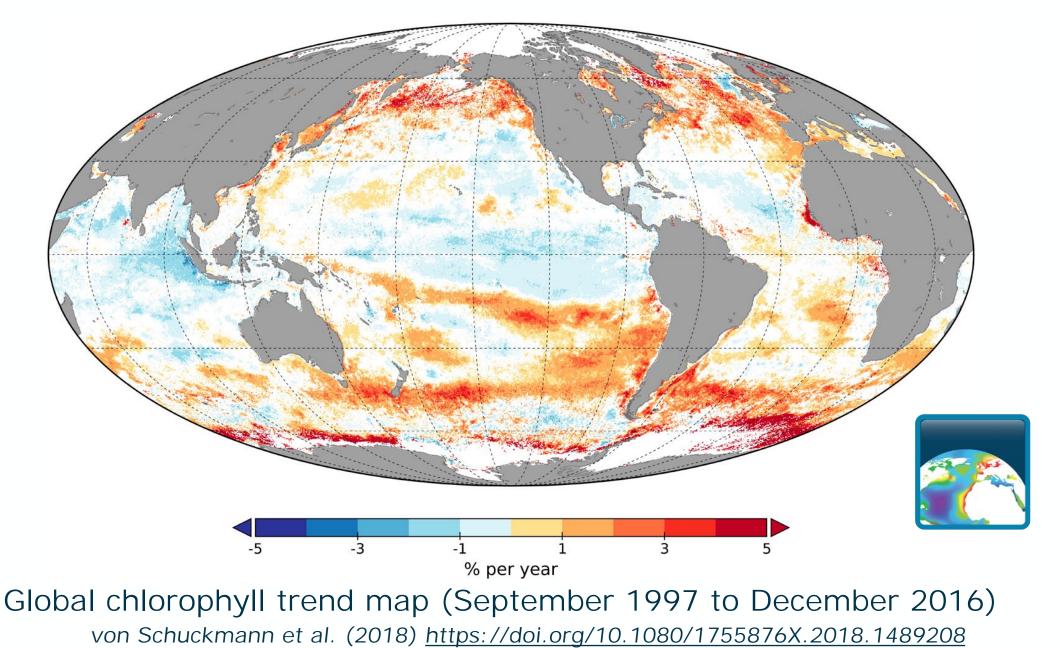


Results from merging CPR data with CZCS and SeaWiFS reflect a clear increasing pattern of Chl-a in the North East Atlantic

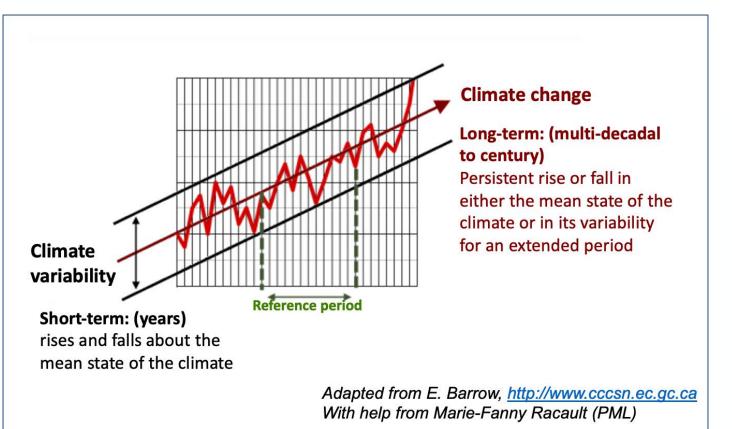
> Raitsos et al .(2013) https://doi.org/10.1111/gcb.12457



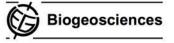
4) Responses: Total phytoplankton biomass may change Cesa



4) Responses: Total phytoplankton biomass may change esa

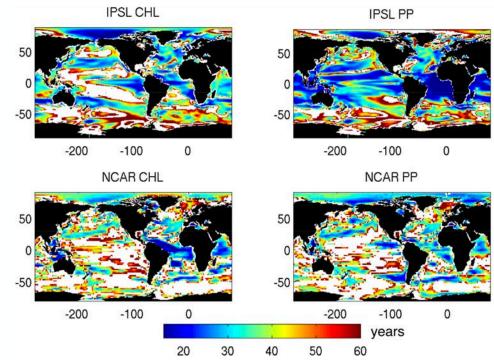


Biogeosciences, 7, 621–640, 2010 www.biogeosciences.net/7/621/2010/ © Author(s) 2010. This work is distributed under the Creative Commons Attribution 3.0 License.



