

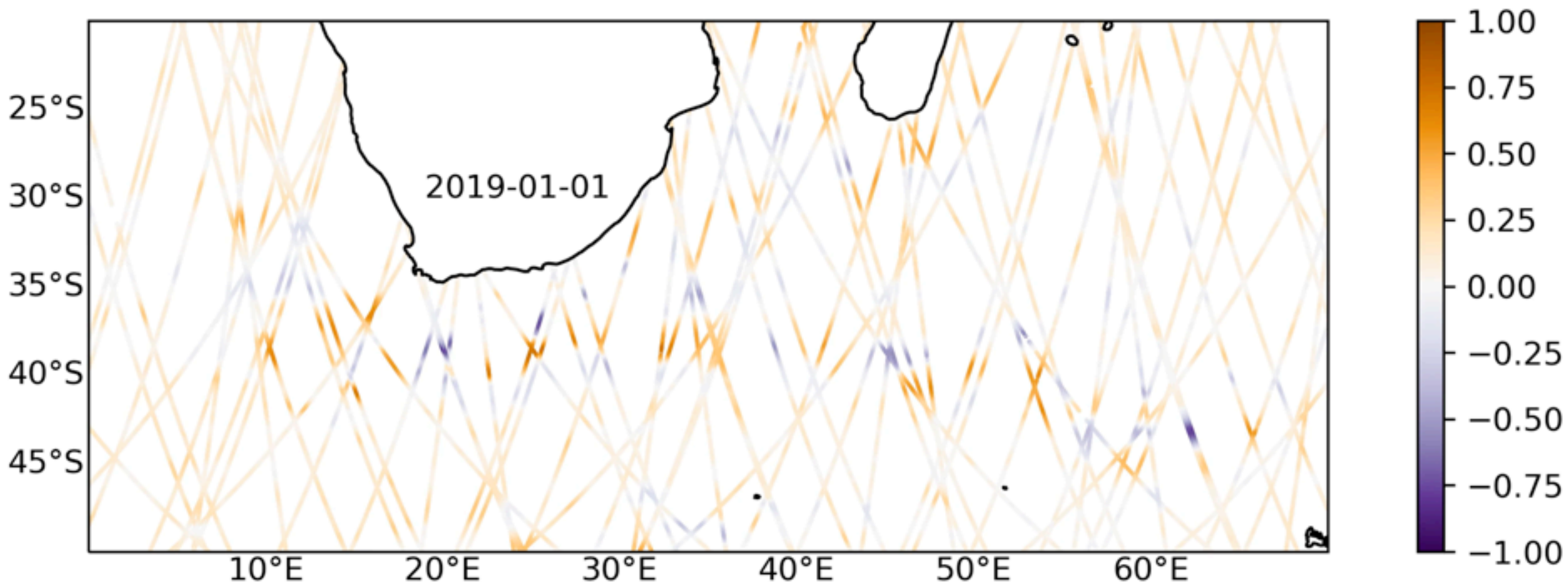


High resolution gridded altimetric maps by  
the Back-and-Forth Nudging (BFN) method

Florian Le Guillou - ESA

# Ocean circulation estimated from altimetry

Several altimeters provide Sea Level Anomaly (SLA) observations along **one-dimensional tracks**

