

# Geodetic mass balance of glaciers

Marin Kneib, glaciologist at Institut des Géosciences de l'Environnement (Grenoble) & University of Innsbruck

# Programme

- 1 hour lecture (17/09 14:00-15:00)
- 2 hours exercise in smaller groups:
  - 17/09 16:00-18:00
  - 18/09 08:30-10:30
  
- Contact: [marin.kneib@univ-grenoble-alpes.fr](mailto:marin.kneib@univ-grenoble-alpes.fr)

# Programme

1. Introduction to geodetic mass balance (what for?)
2. Principles
3. Stereo-processing for the generation of DEMs
4. DEM co-registration & differencing
5. Volume & mass changes calculations
6. Uncertainty estimations
7. Perspective: linking geodetic & glaciological mass balance
- 8. *Exercices with geoutils & xdem (python, Jupyter notebook)***

## Objectives:

- 1. Get familiar with the overall processing chain from satellite image acquisitions to mass change calculation along with the underlying assumptions & limitations**
- 2. Calculate your own geodetic mass balance of a glacier based on two DEMs**

# Geodetic mass balance? What for?



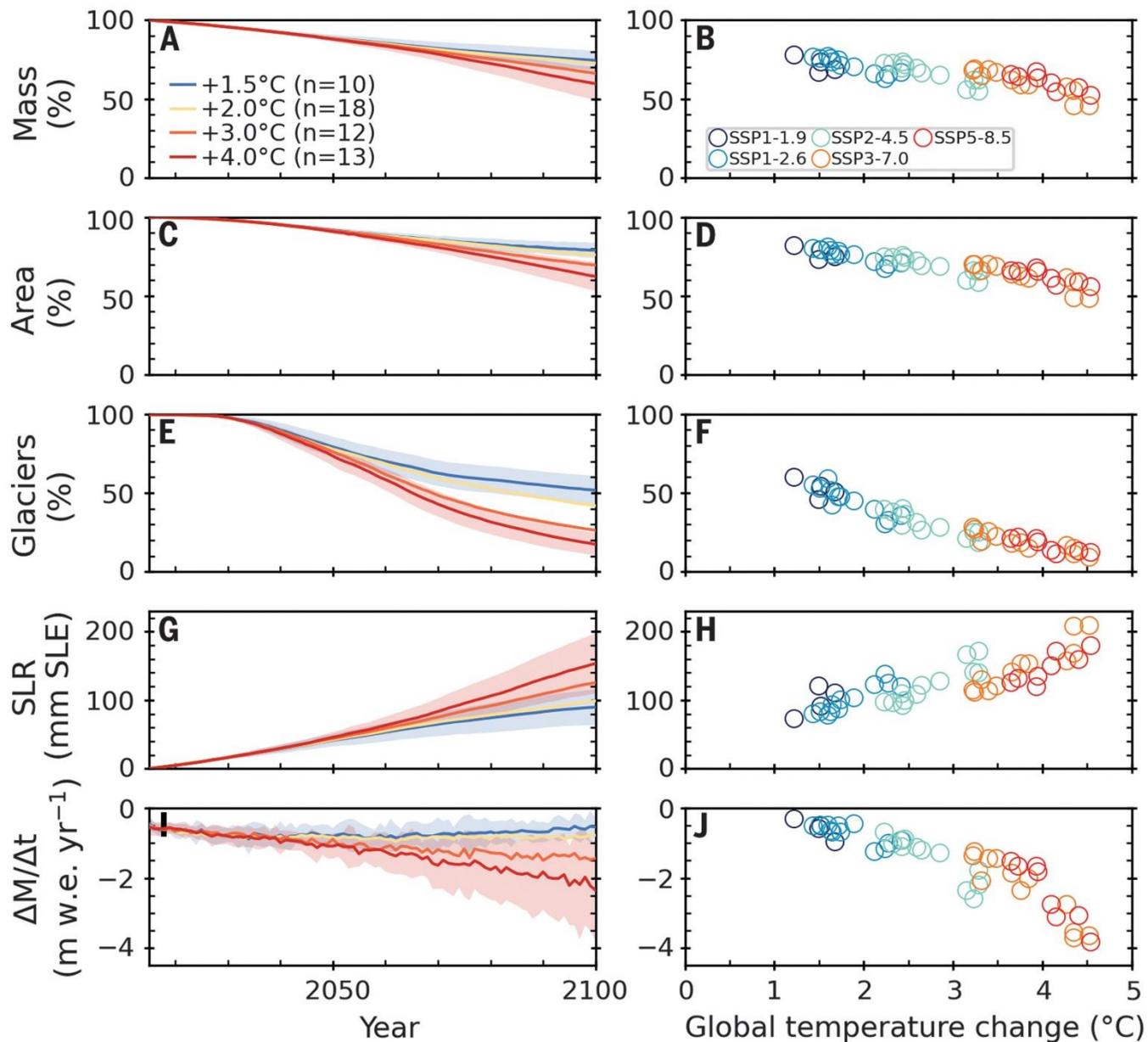
# Geodetic mass balance? What for?

## Global glacier change in the 21st century: Every increase in temperature matters

DAVID R. ROUNCE , REGINE HOCK , FABIEN MAUSSON , ROMAIN HUGONNET , WILLIAM KOCHTITZKY , MATTHIAS HUSS , ETIENNE BERTHIER 

DOUGLAS BRINKERHOFF, LORIS COMPAGNO, [...] AND ROBERT W. MCNABB  [+3 authors](#) [Authors Info & Affiliations](#)

SCIENCE · 5 Jan 2023 · Vol 379, Issue 6627 · pp. 78-83 · DOI: 10.1126/science.abo1324



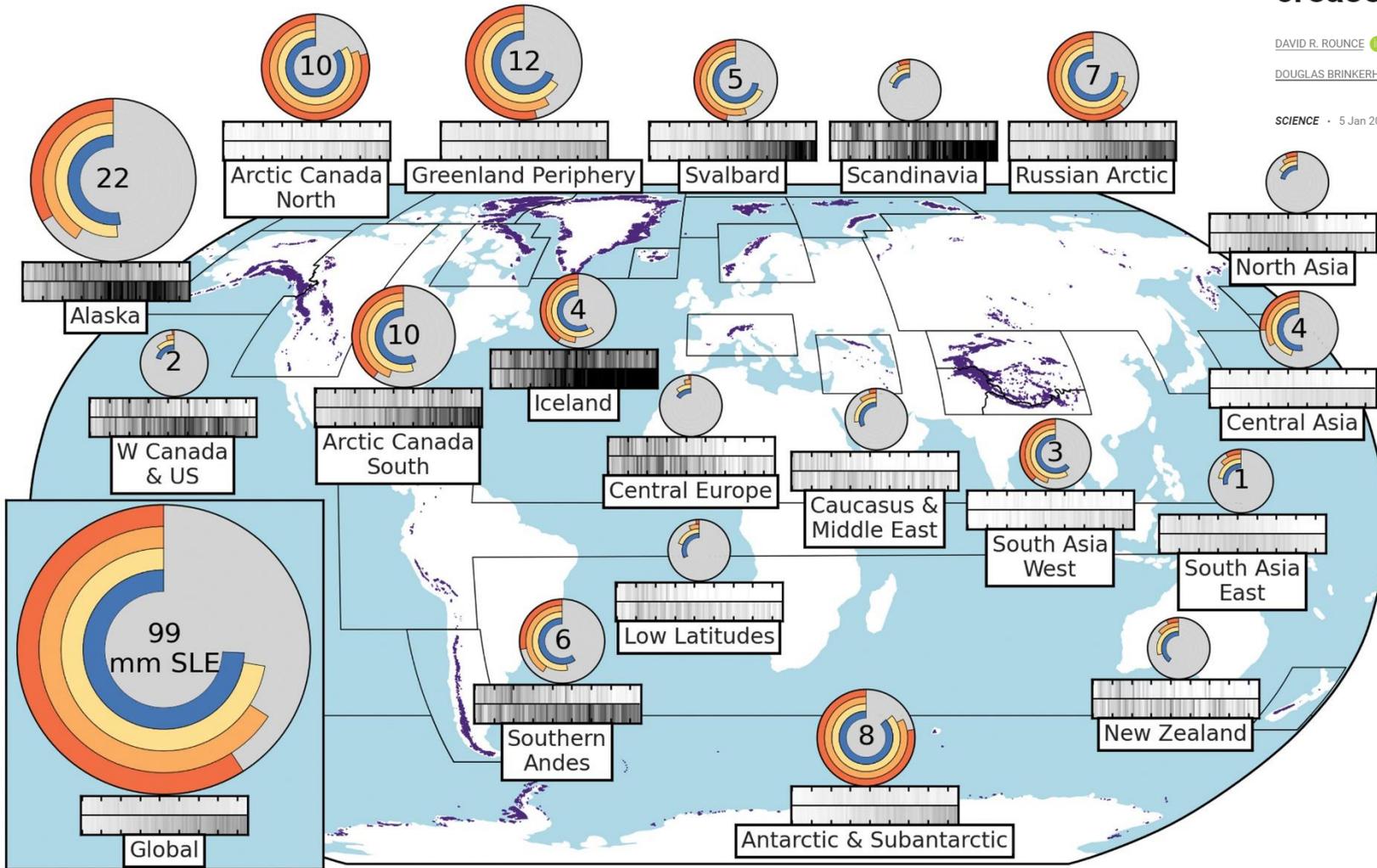
# Geodetic mass balance? What for?

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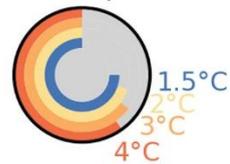
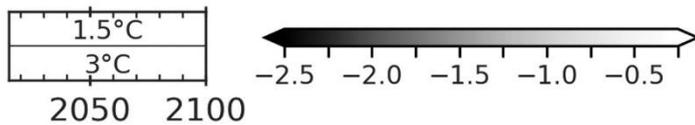
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Annual mass balance (m w.e.)

Mass at 2100 (rel. to 2015)



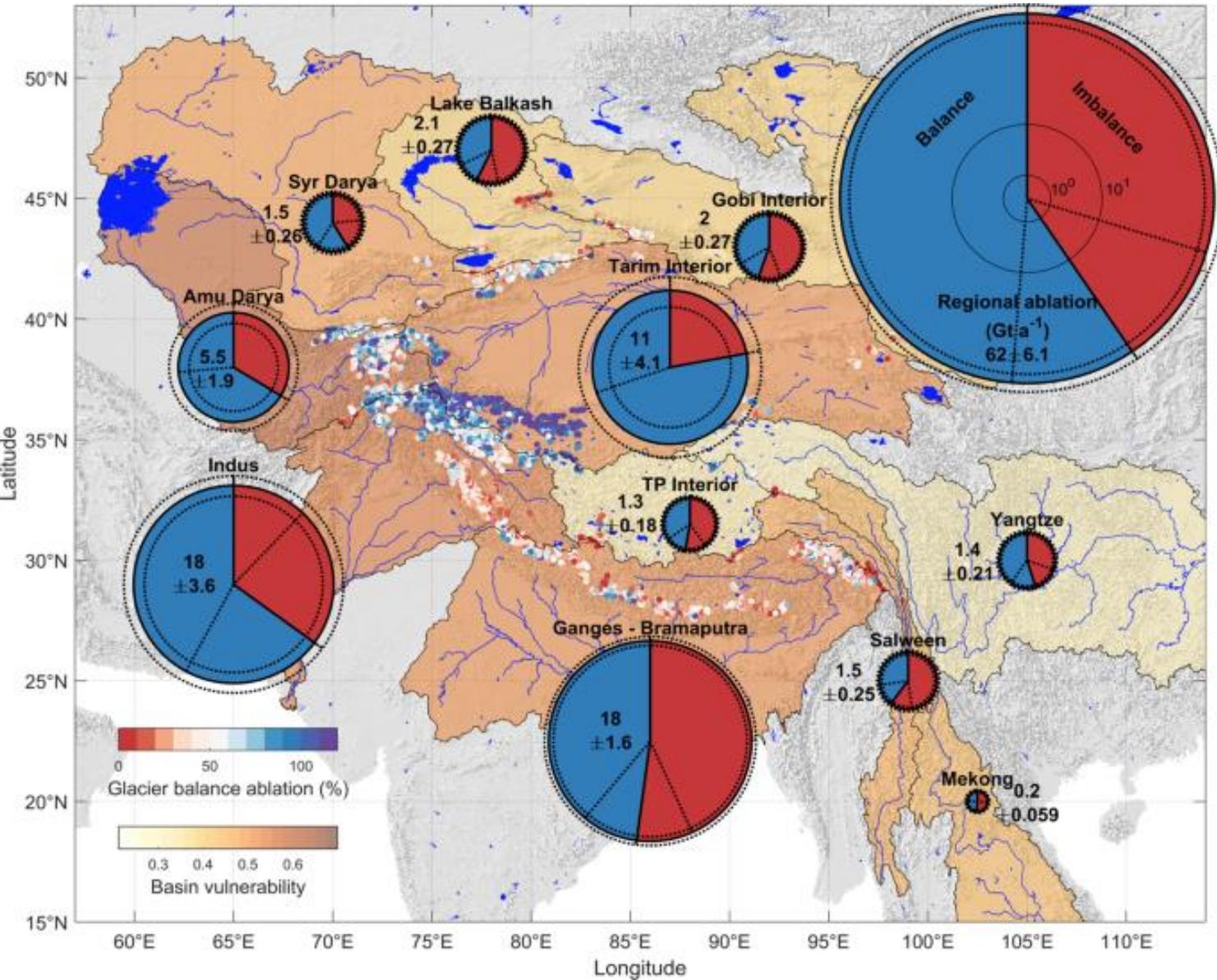
# Geodetic mass balance? What for?

## Health and sustainability of glaciers in High Mountain Asia

[Evan Miles](#) , [Michael McCarthy](#), [Amaury Dehecq](#), [Marin Kneib](#), [Stefan Fugger](#) & [Francesca Pellicciotti](#)

*Nature Communications* **12**, Article number: 2868 (2021) | [Cite this article](#)

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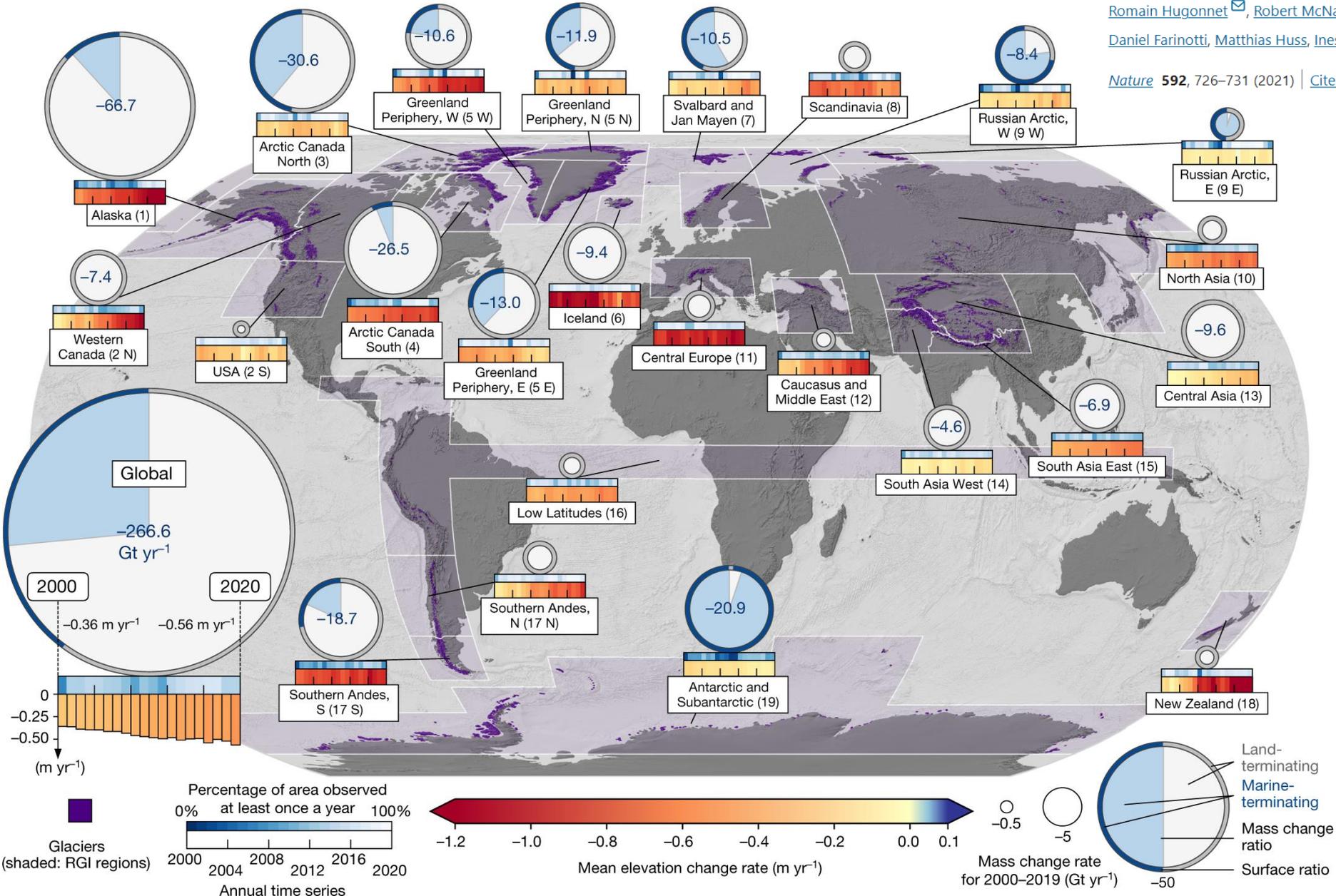
# Geodetic mass balance? What for?