Detection of anthropogenic climate change in satellite records of ocean chlorophyll and productivity

S. A. Henson^{1,*}, J. L. Sarmiento¹, J. P. Dunne², L. Bopp³, I. Lima⁴, S. C. Doney⁴, J. John², and C. Beaulieu¹



4) Responses: Total phytoplankton biomass may changeesa

10

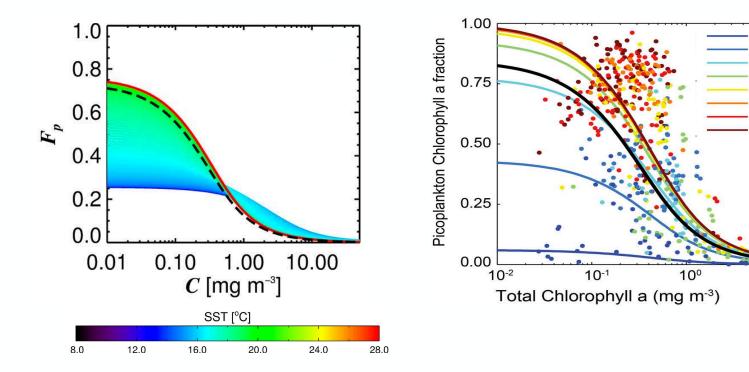
14

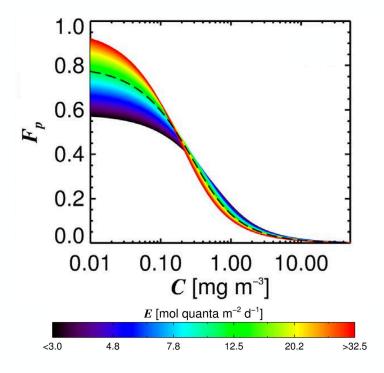
18

22

26

10¹





Fraction of picoplankton to total chl-a in the North Atlantic as a function of SST.

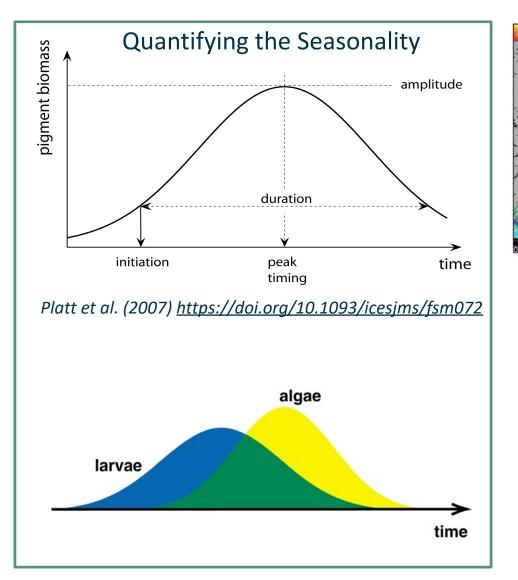
Brewin et al. (2017) https://doi.org/10.3389/fmars.2 017.00104 Fraction of picoplankton to total chl-a in the Global Ocean as a function of SST.