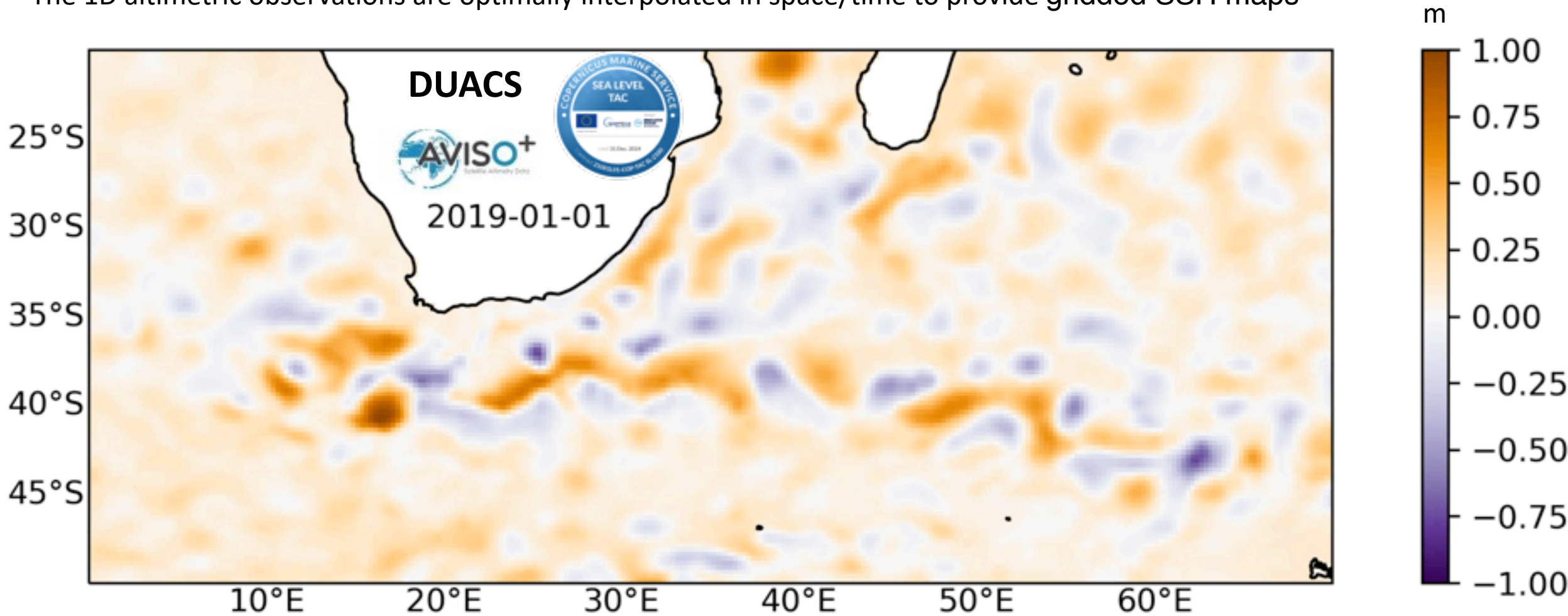


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# Ocean circulation estimated from altimetry

The 1D altimetric observations are optimally interpolated in space/time to provide gridded SSH maps

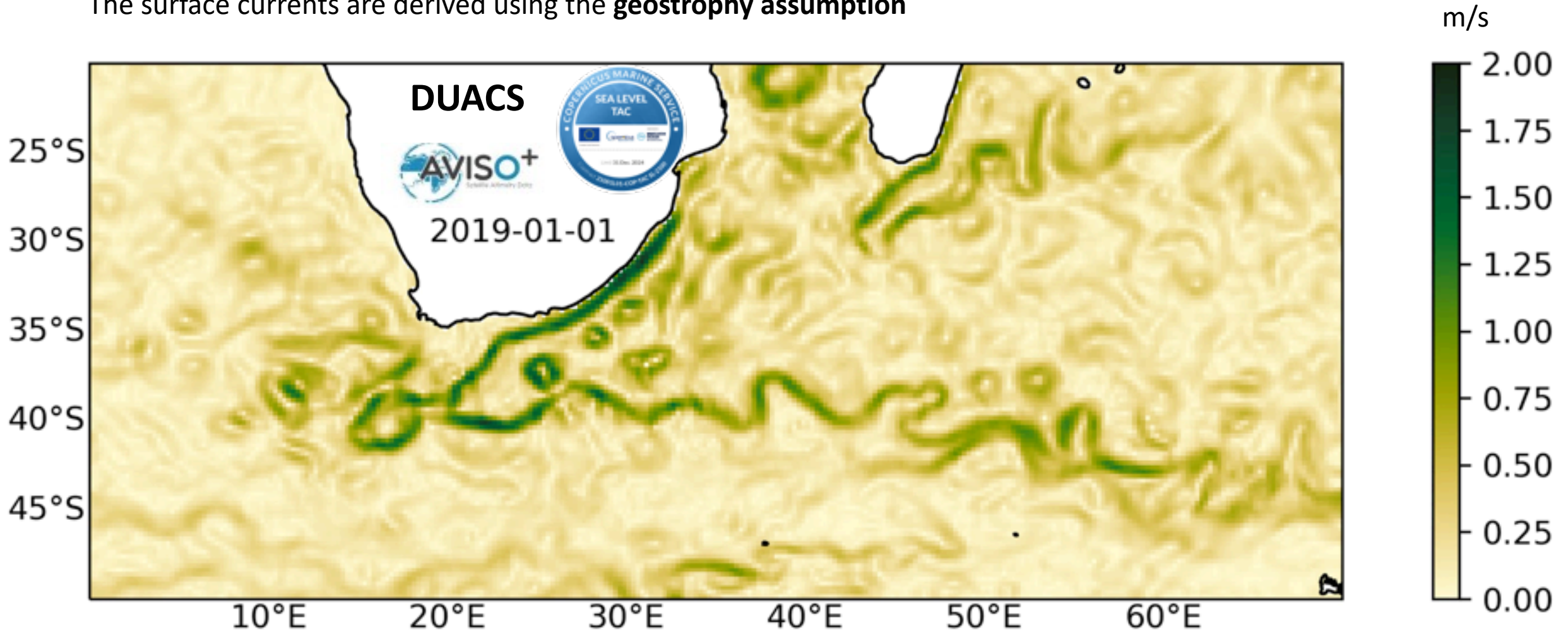


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# Ocean circulation estimated from altimetry