## Accelerated global glacier mass loss in the early twenty-first century

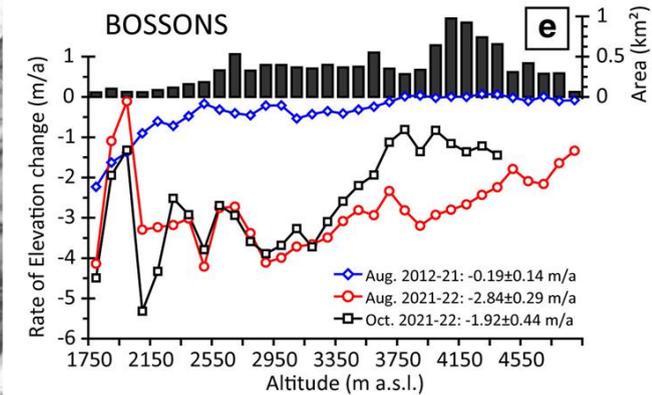
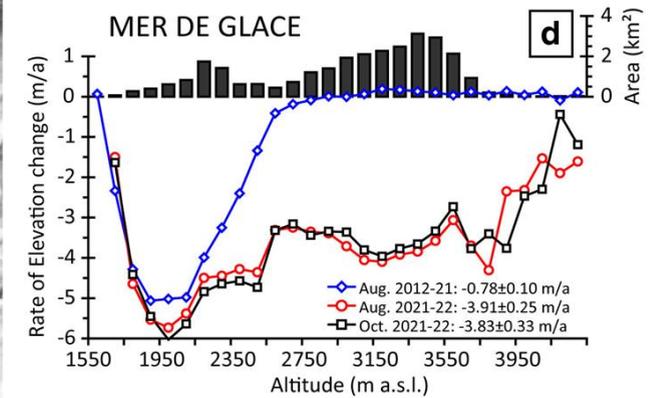
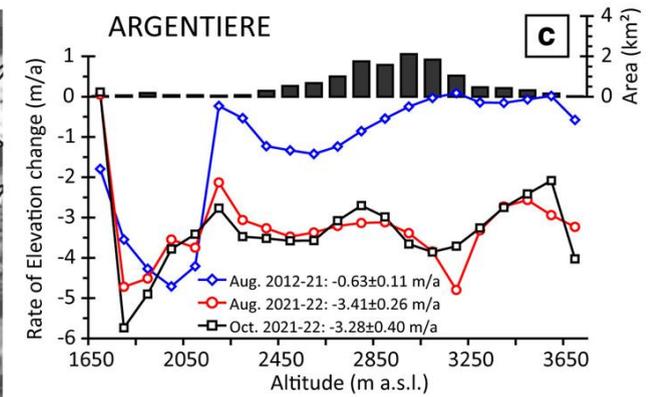
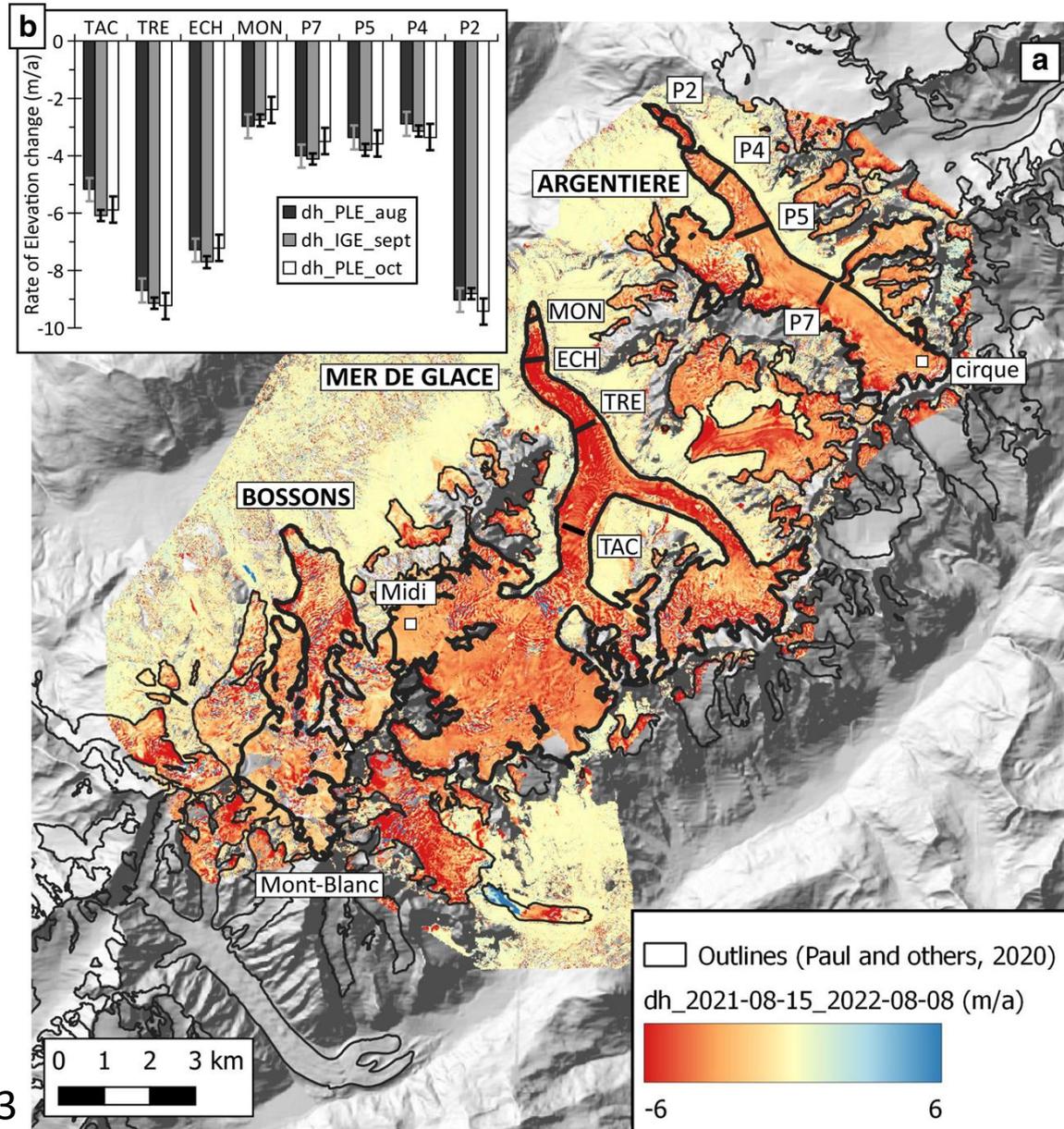
Romain Hugonnet , Robert McNabb, Etienne Berthier, Brian Menounos, Christopher Nuth, Luc Girod,

Daniel Farinotti, Matthias Huss, Ines Dussailant, Fanny Brun & Andreas Kääb

*Nature* 592, 726–731 (2021) | [Cite this article](#)



# Geodetic mass balance? What for?



# Geodetic mass balance? What for?

**Geodetic mass balance:** Mass change of glaciers calculated from a difference of digital elevation models acquired multi years apart

**Digital elevation model (DEM):** gridded representation of topography

# Principles of geodetic mass balance



# Measurements of mass/volume changes from space

## DEM differencing

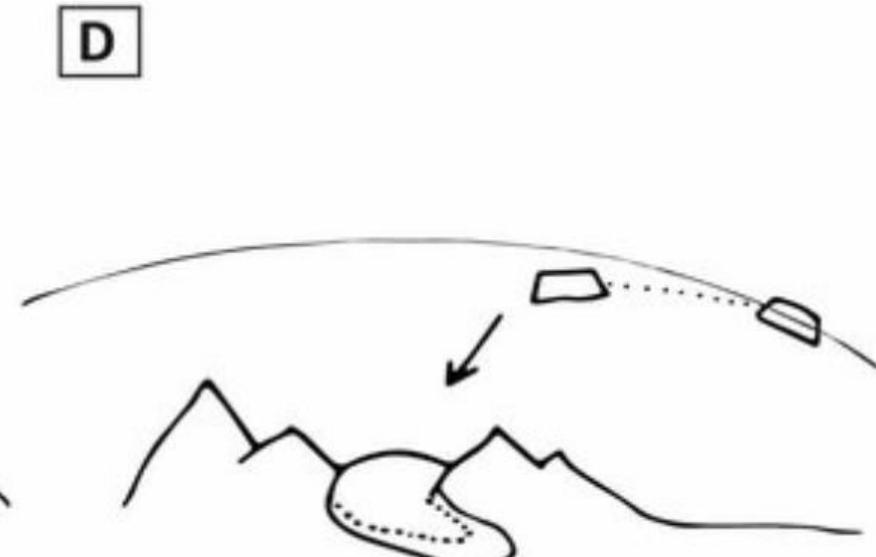
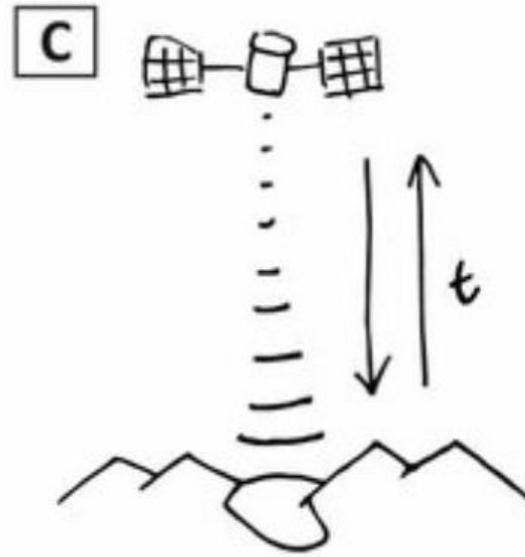
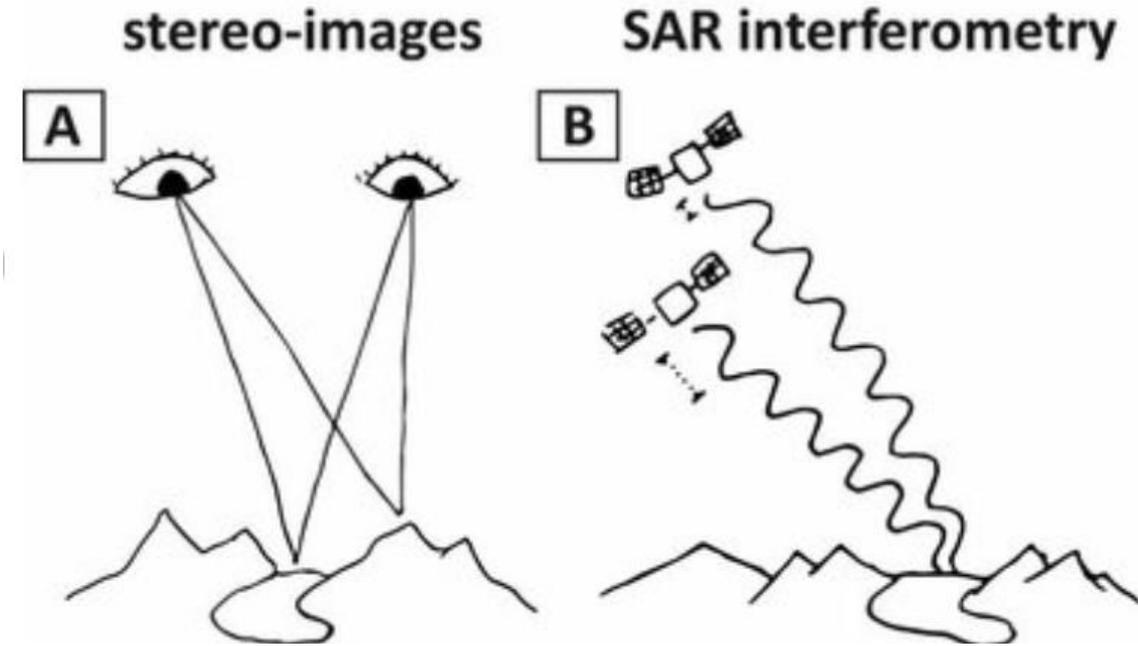
stereo-images

SAR interferometry

## Altimetry

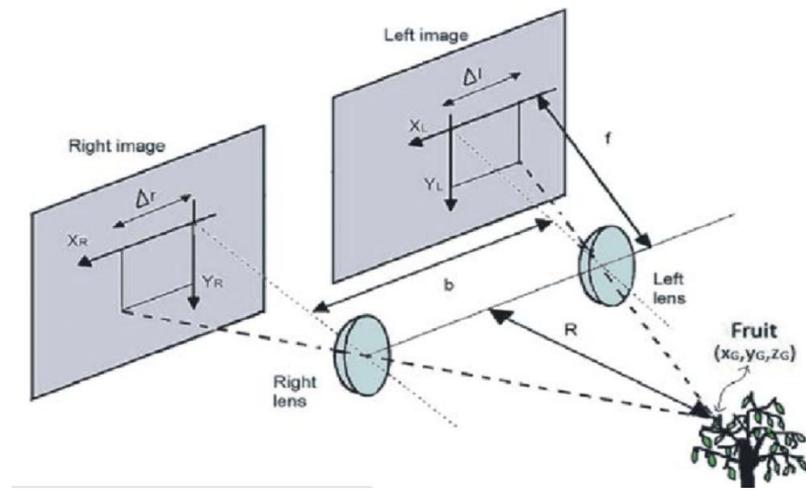
laser, radar

## Gravimetry



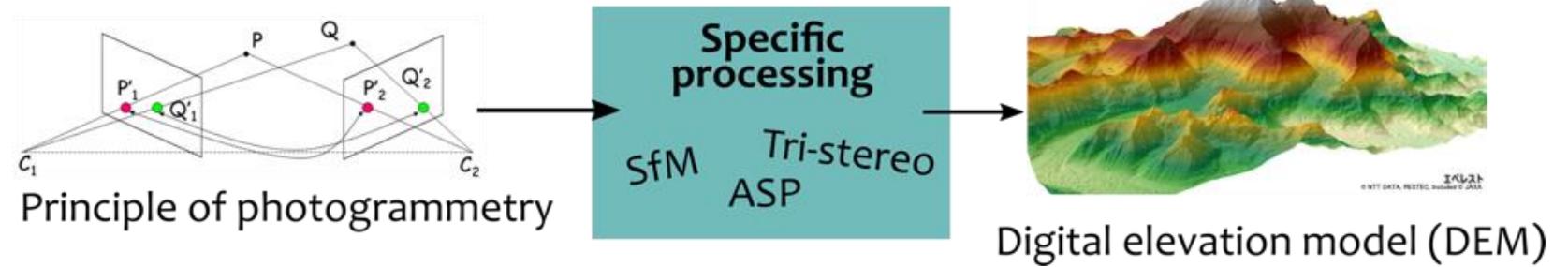
# Optical stereo data

- Advantages:
  - High spatial resolution (up to 50 cm of ground sampling distance)
  - Based on images in visible and near infra-red -> easy to interpret by humans



- Drawbacks:
  - Obscured by clouds and polar night. Some problems in the steep faces in the shade.
  - Sensor specific problems (e.g. jitter, undulations, tilt...)