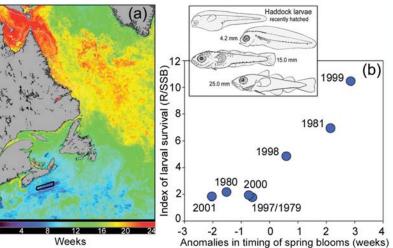
Ward et al. (2015) https://doi.org/10.1371/journal .pone.0135581 Fraction of picoplankton to total chl-a in the Global Ocean as a function of light

Brewin et al. (2015) https://doi.org/10.1016/j.rse.20 <u>15.07.004</u>

4) Responses: Phytoplankton phenology







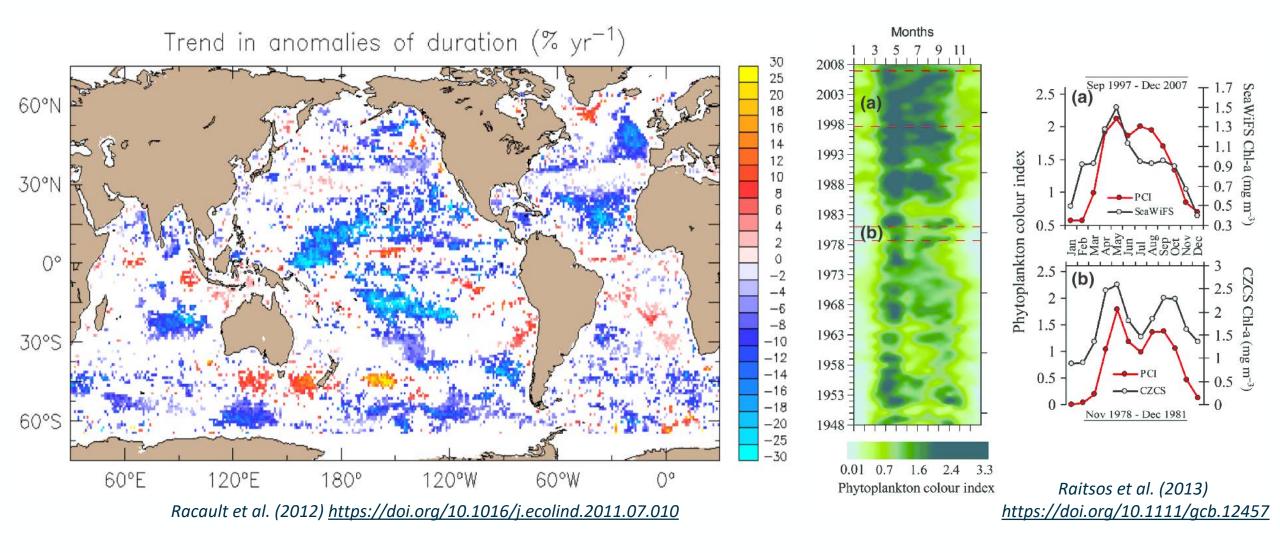
Inter-annual variations in the timing of the spring bloom impacts survival of larval fish