The surface currents are derived using the **geostrophy assumption**



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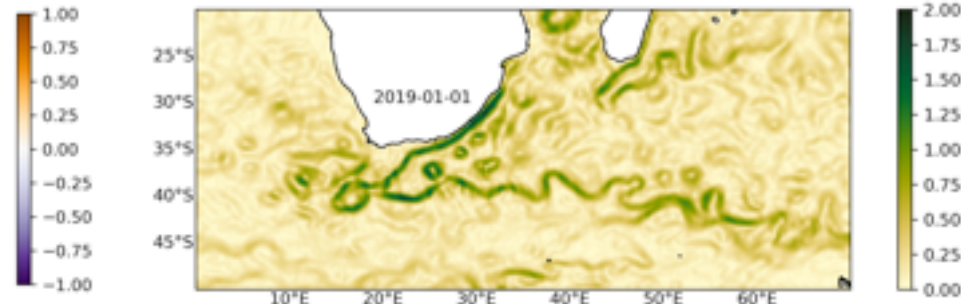
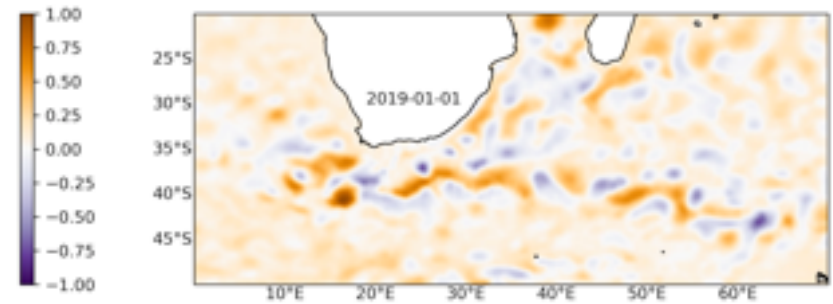
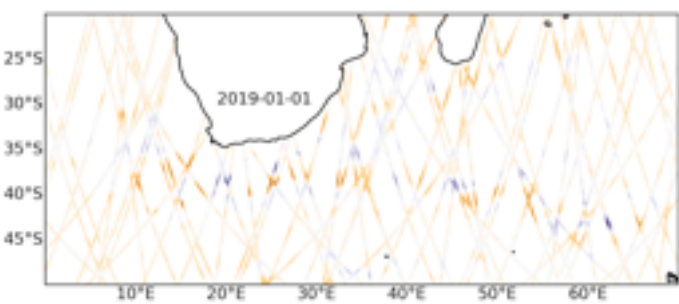


# Limits of the operational DUACS products

Along-track SLA

Mapped SLA

Surface currents



Optimal (linear) interpolation

Geostrophy assumption

**Limit:** Space-time resolution limited to 200km – 10 days  
(Ballarotta et al., 2019)

**Challenge:** Improve the mapping algorithm to take into account non linear dynamics

**Opportunity:** Add non linear dynamical constrain in the mapping procedure (e.g. Ubelmann et al. 2015 and **this talk**)

**Limit:** Only the geostrophic circulation is estimated

**Challenge:** Include ageostrophic processes (e.g. Ekman, tidal, inertial currents)

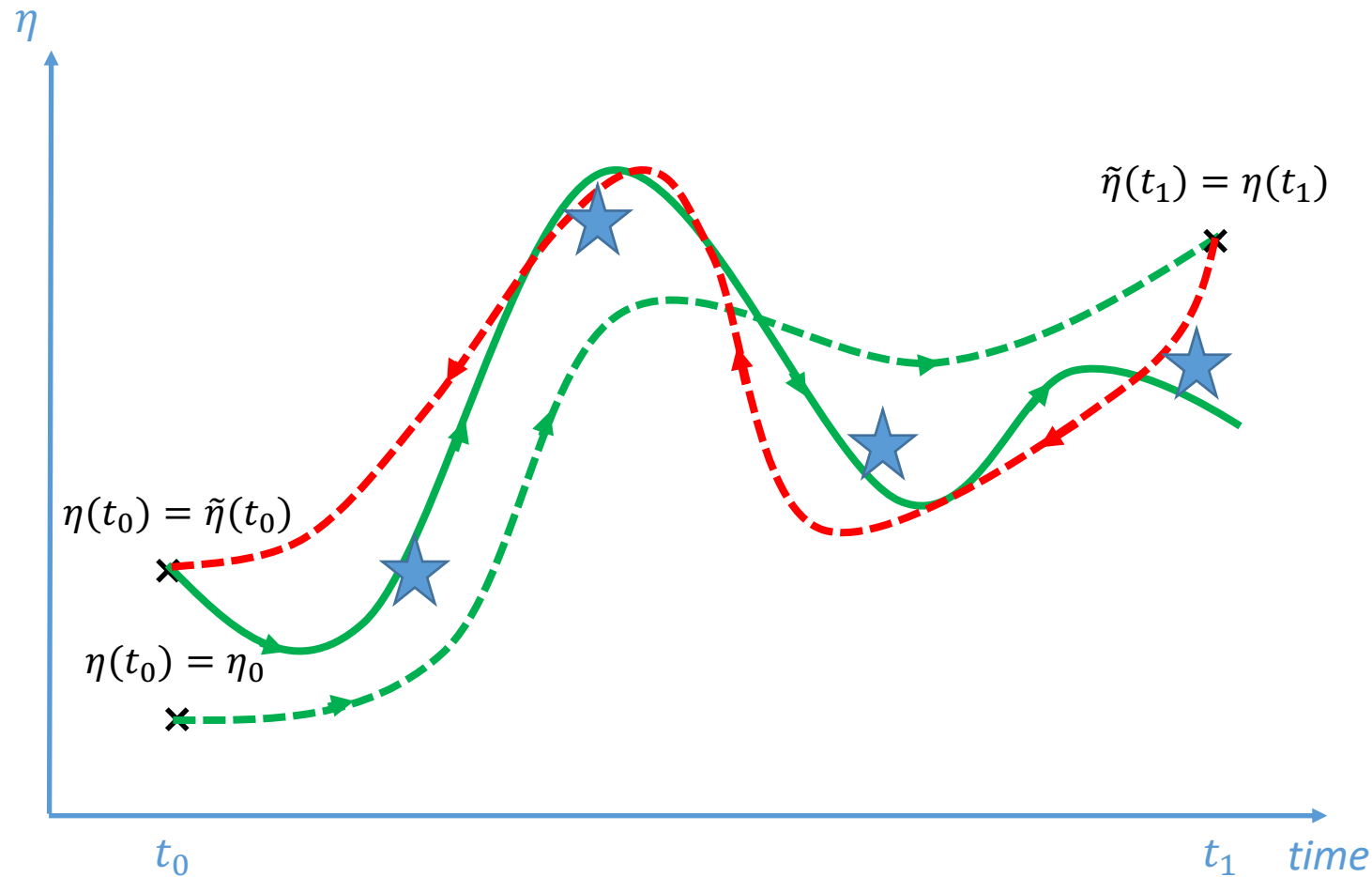
**Opportunity:** Use other spaceborne and/or *in situ* observations, e.g. Drifters (**next talk** by Clement Ubelmann)