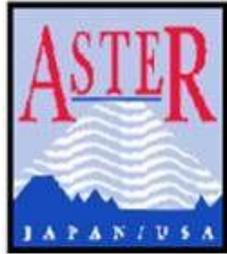
- Some examples:
  - ASTER
  - Pléiades
  - KH9
  - Worldview (HMA DEM)



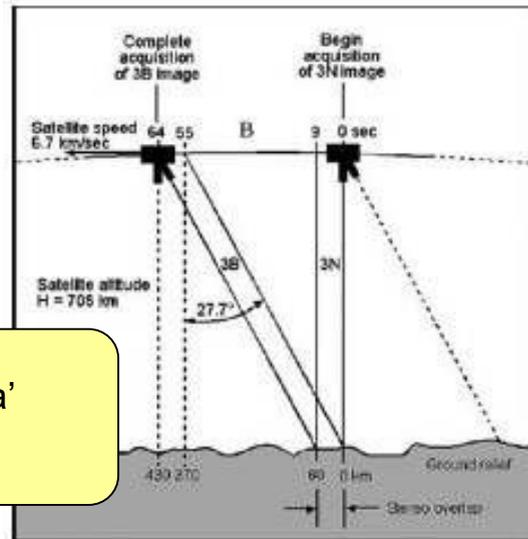
- Processing:
  - 2-3 images
  - Stereo-processing
  - Software: AMES Stereo Pipeline (ASP), Erdas Imagine, Agisoft Metashape...



# 2 sources of stereo-images/DEMs



'public data' sensor



ASTER 3N & 3B, 15 m  
Along track stereo using two telescopes

↓  
30 m DEM  
generated using the Ames Stereo Pipeline (ASP)



Commercial satellites

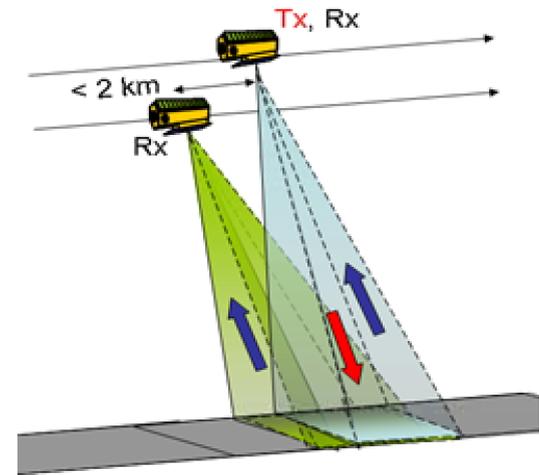
Pléiades 1A & 1B, 0.5 m Along-track stereo satellite agility

↓  
DEM generation using the Ames Stereo Pipeline

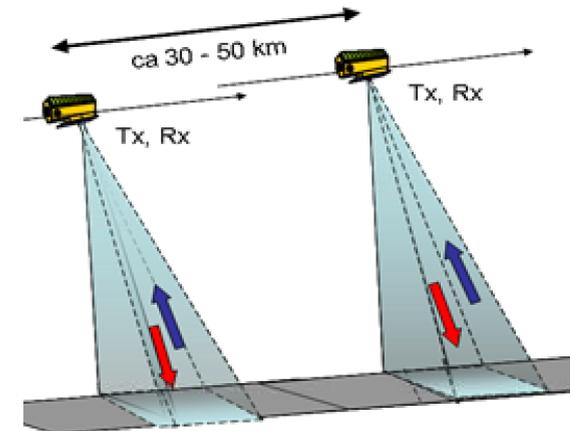
# InSAR data

- Advantages:
  - Insensitive to clouds and night
- Drawbacks:
  - Penetration of the radar signal into snow and ice
  - Potential large errors in steep topography
- Some examples:
  - SRTM
  - TanDEM-X

a) bistatic mode



b) pursuit nonostatic node

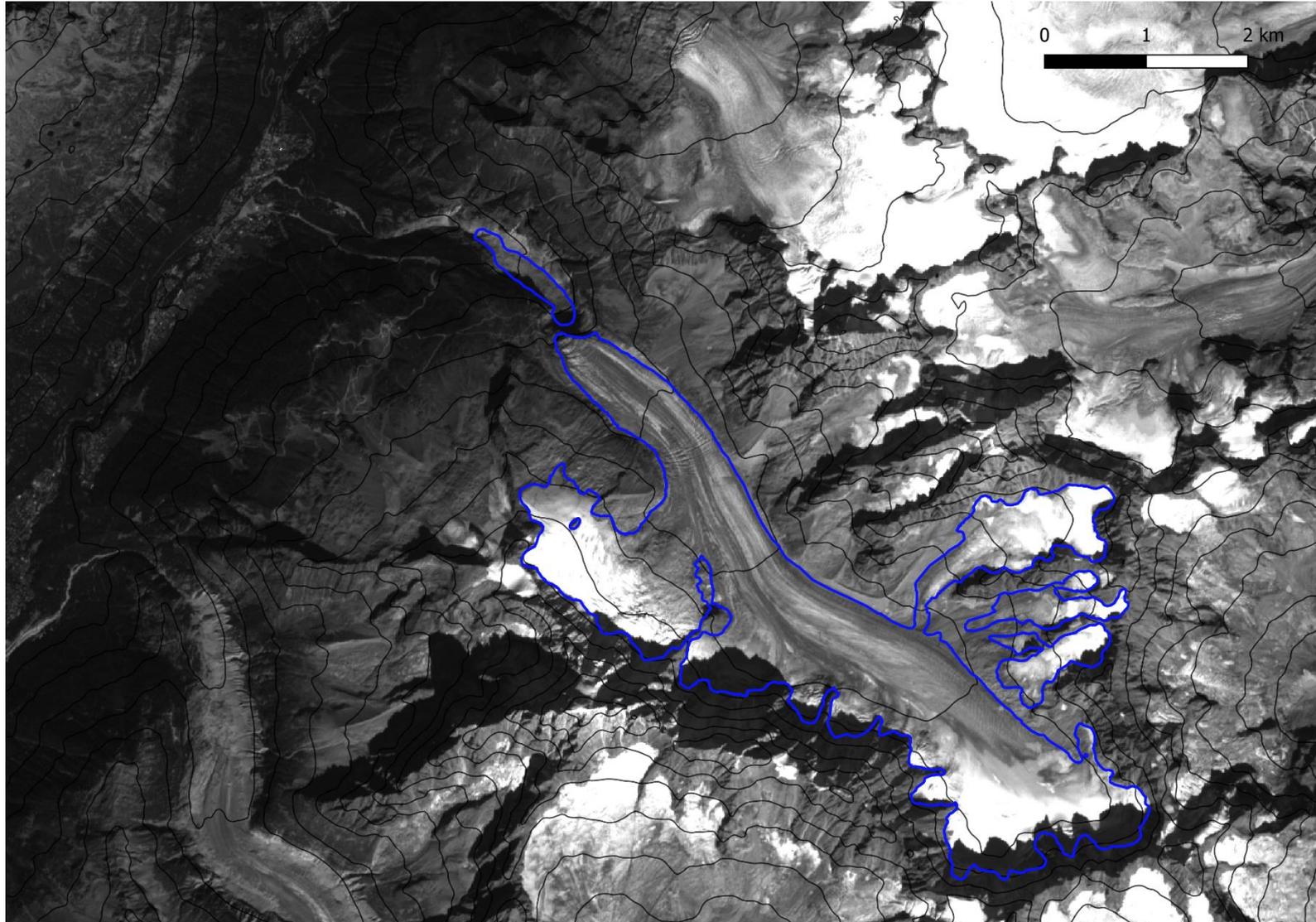


# UAV data

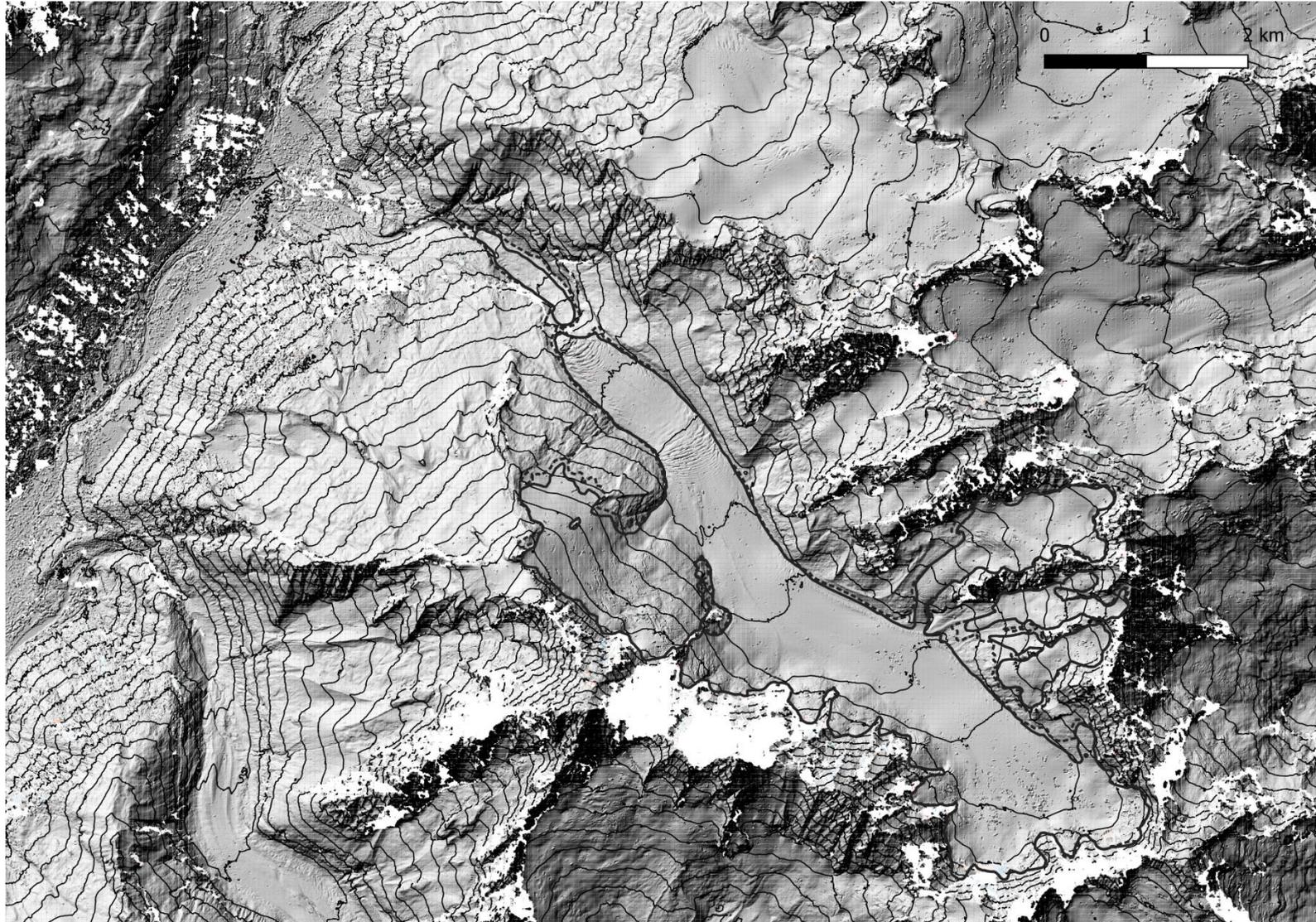
- Advantages:
  - Very high resolution (<10 cm)
  - Constrained only by conditions in the field
  - Possibility of repeat flights
- Drawbacks:
  - Limited spatial coverage
  - Needs some ground control to constrain geometry
- Some examples:
  - Quadcopters
  - Fixed wings/VTOL drones
- Processing:
  - > 100 images
  - Structure-from-motion photogrammetry
  - Software: Agisoft Metashape, Pix4d...



# Example: Argentière Glacier



# Example: Argentière Glacier

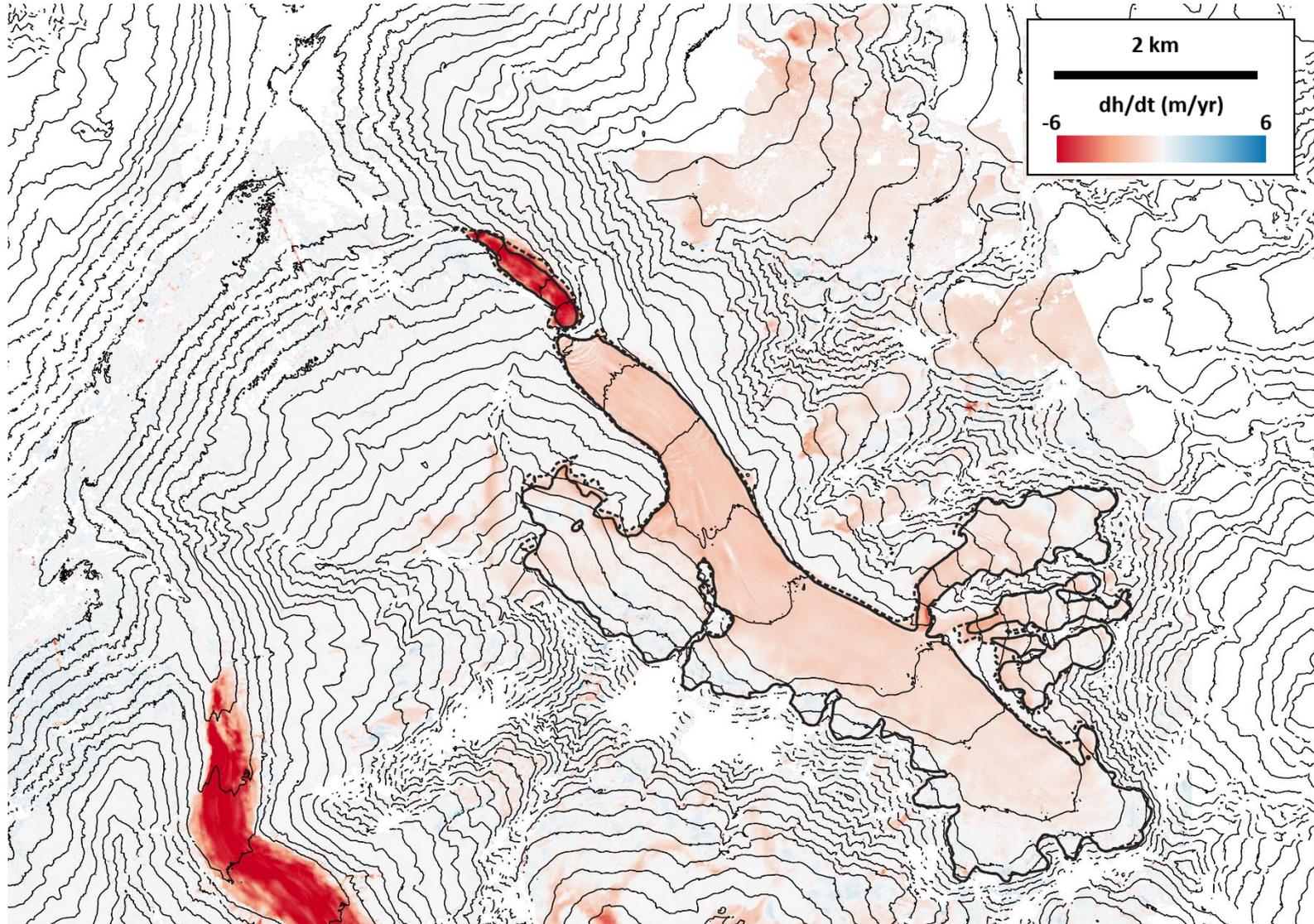


Geodetic mass balance (2012-2021):