> Platt et al. (2003) https://doi.org/10.1038/423398b

Koeller et al. (2009) <u>http://doi.org/10.1126/science.1170987</u>

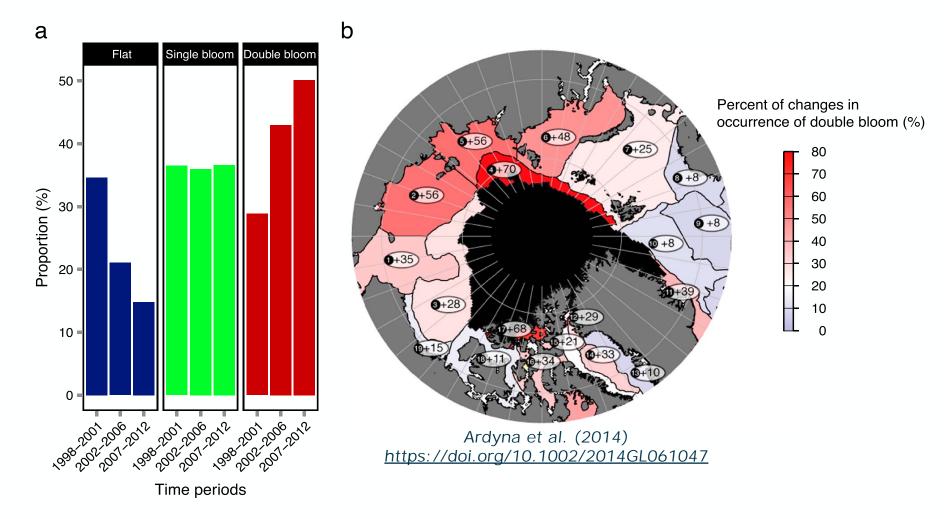
4) Responses: Phytoplankton phenology





4) Responses: Changing polar regions



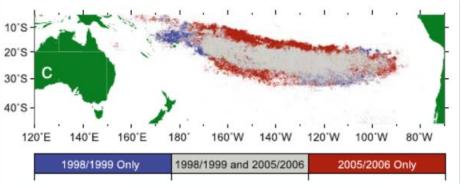


In a changing Arctic Ocean, a phenology of biological productivity with two phytoplankton blooms and two peaks of sedimentation may become prevalent.

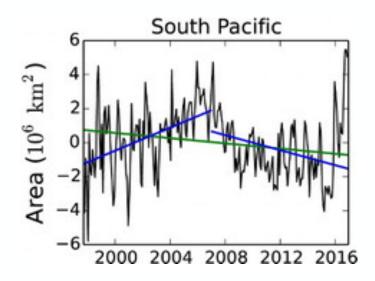
4) Responses: Other



Oligotrophic regions

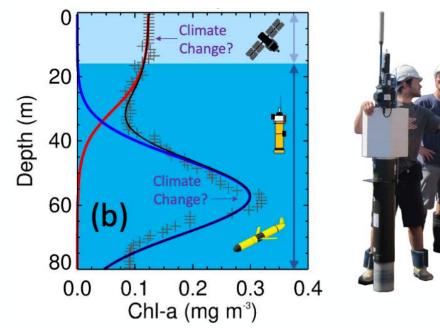


Polovina et al. (2008) Geophys. Res. Lett. https://doi.org/10.1029/2007GL031745

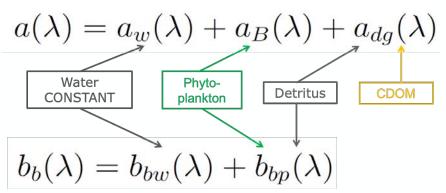


von Schuckmann et al. (2018) https://doi.org/10.1080/1755876X.2018.1489208

Vertical structure



Changes in other optically-active substances



5) Other applications Fisheries

Water Quality

Harmful algal blooms

Aquaculture

Marine pollution

Bio-feedback mechanisms



in Marine Science



Ocean Colour Bibliography

Home » Resources » Ocean Colour Bibliography

The IOCCG bibliography is updated periodically with new references submitted by readers. Another useful ocean colour bibliography is the searchable Historic Ocean Colour Archive assembled by Marcel Wernand, with articles and books written between the 17th and early 20th century.

http://ioccg.org/what-we-do/ioccg-publications/

http://ioccg.org/resources/ocean-colour-bibliography/

http://ioccg.org/what-we-do/ioccg-publications/ioccgreports/

> REVIEW published: 29 August 2019 doi: 10,3389/fmars.2019.00485

U
Check for updates

Groom et al. (2019) https://doi.org/10.3389/f mars.2019.00485

Marine biodiversity and function

Satellite Ocean Colour: Current Status and Future Perspective

Steve Groom^{1,2*}, Shubha Sathyendranath^{1,2}, Yai Ban³, Stewart Bernard⁴,



Part 3: Outline



- 1) The Greenhouse effect
- 2) Essential Climate Variables
- 3) Ocean colour datasets requirements for climate
- 4) Responses of the marine ecosystem to climate change
 - Total phytoplankton biomass
 - Phytoplankton community structure
 - Phytoplankton phenology
 - Other responses
- 5) Other applications