# The Back-and-Forth-Nudging (BFN) method

Quasi-Geostrophic (QG) dynamics

$$\left. \begin{aligned} q &= \frac{g}{f} \nabla^2 \eta - \frac{f}{H_e} \eta \\ \frac{\partial q}{\partial t} + u_g \frac{\partial q}{\partial x} + v_g \frac{\partial q}{\partial y} &= 0 \end{aligned} \right\} \frac{\partial \eta}{\partial t} = \mathcal{M}(\eta, t)$$

Let's consider some observations  $\eta^{obs}$  over  $[t_0, t_1]$



*Forward nudging*

$$\left\{ \begin{aligned} \eta(t_0) &= \eta_0 \\ \frac{\partial \eta}{\partial t} &= \mathcal{M}(\eta, t) + K(\eta^{obs} - \eta) \end{aligned} \right.$$

*Backward nudging*

$$\left\{ \begin{aligned} \tilde{\eta}(t_0) &= \tilde{\eta}_0 \\ \frac{\partial \tilde{\eta}}{\partial t} &= \mathcal{M}(\tilde{\eta}, t) - K(\eta^{obs} - \tilde{\eta}) \end{aligned} \right.$$

# Experimental set-up

**Study region:** retroflexion of the Agulhas Current

**Time periods:** 2010-2019

**Input dataset:** all available altimetry data

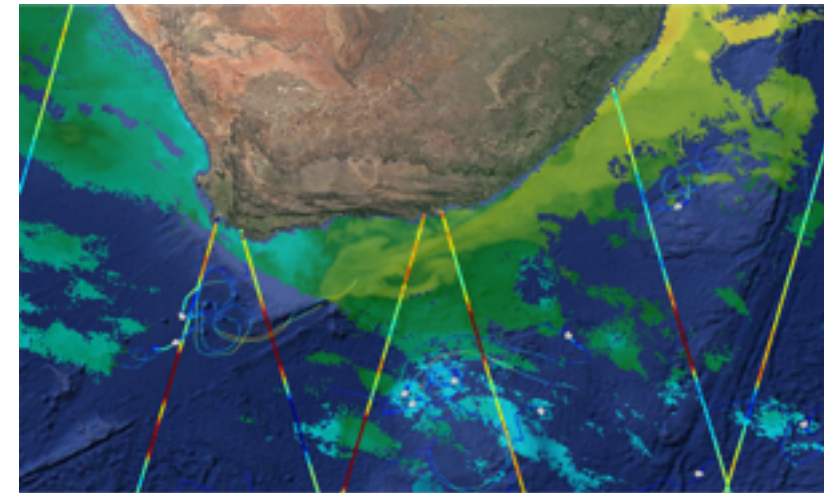
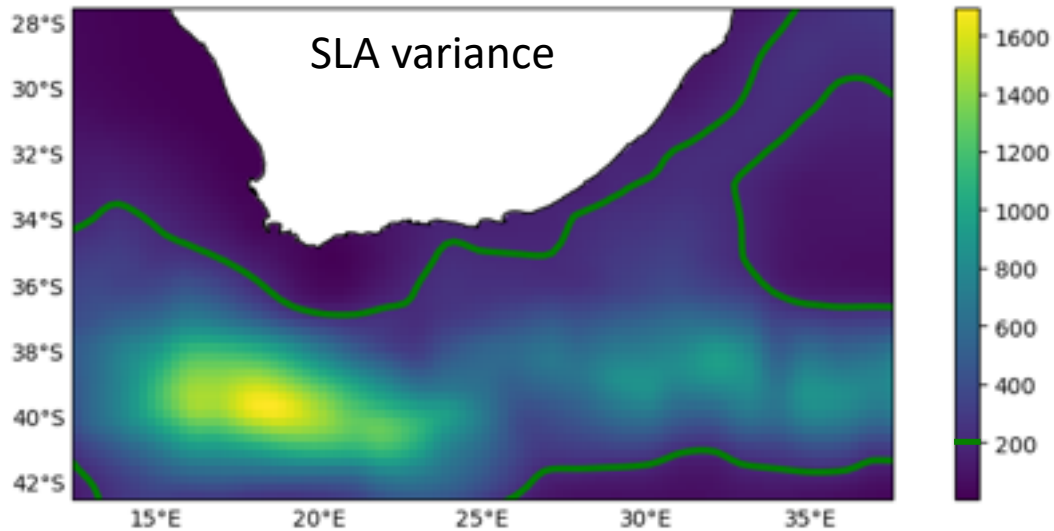
**Methods:** DUACS & BFN-QG