# Example: Argentière Glacier

$dh = -0.85 \pm 0.07 \text{ m/yr}$

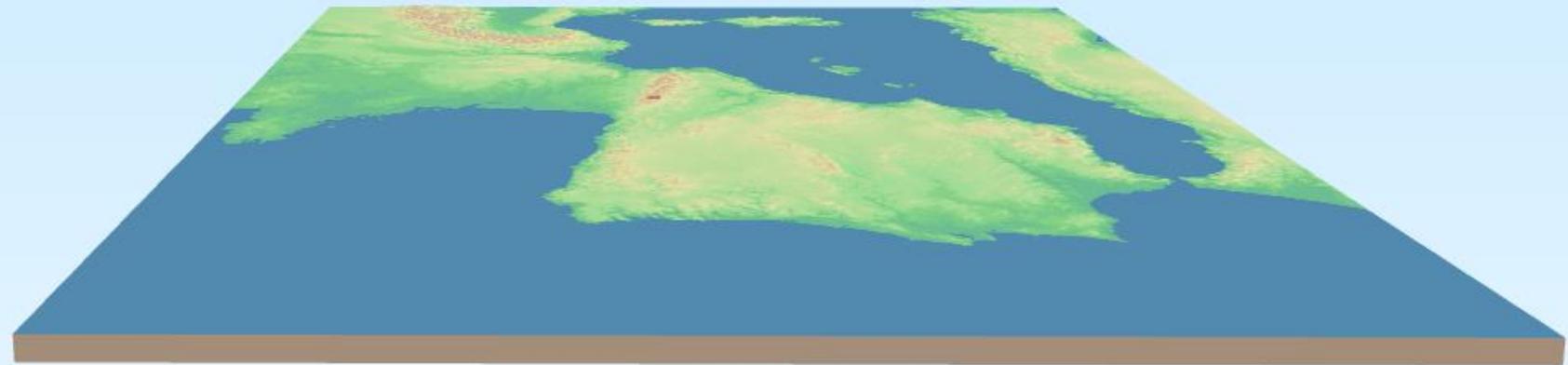
$dh_{we} = -0.72 \pm 0.08 \text{ m w.e./yr}$



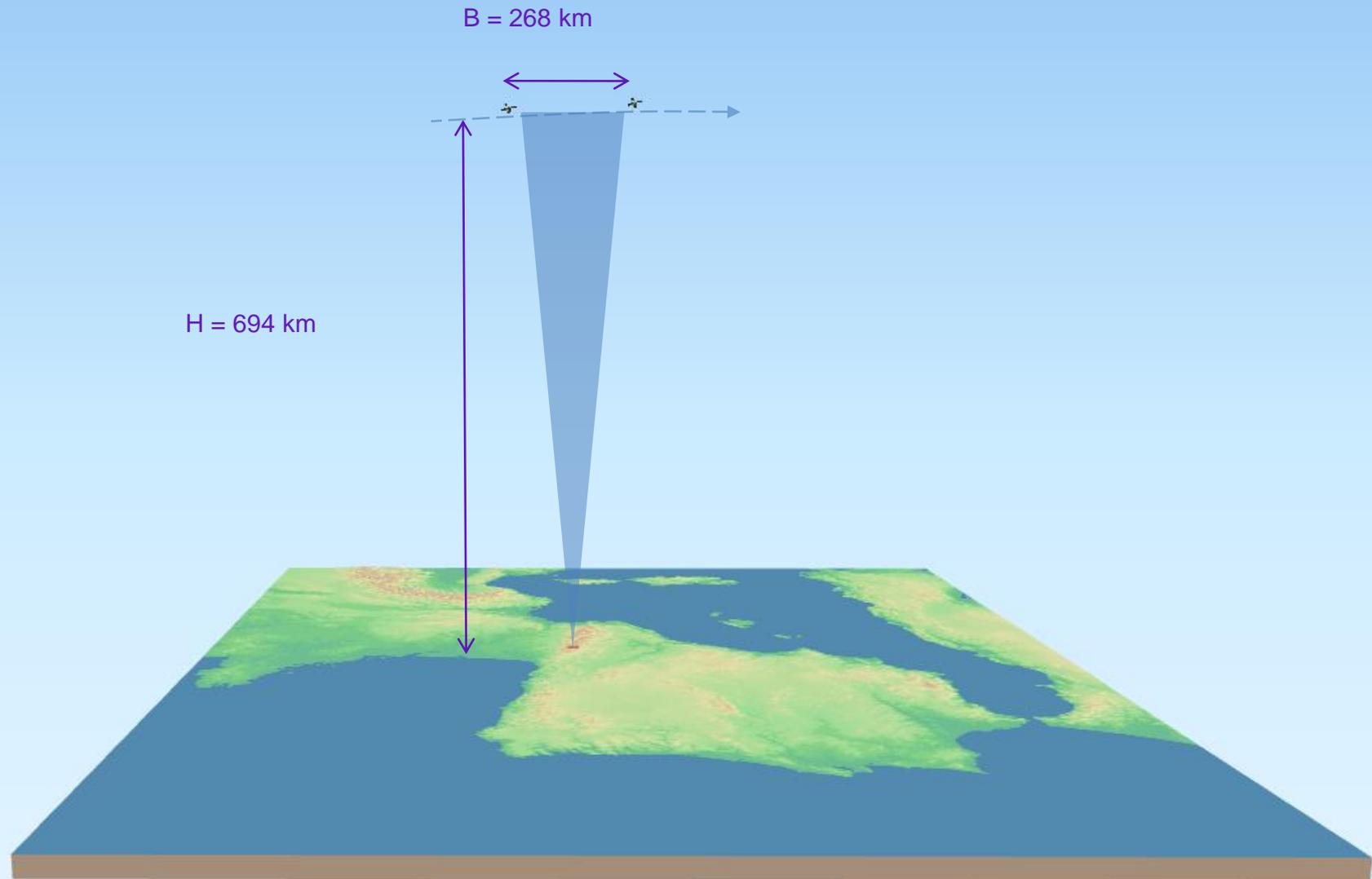
# Stereo-processing: a brief overview



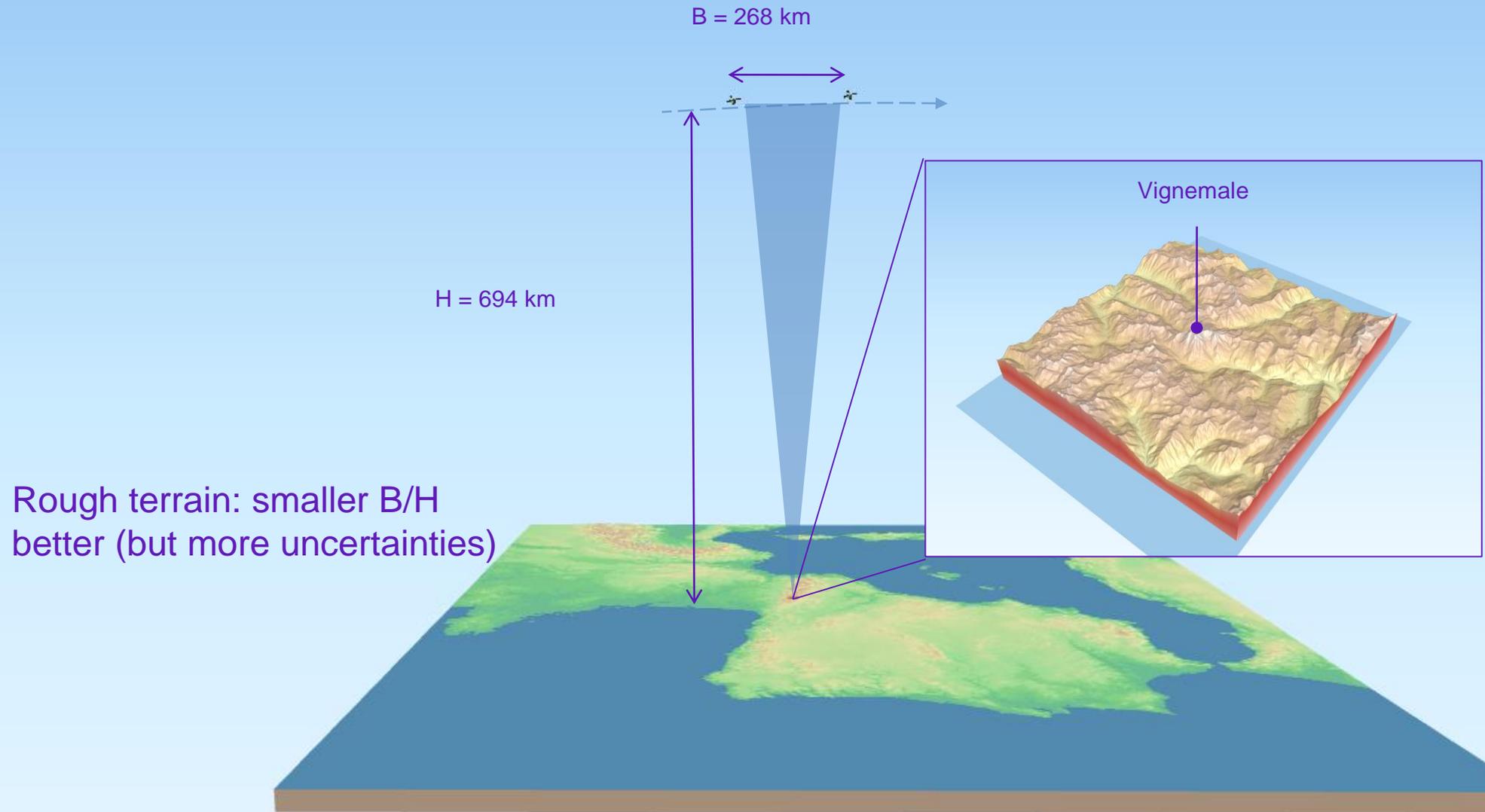
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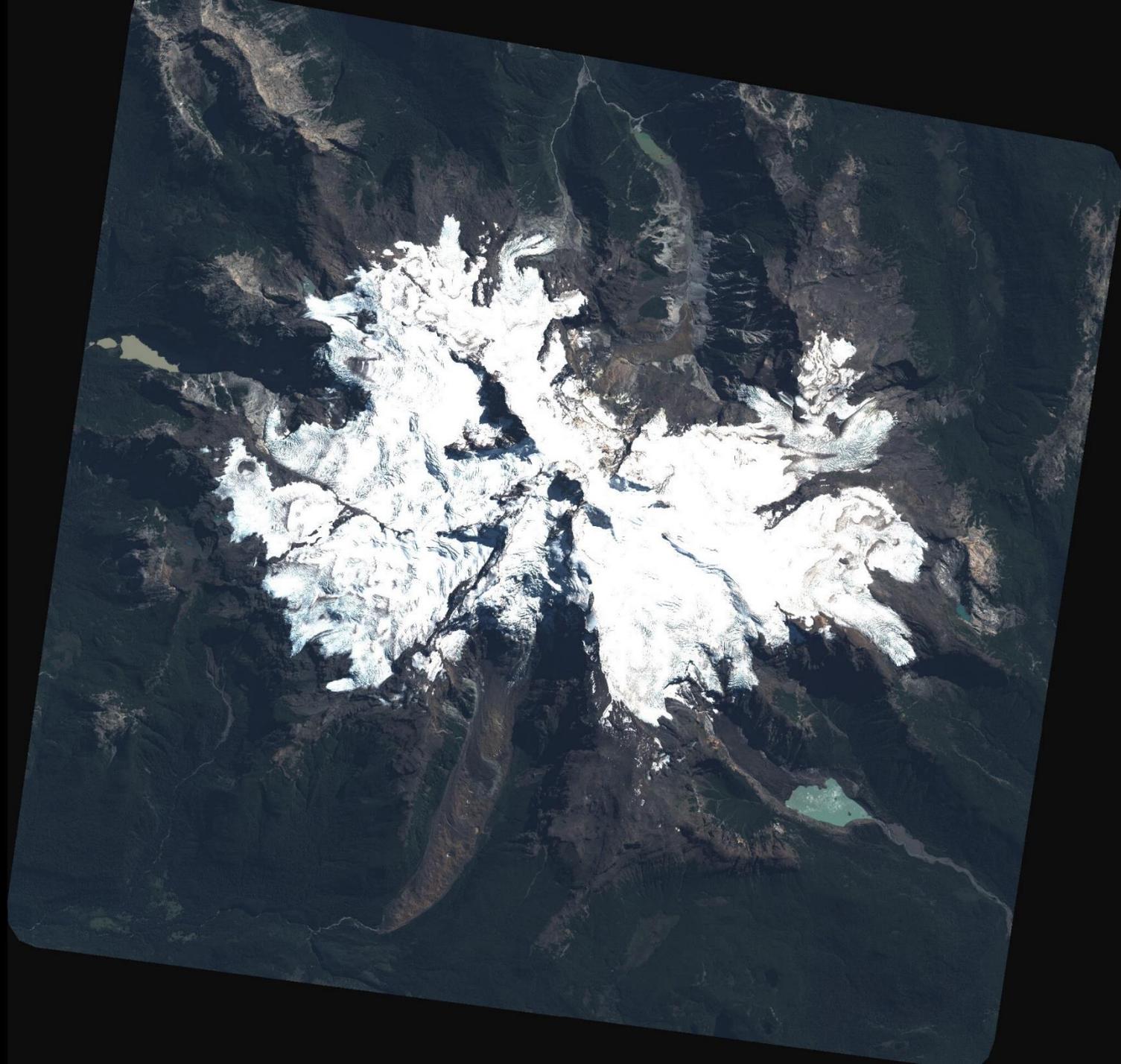
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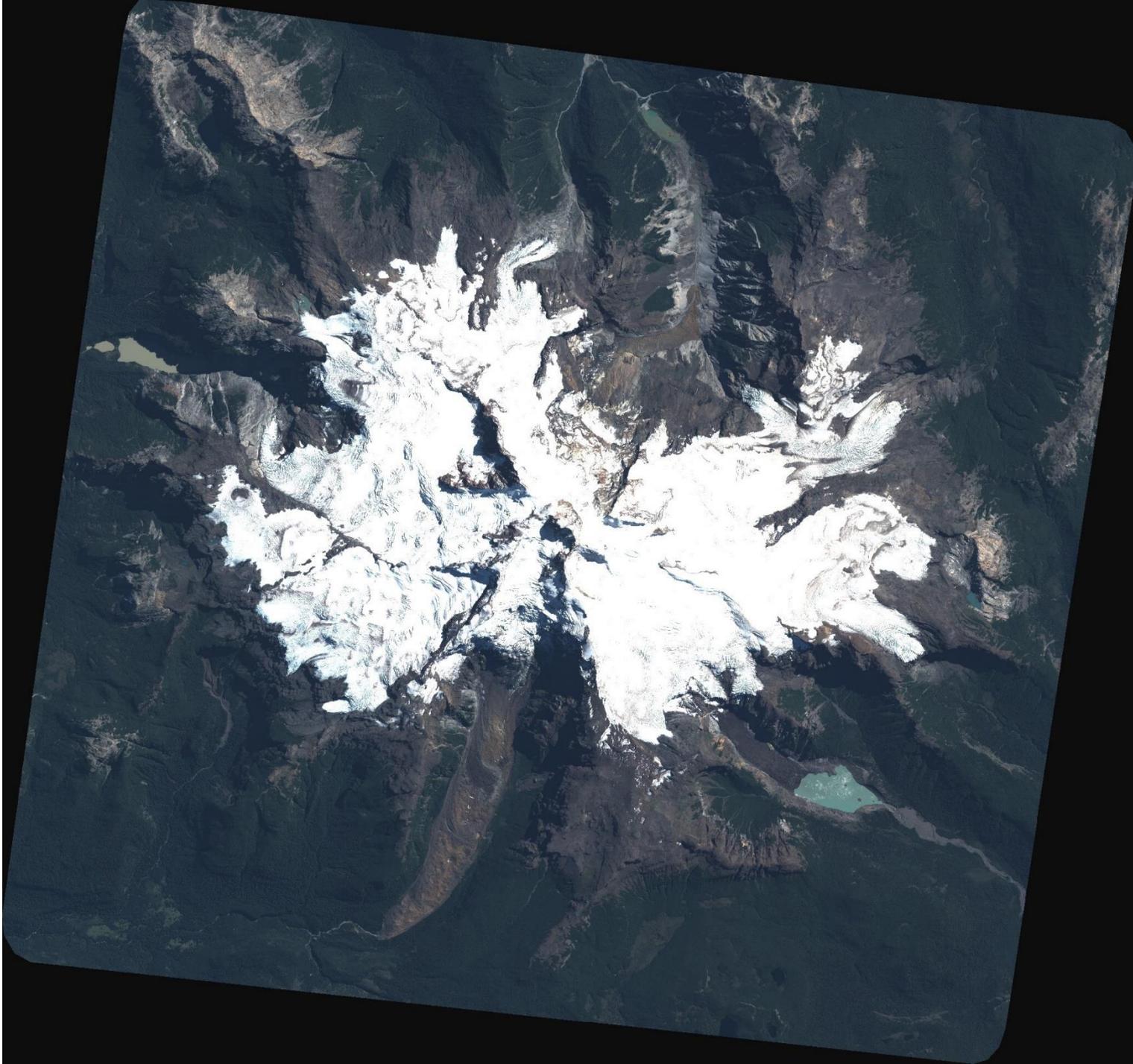
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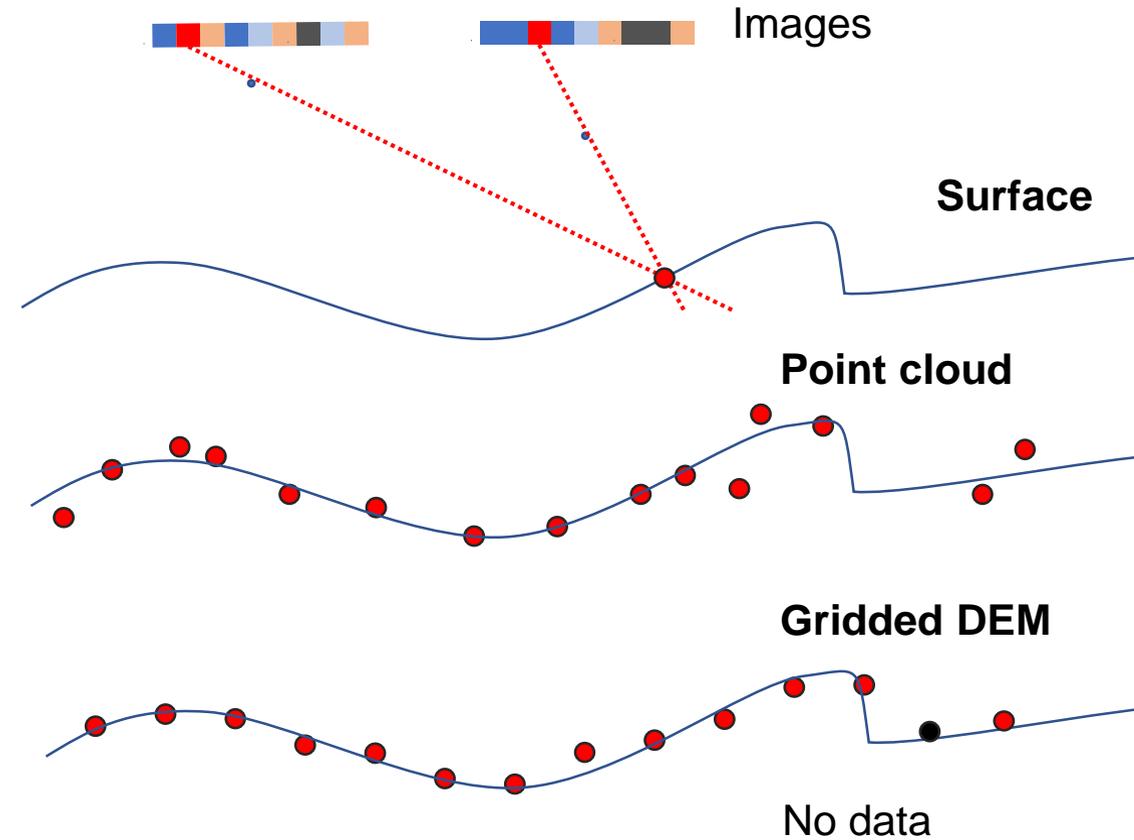
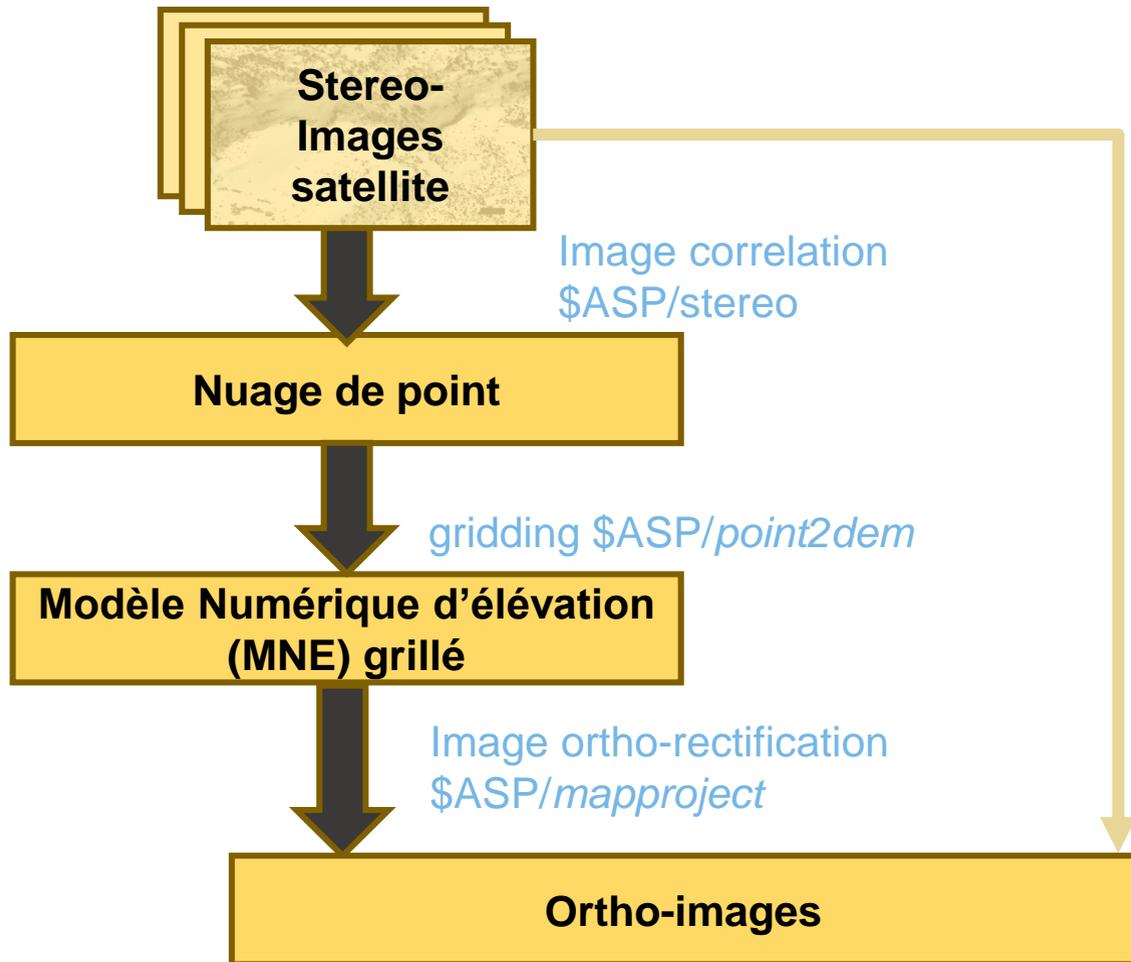
# Stereo images