**Independent dataset for validation:**

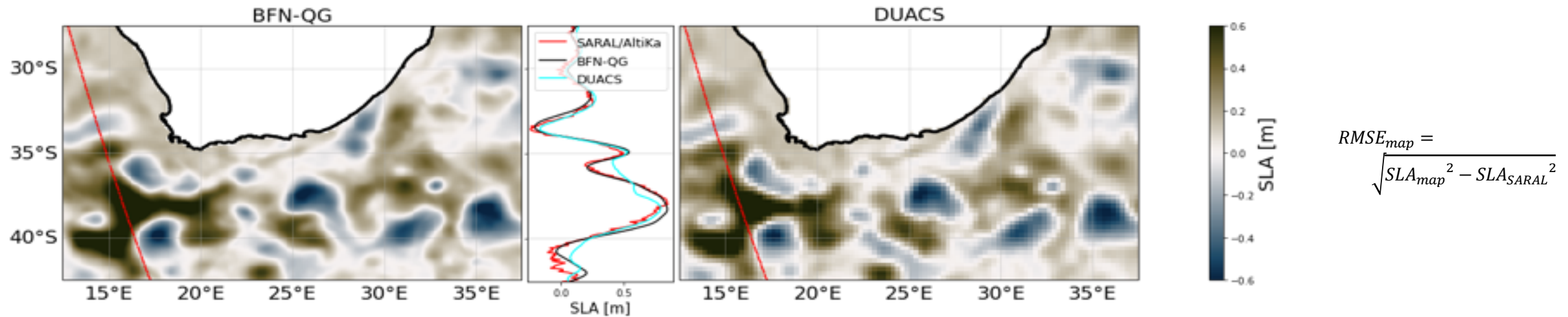
SLA from SARAL/AltiKa, removed from the input dataset (only for year 2019)

SST Seviri L3C (Eumetsat)

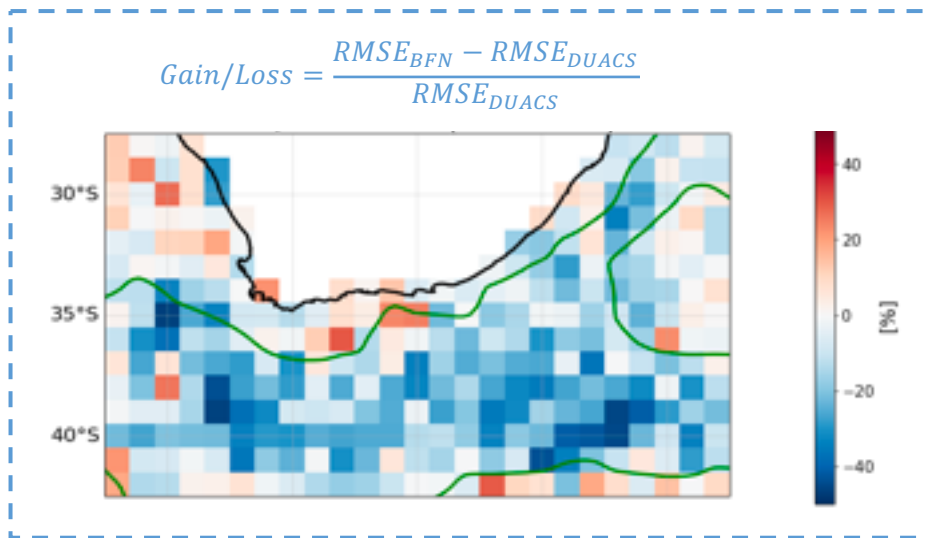
Drifters 15m drogue (Coriolis, CMEMS)



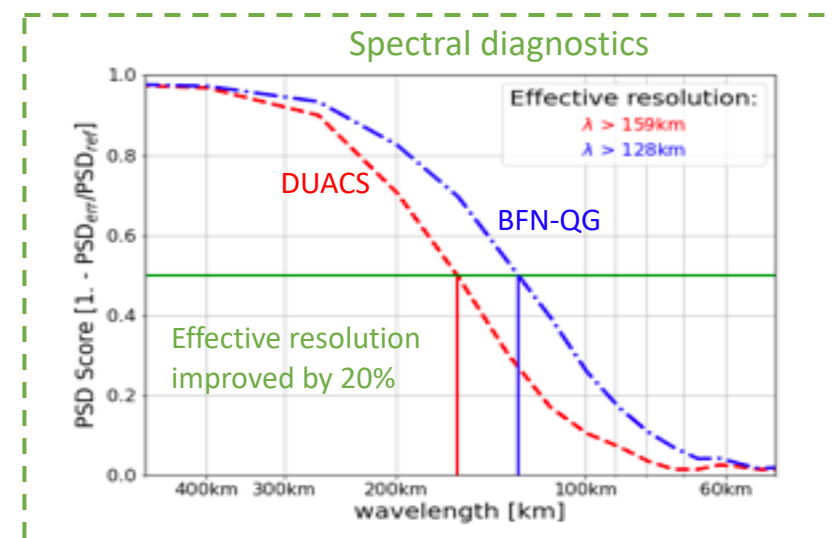
# Results: performances in SLA mapping



$$RMSE_{map} = \sqrt{SLA_{map}^2 - SLA_{SARAL}^2}$$



Average reduction of error ~10%

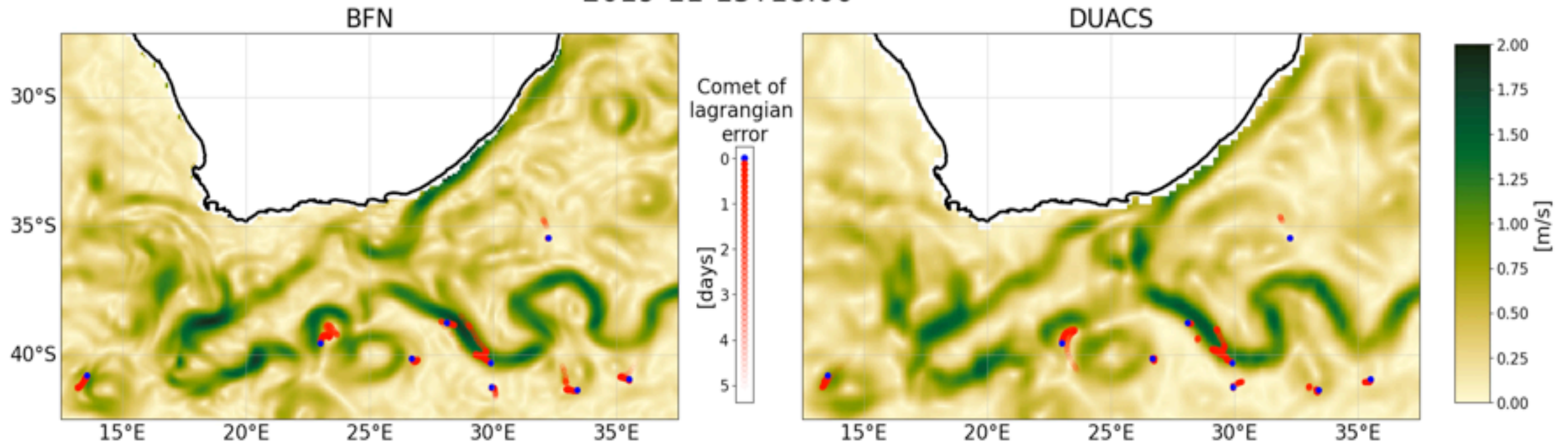




# Results: performances in ocean currents retrieval

## Comparisons with Drifter data

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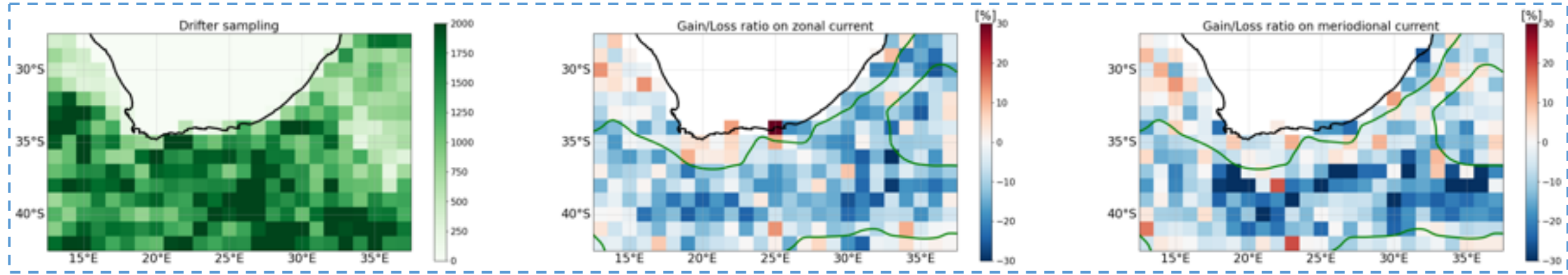