# Stereo images



# Stereo-processing: a brief overview



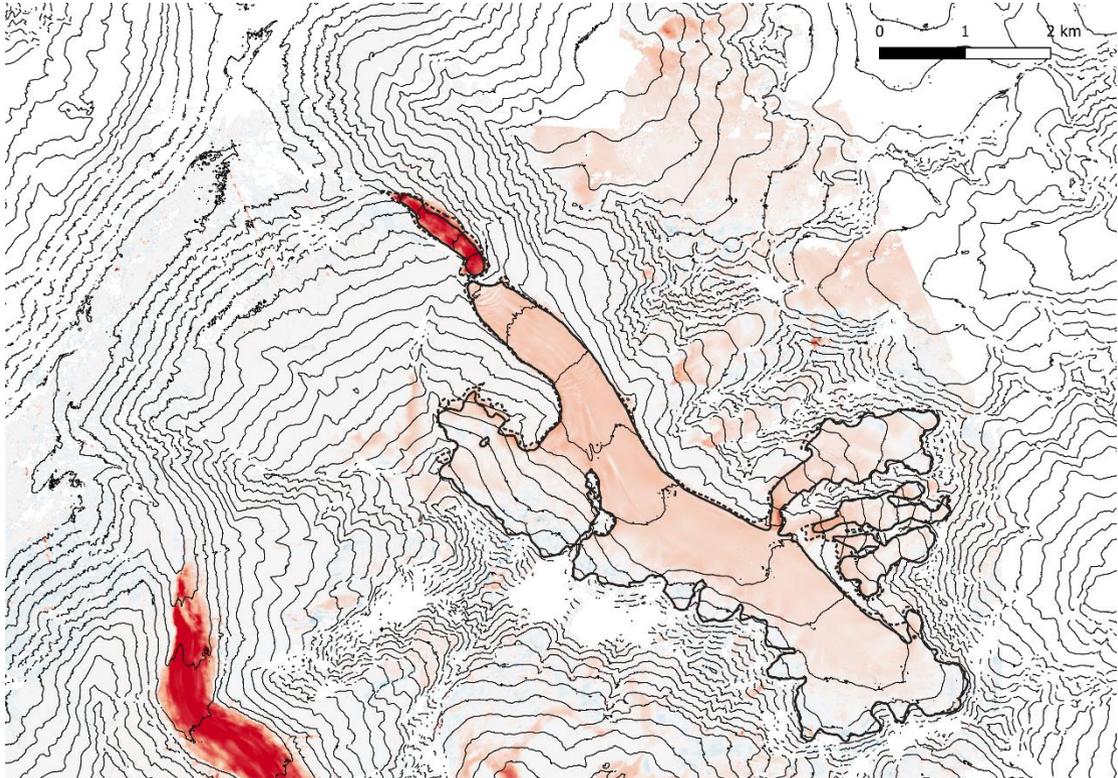
Recommended software: AMES stereo-pipeline

# DEM differencing & co-registration

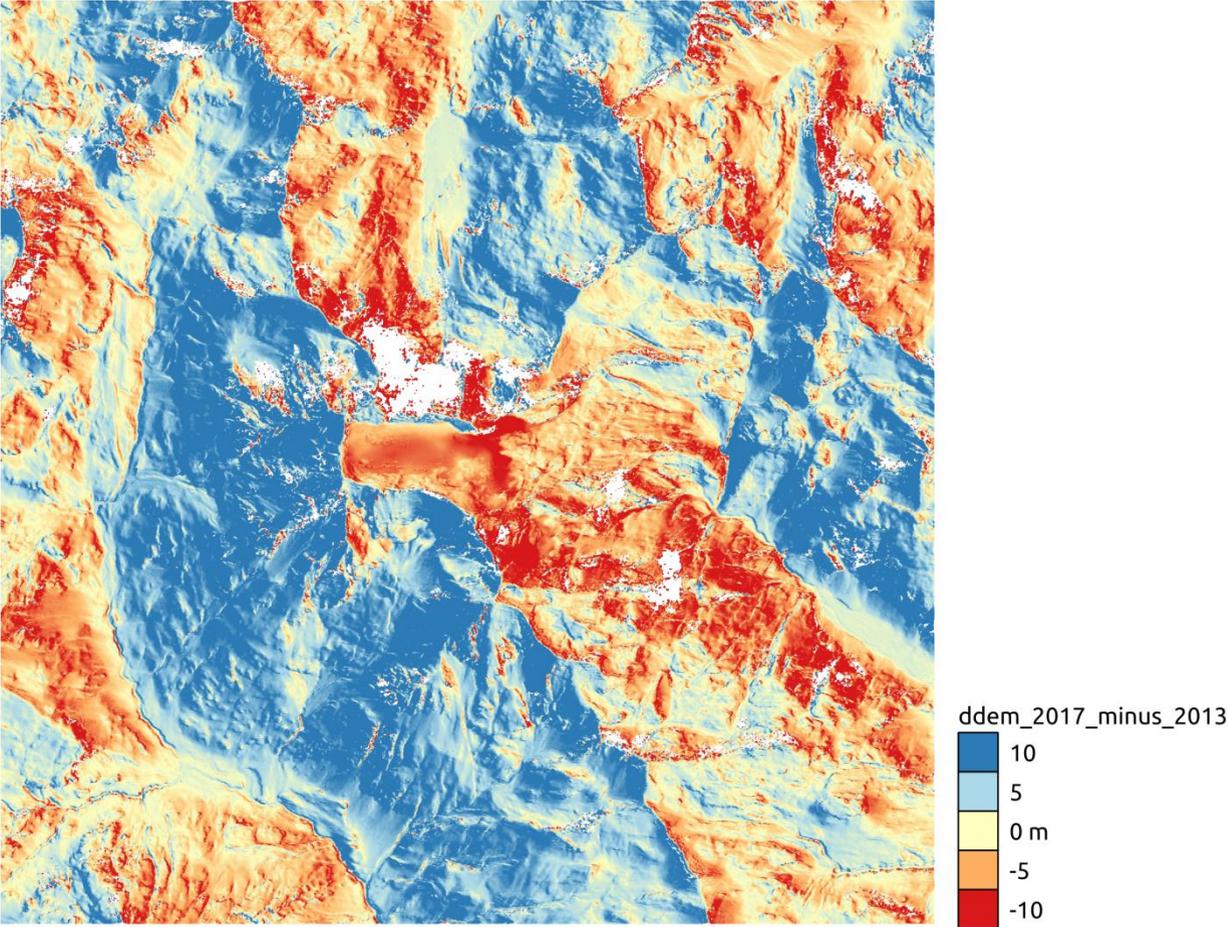


# DEM differencing & co-registration

What we want:

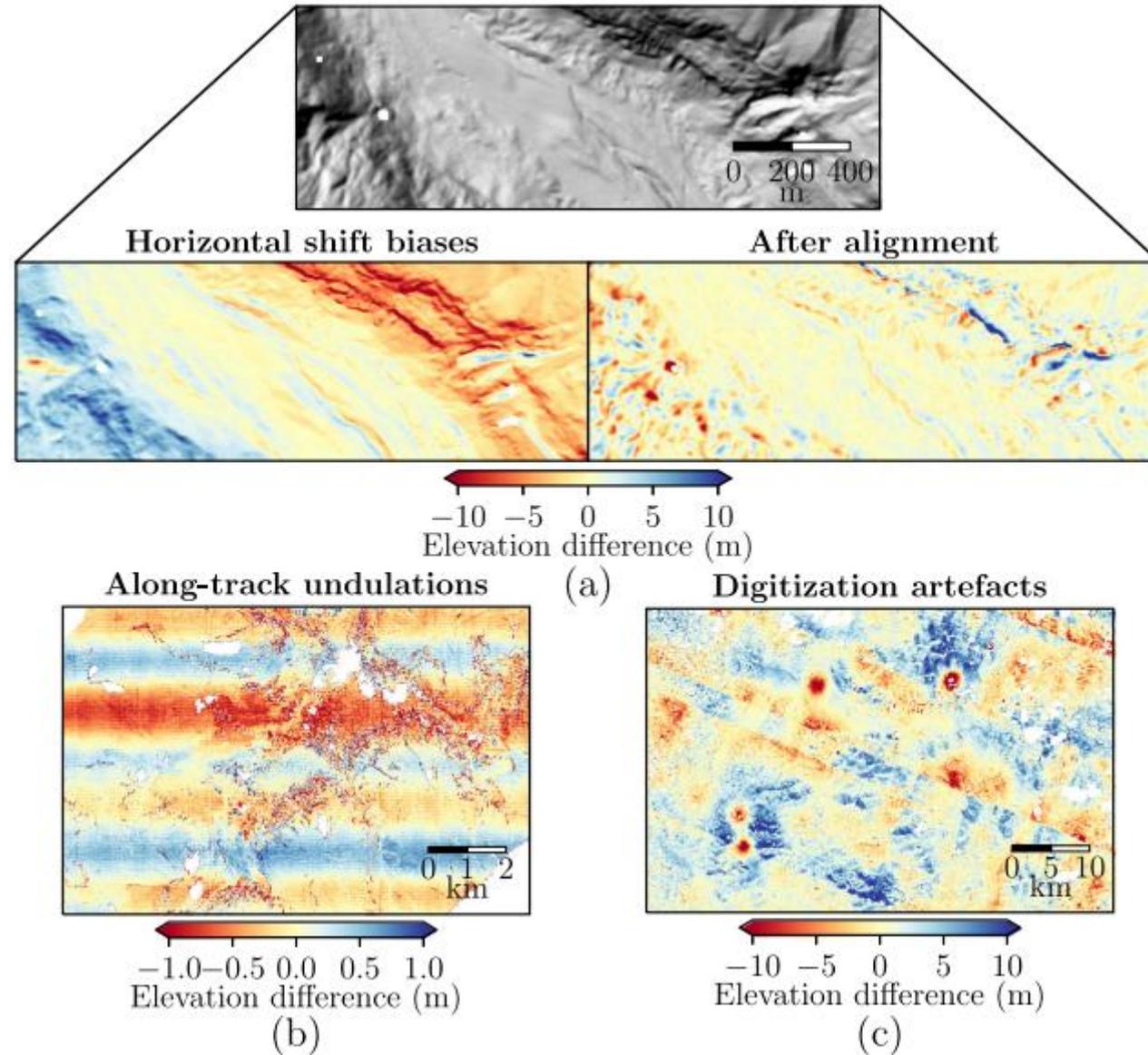


What we get:



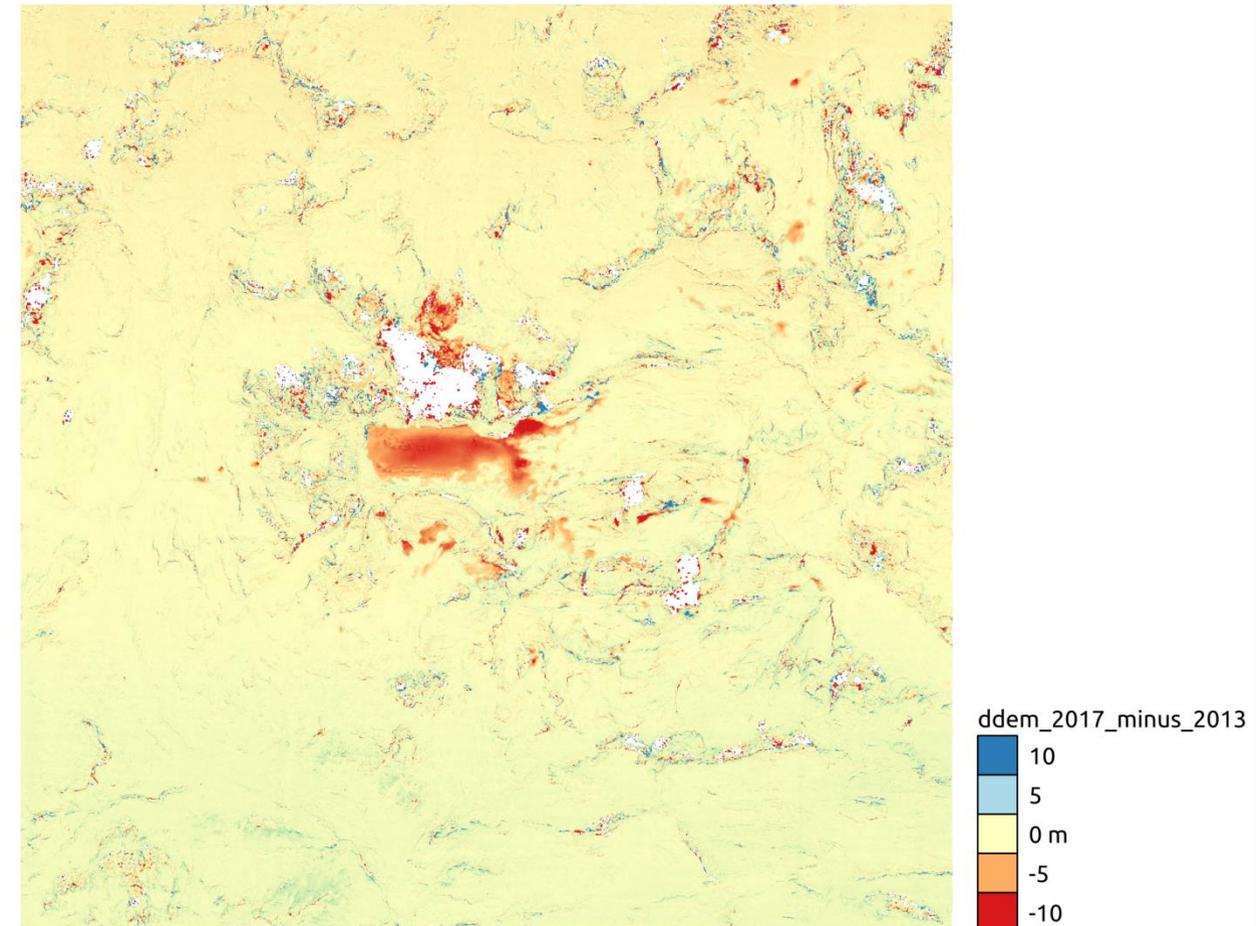
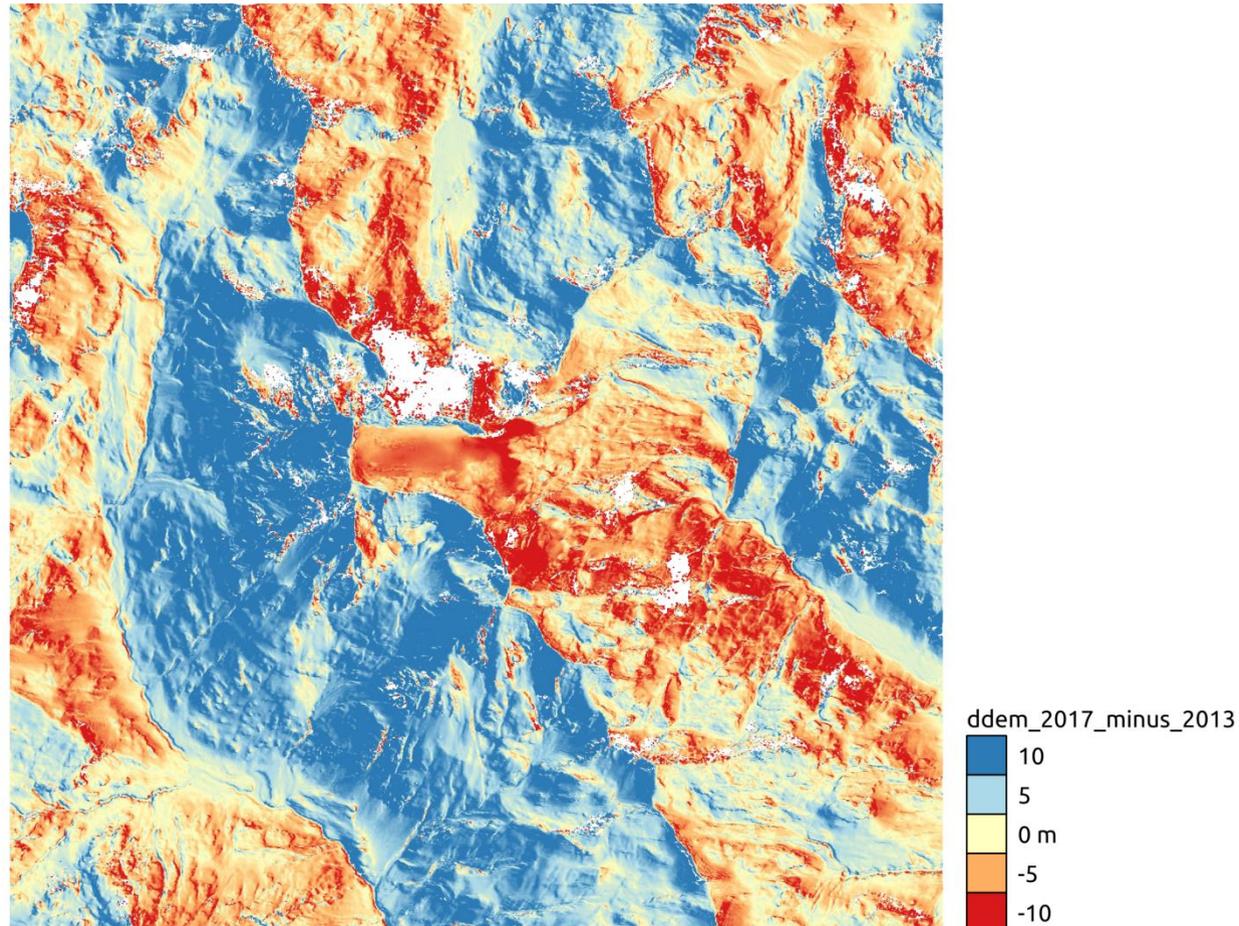
DEM difference on Ossoue Glacier (Gascoin et René 2018)

# DEM differencing & co-registration



# DEM differencing & co-registration

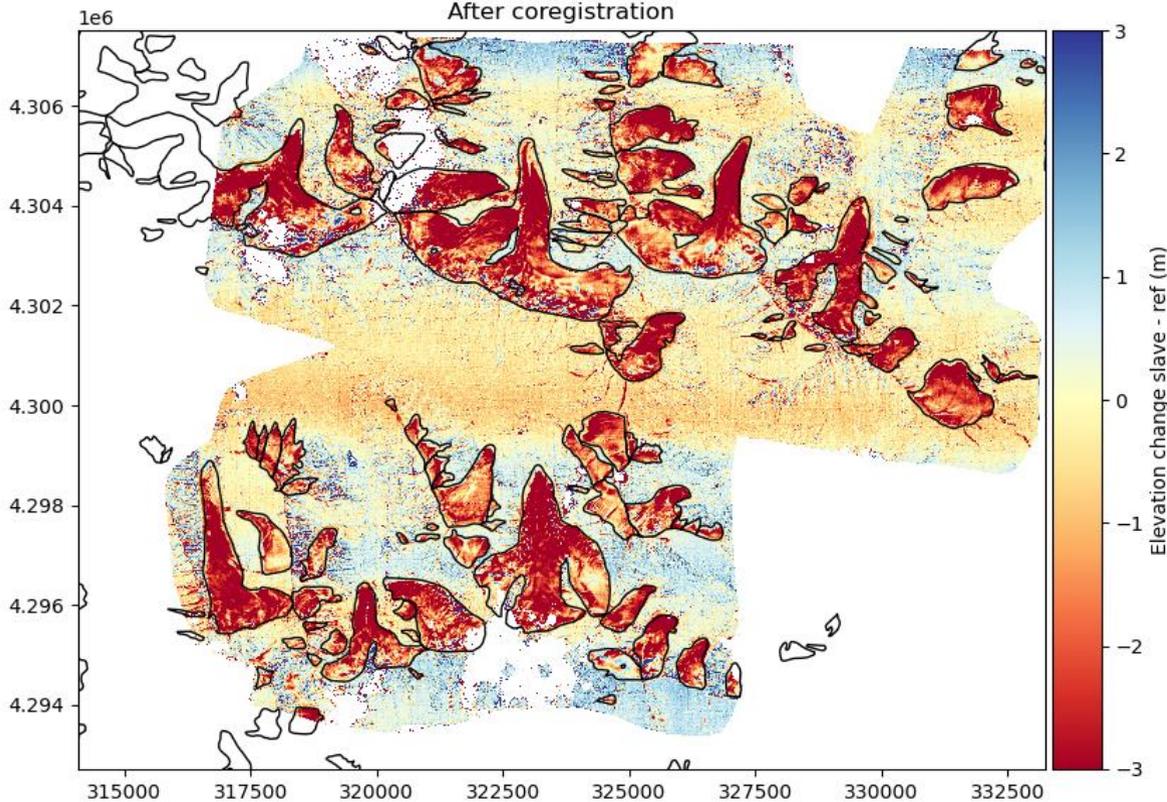
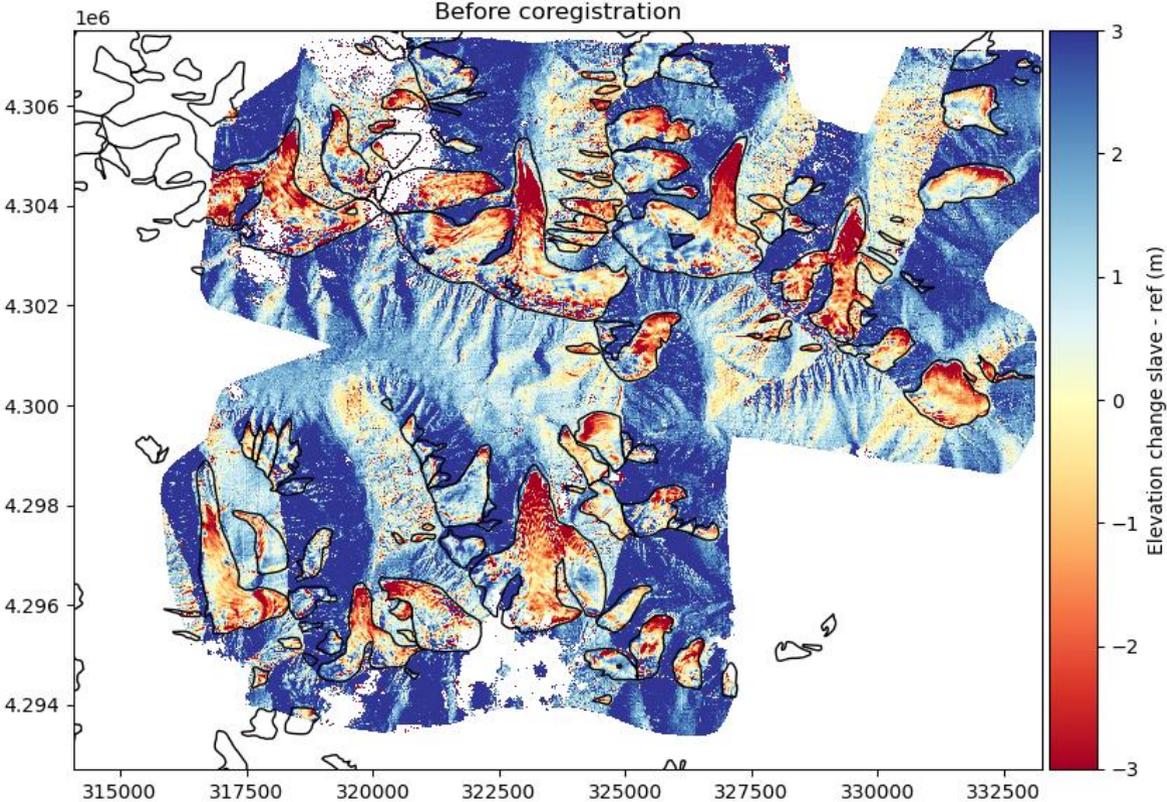
Idea: use stable terrain to co-register DEMs



DEM difference on Ossoue Glacier (Gascoin et René 2018)