# Results: performances in ocean currents retrieval

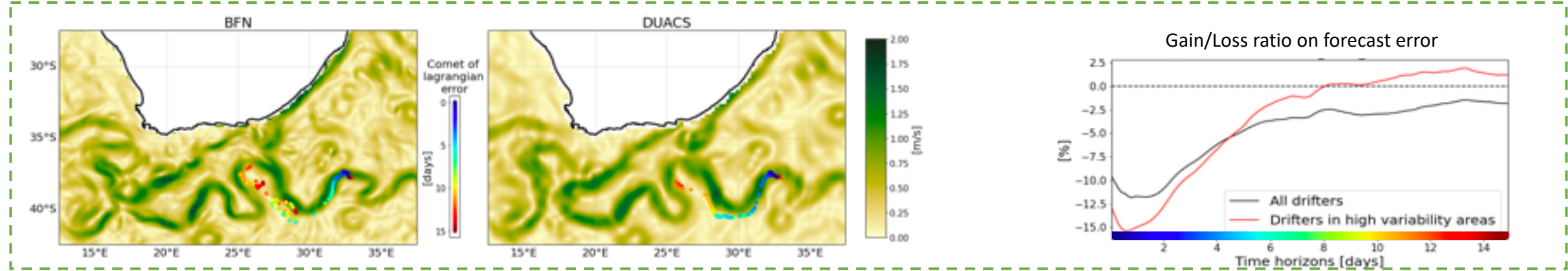


## Comparisons with Drifter data

### Diagnostics on Eulerian velocities



### Diagnostic on Lagrangian velocities



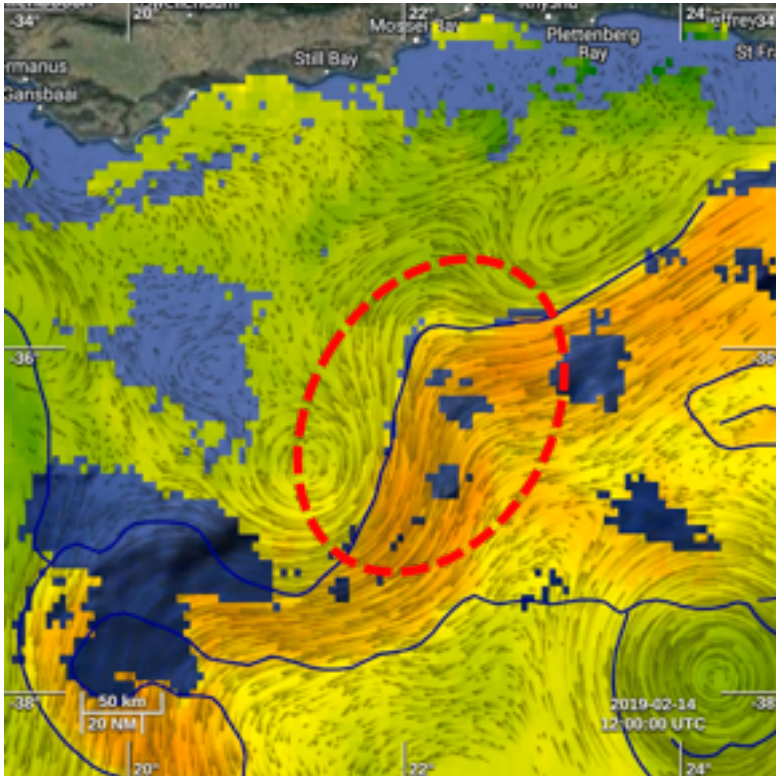
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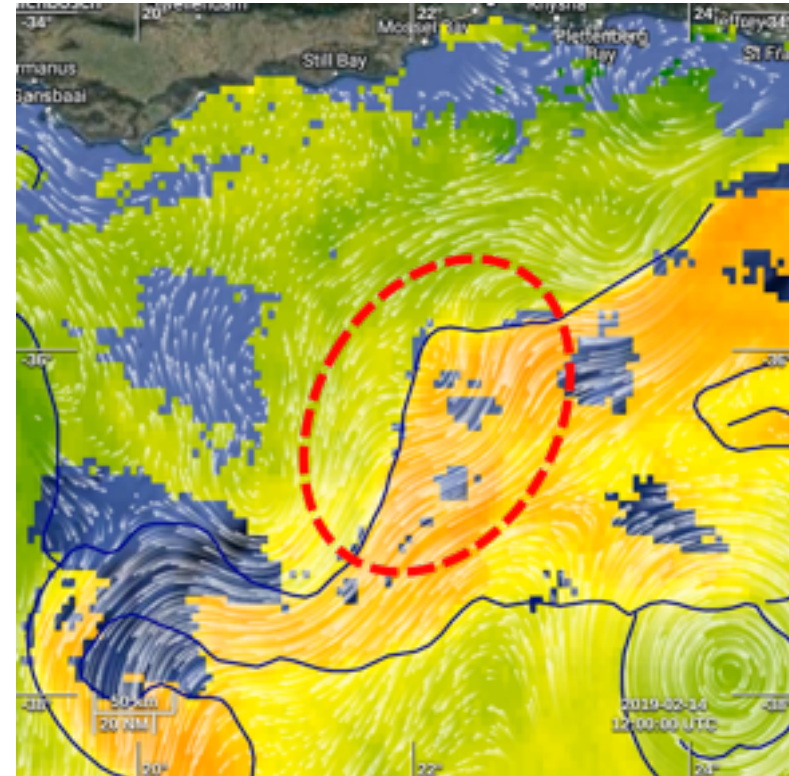
# Results: performances in ocean currents retrieval

## Comparisons with SST data

BFN-QG



DUACS

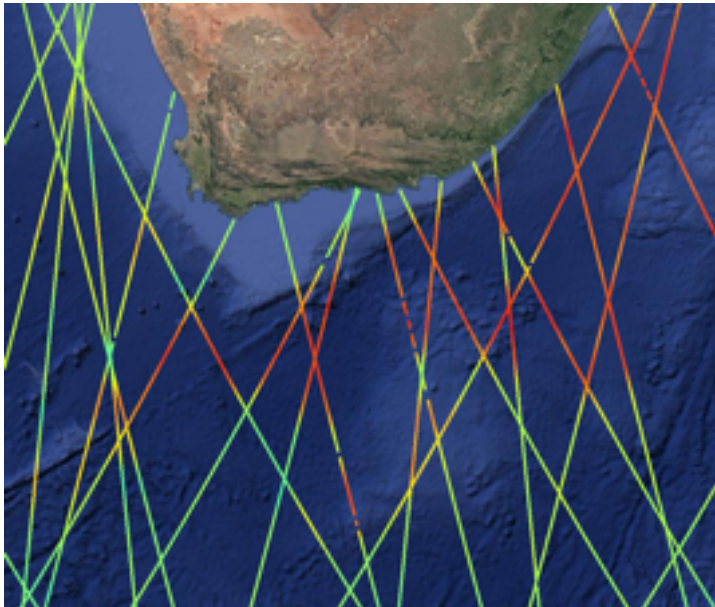


For quantitative assessment, see [Lucile Gautier's talk this afternoon](#)

- BFN-QG improves the quality of SLA/currents retrievals in the Agulhas current
  - Space resolutions of SLA maps improved by 20%
  - Both Eulerian zonal/meridional currents improved by 10%
  - Lagrangian prediction improved for lead times < 5-10 days
  
- What remains to be improved:
  - Performances depend on the dynamical regime (where QG assumptions hold)
  - SLA only gives the geostrophic currents

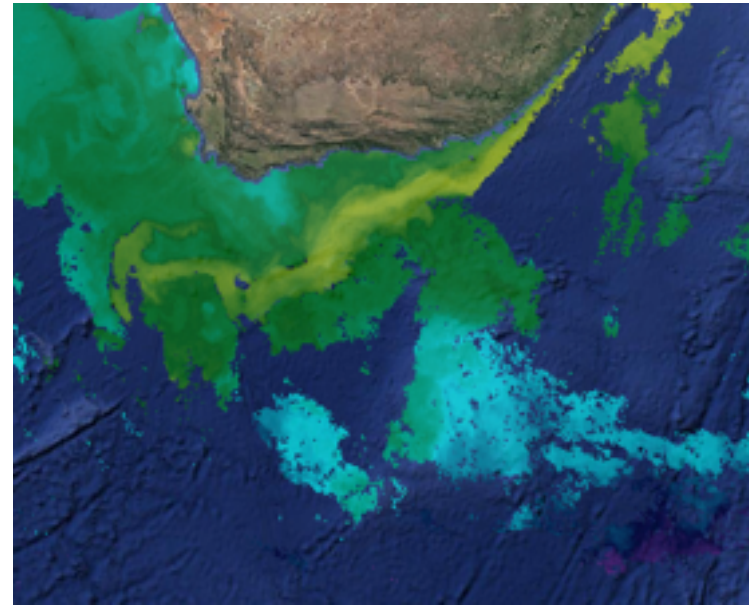
## Multi-sensors assimilation

Altimetry



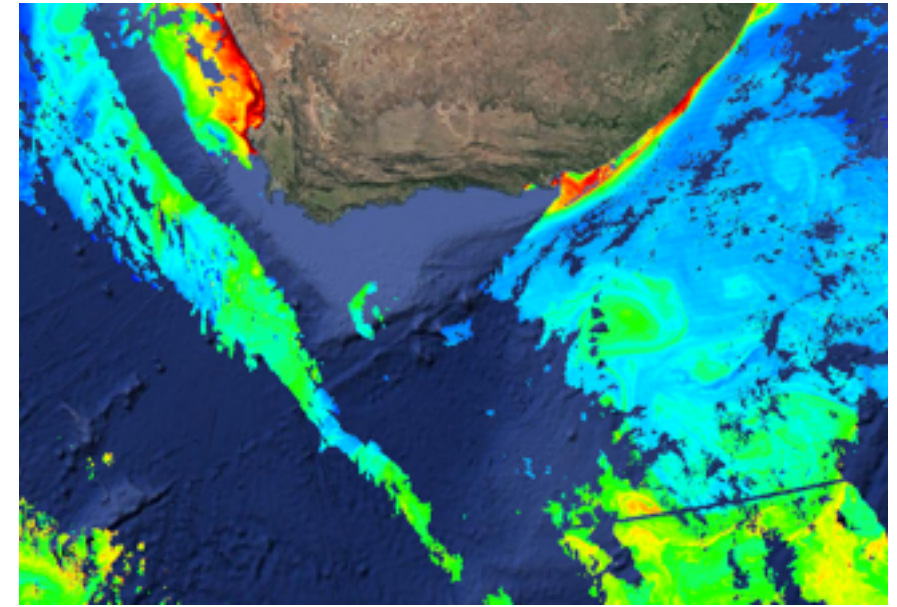
$$\frac{\partial q}{\partial t} + \mathbf{U}_g \cdot \nabla q = 0$$

Sea Surface Temperature



$$\frac{\partial SST}{\partial t} + (\mathbf{U}_g + \mathbf{U}_a) \cdot \nabla SST = F_{SST}$$

Ocean Color



$$\frac{\partial Chl}{\partial t} + (\mathbf{U}_g + \mathbf{U}_a) \cdot \nabla Chl = F_{Chl}$$



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