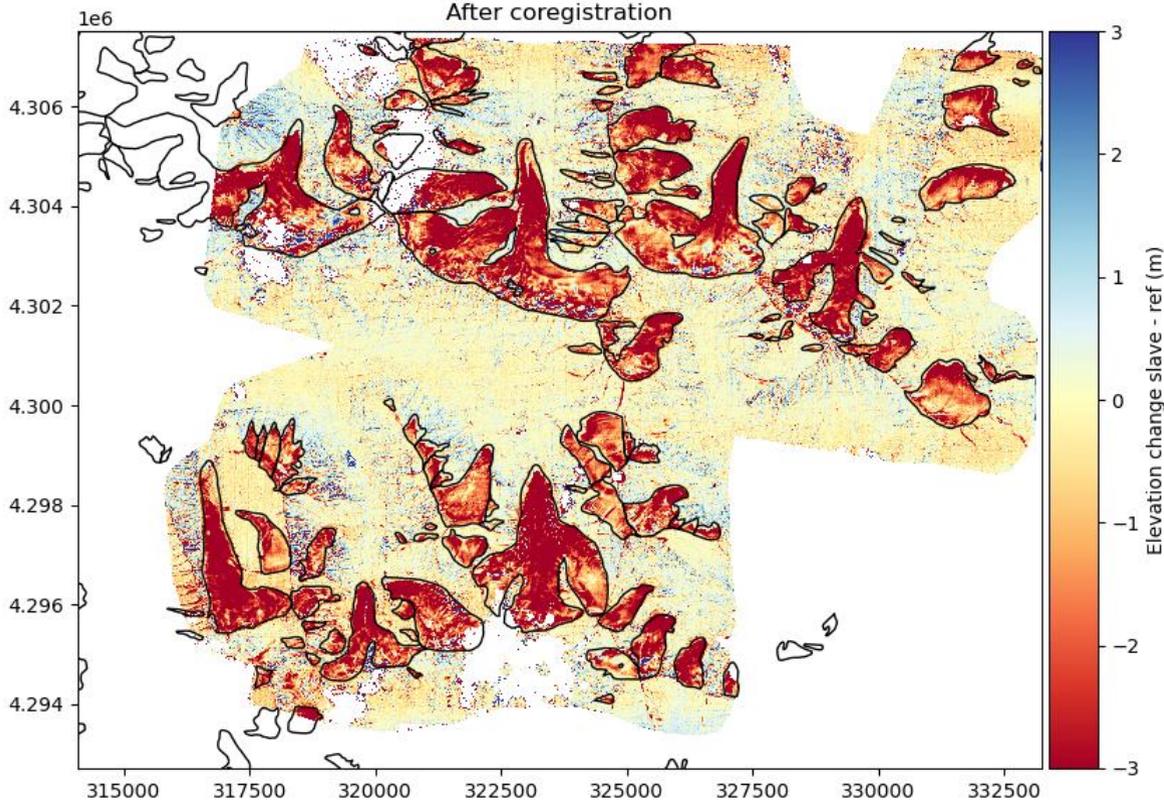
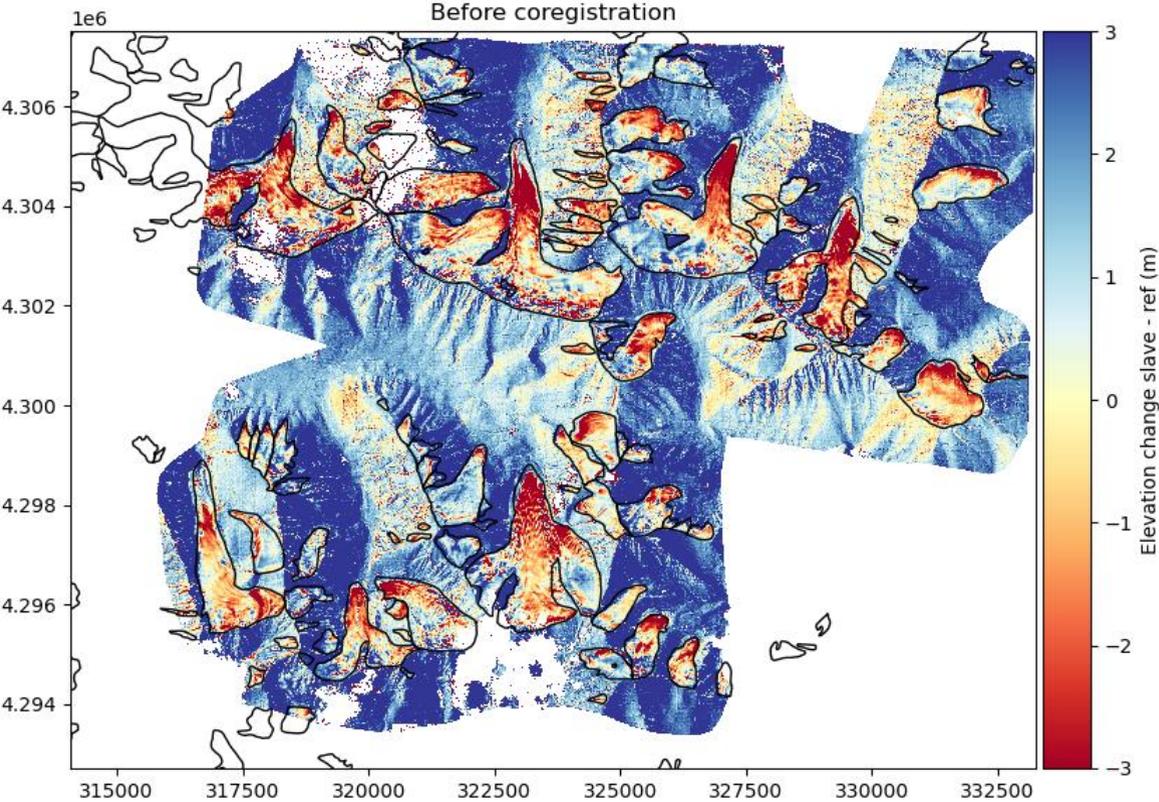
# DEM differencing & co-registration

08/2019-08/2024, Tajikistan

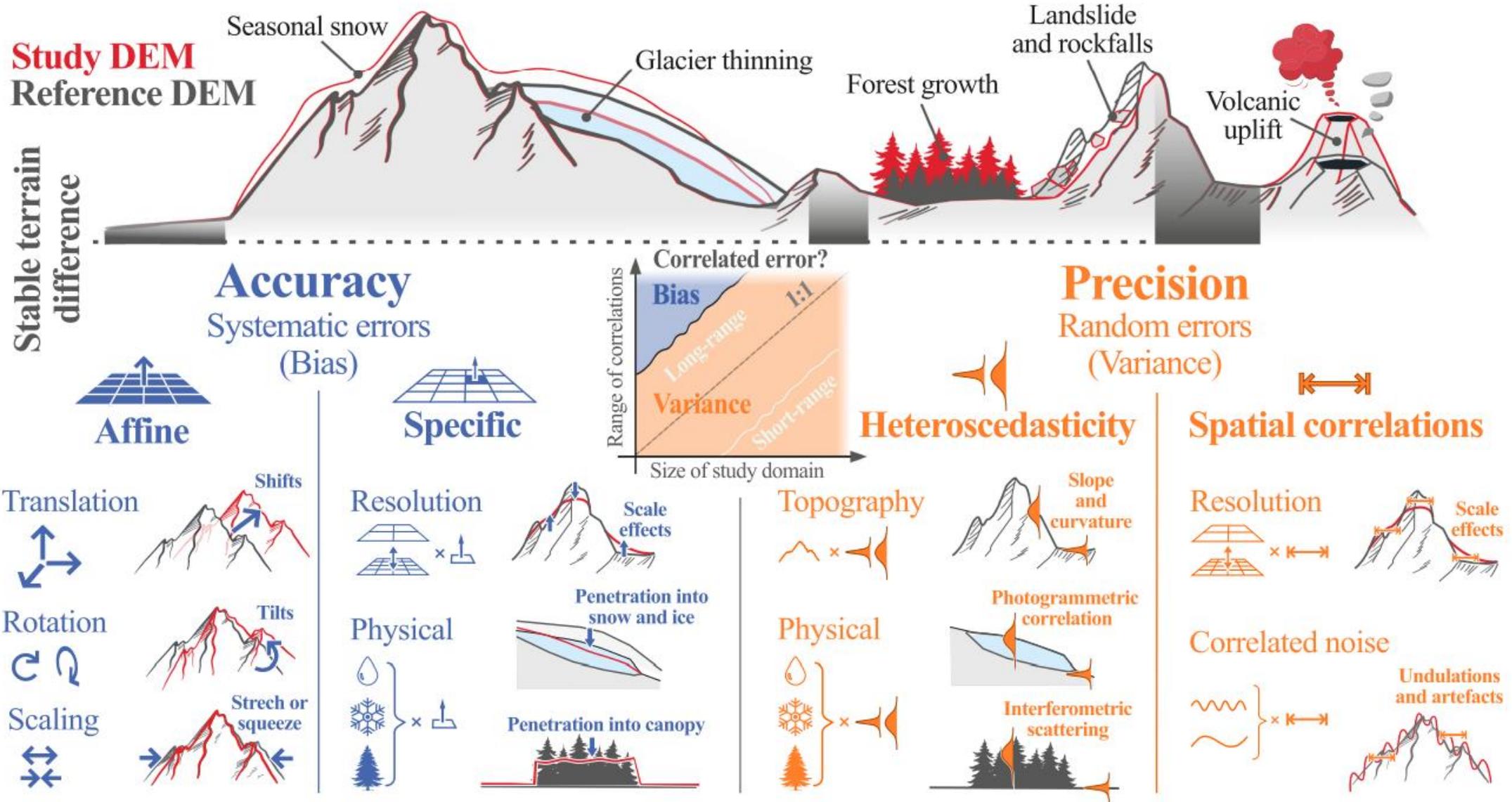


# DEM differencing & co-registration

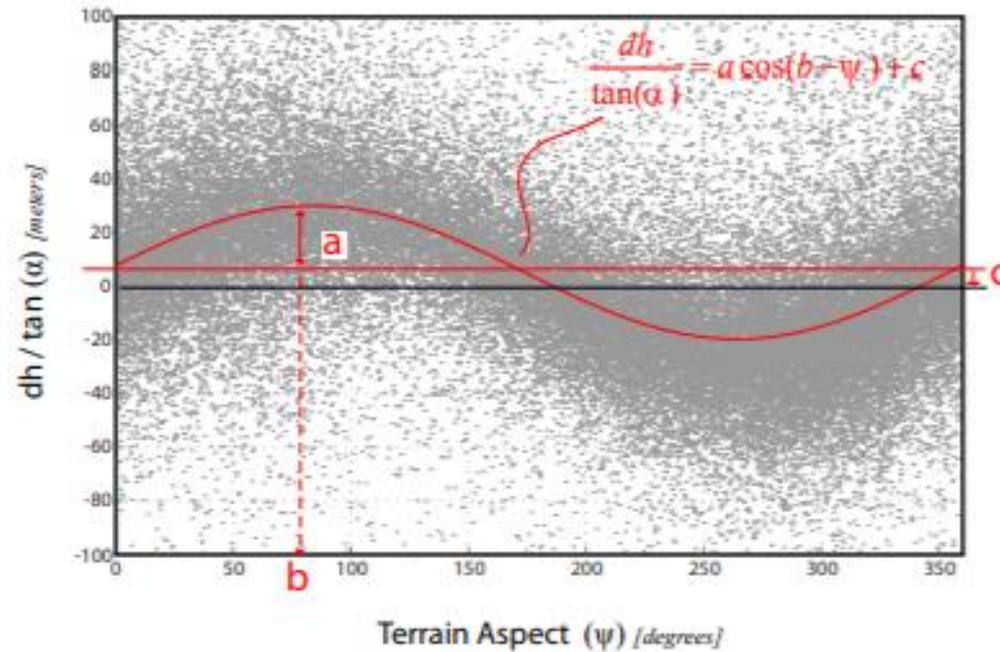
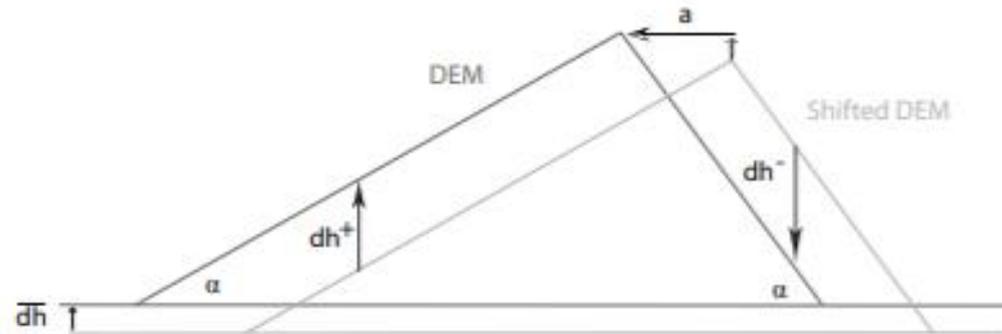
08/2019-08/2024, Tajikistan



# DEM differencing & co-registration

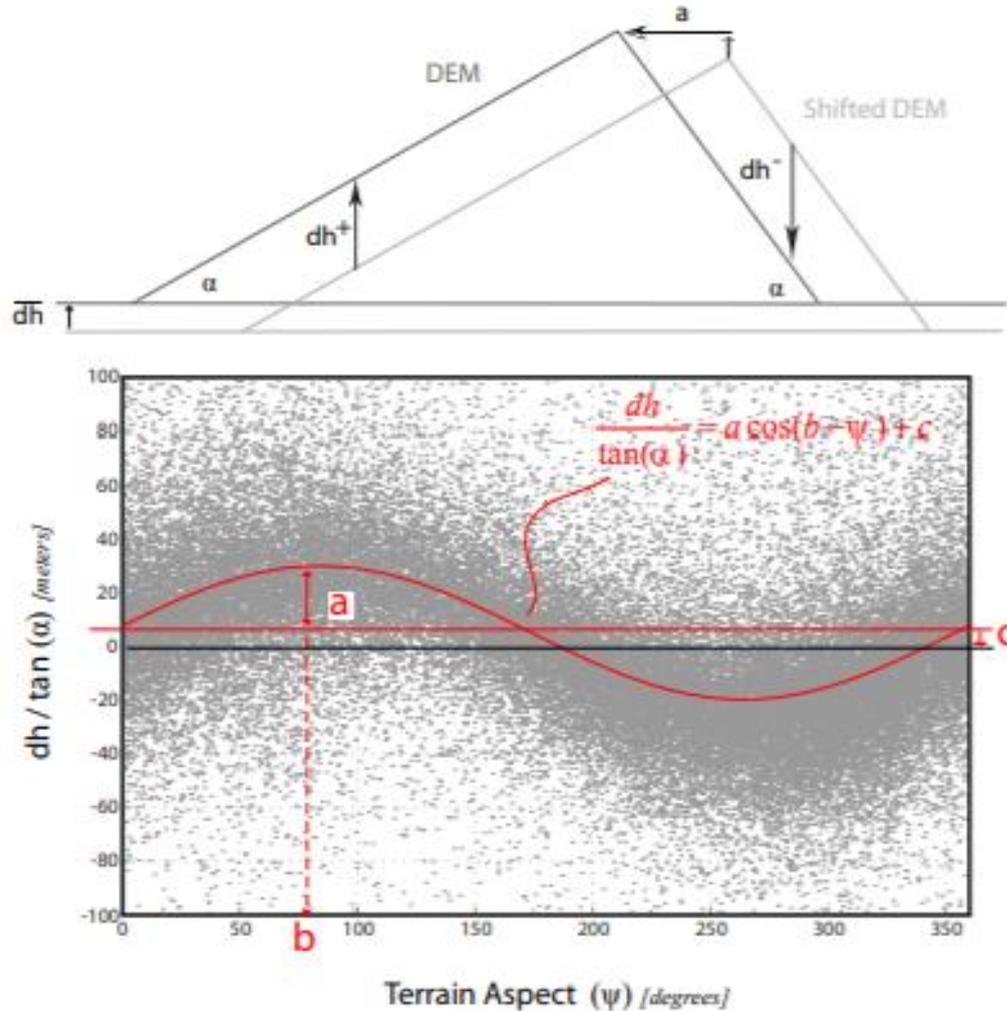


# DEM differencing & co-registration



# DEM differencing & co-registration

Nuth & Kääb, 2011



Possible methods:

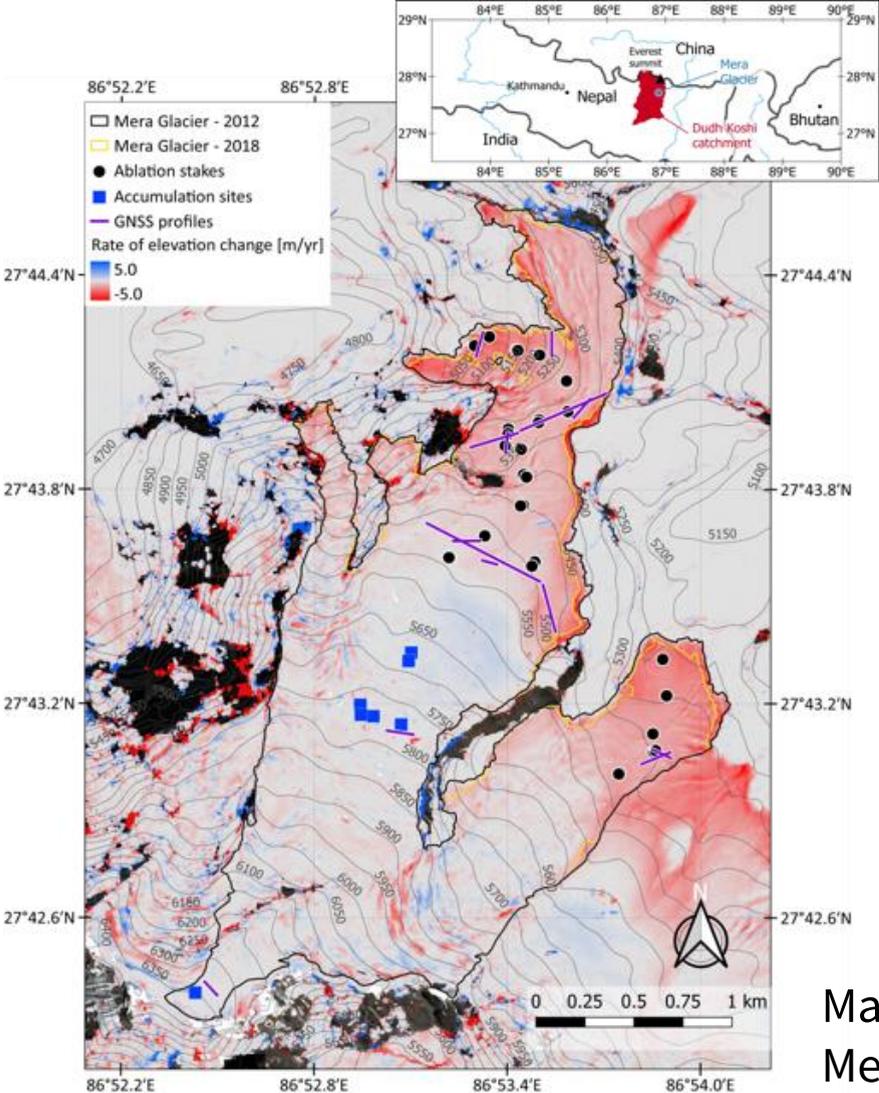
- Nuth & Kääb, 2011: translations & vertical shifts
- Iterative Closest Points (ICP): useful for rotations, scaling
- De-ramping: for 2D-plane tilt correction
- Vertical shifts

**Often one needs to apply a combination of several methods**

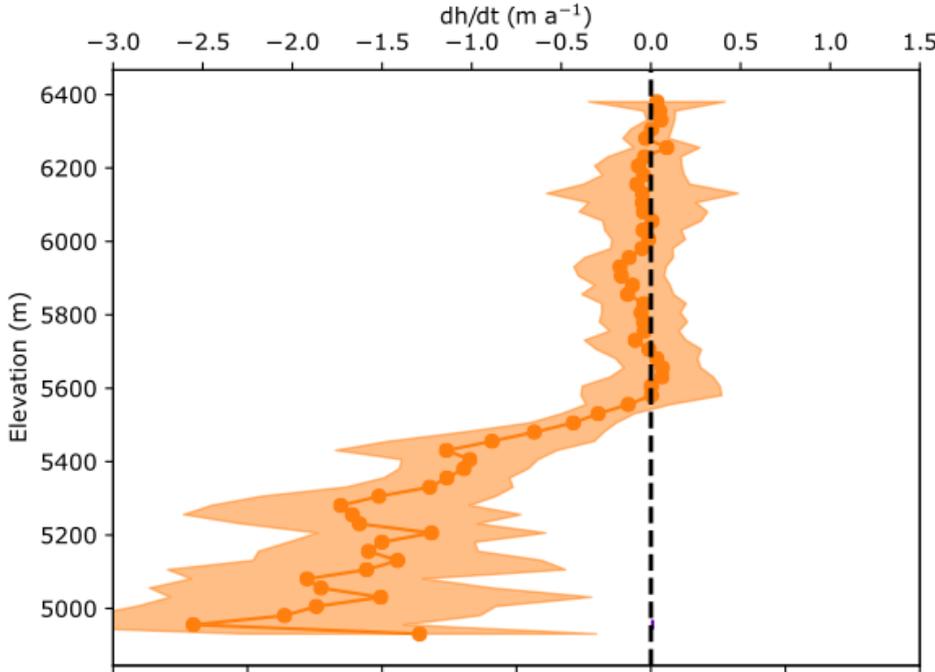
# Volume & mass change calculation



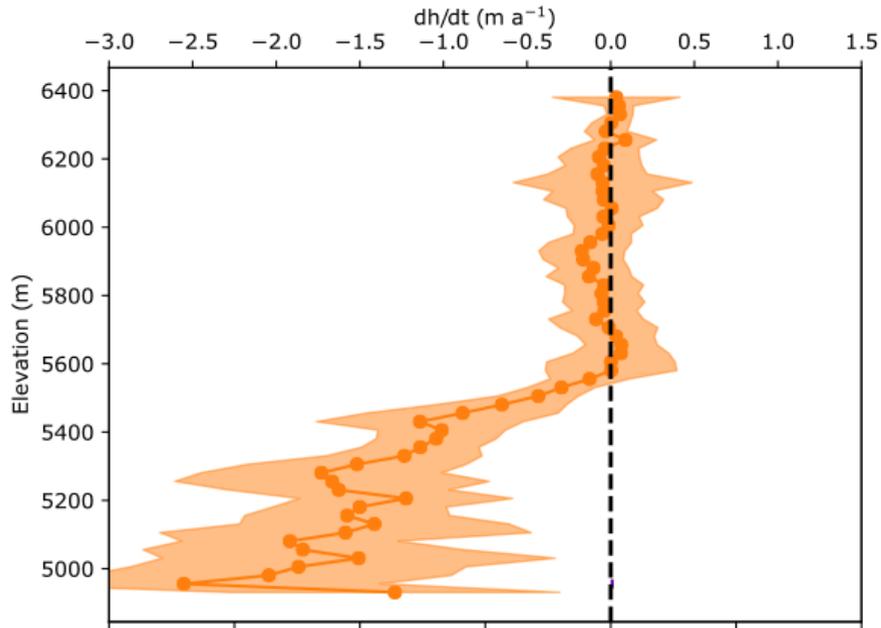
# Volume change calculation



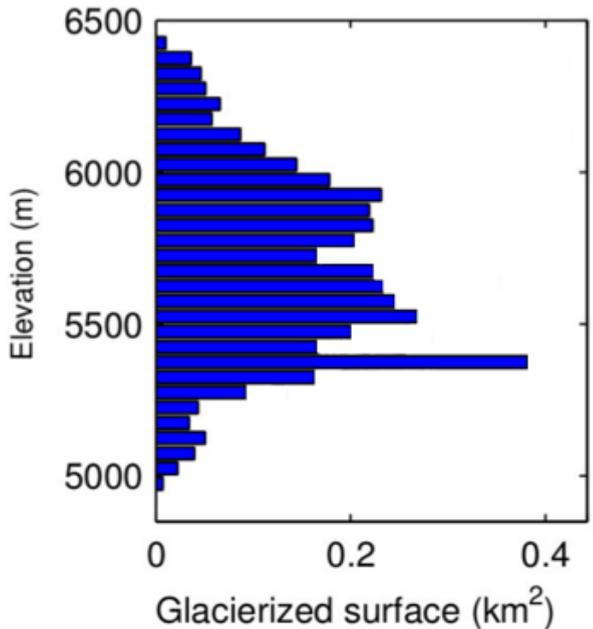
Map of elevation change of Mera Glacier from 2012 to 2018



# Volume change calculation



×



=

Glacier-wide mass balance

# Density conversion

Current reference (but far from perfect): Huss, 2013

$$\frac{\partial M}{\partial t} = \frac{\partial(\rho V)}{\partial t} \longrightarrow \Delta M = \Delta\rho V + \Delta V\rho,$$

$$\Delta M = \left( \frac{\Delta\rho V}{\Delta V} + \rho \right) \cdot \Delta V = f_{\Delta V} \cdot \Delta V$$

$$f_{\Delta V} = 850 \pm 60 \text{ kg/m}^3$$

# Density conversion: where does this come from?



Mera Glacier, November 2023

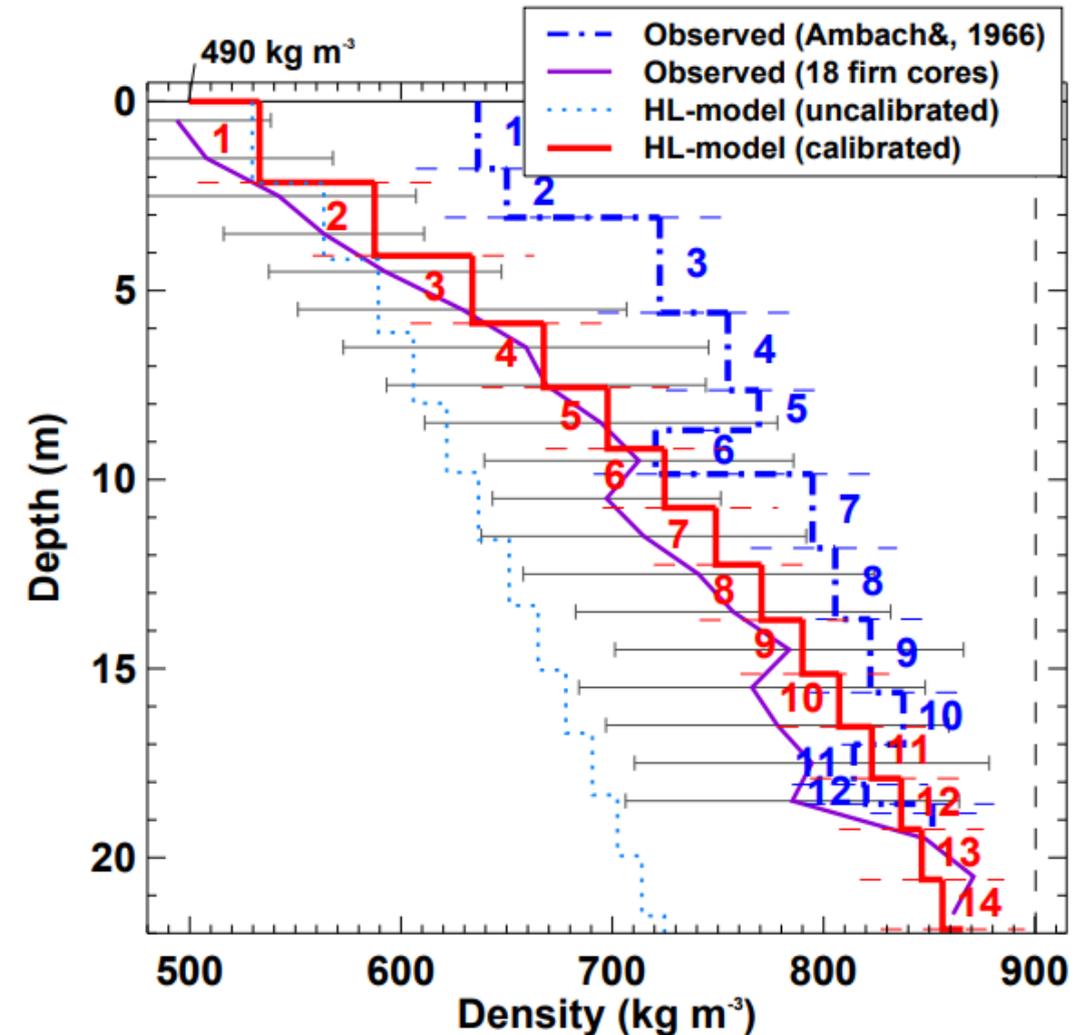


Argentière Glacier, May 2024

# Density conversion: where does this come from?

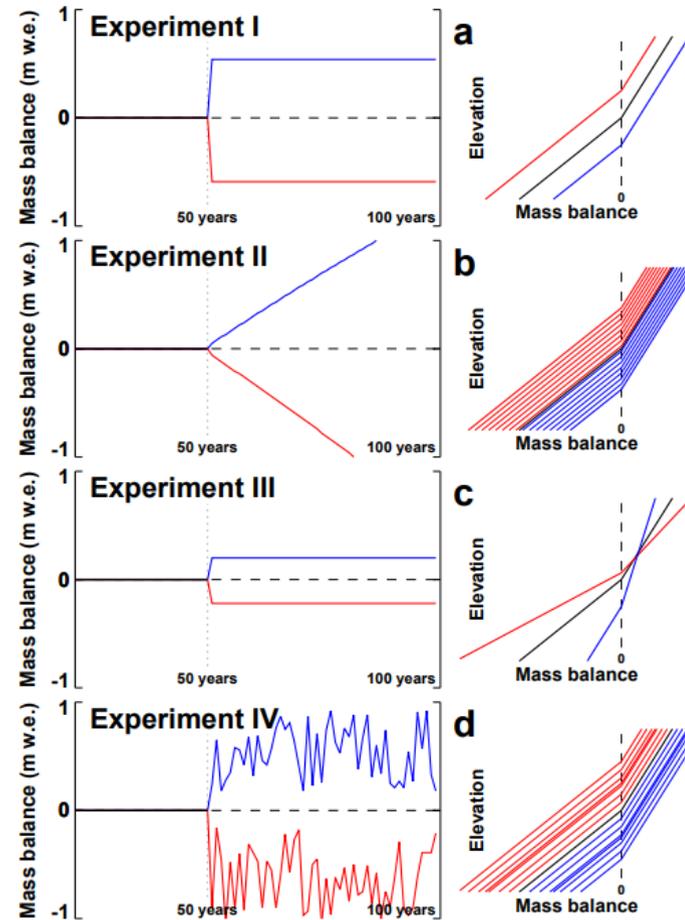
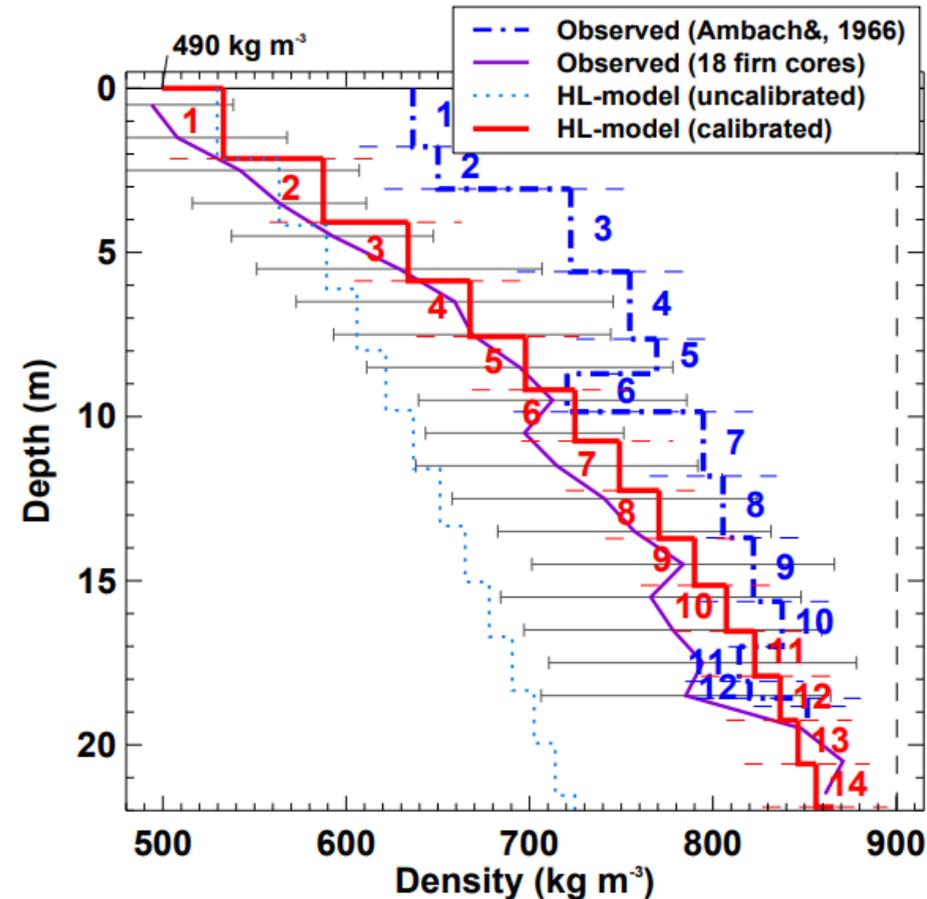
Orders of magnitude:

	Density (kg/m <sup>3</sup> )
Ice (no air bubbles)	917
Ice (mountain glaciers)	~900
Firn (upper snow layer after one year)	500-600
Fresh snow	50-200



# Density conversion: where does this come from?

Experiments on synthetic glaciers to compute  $f_{\Delta V}$ :

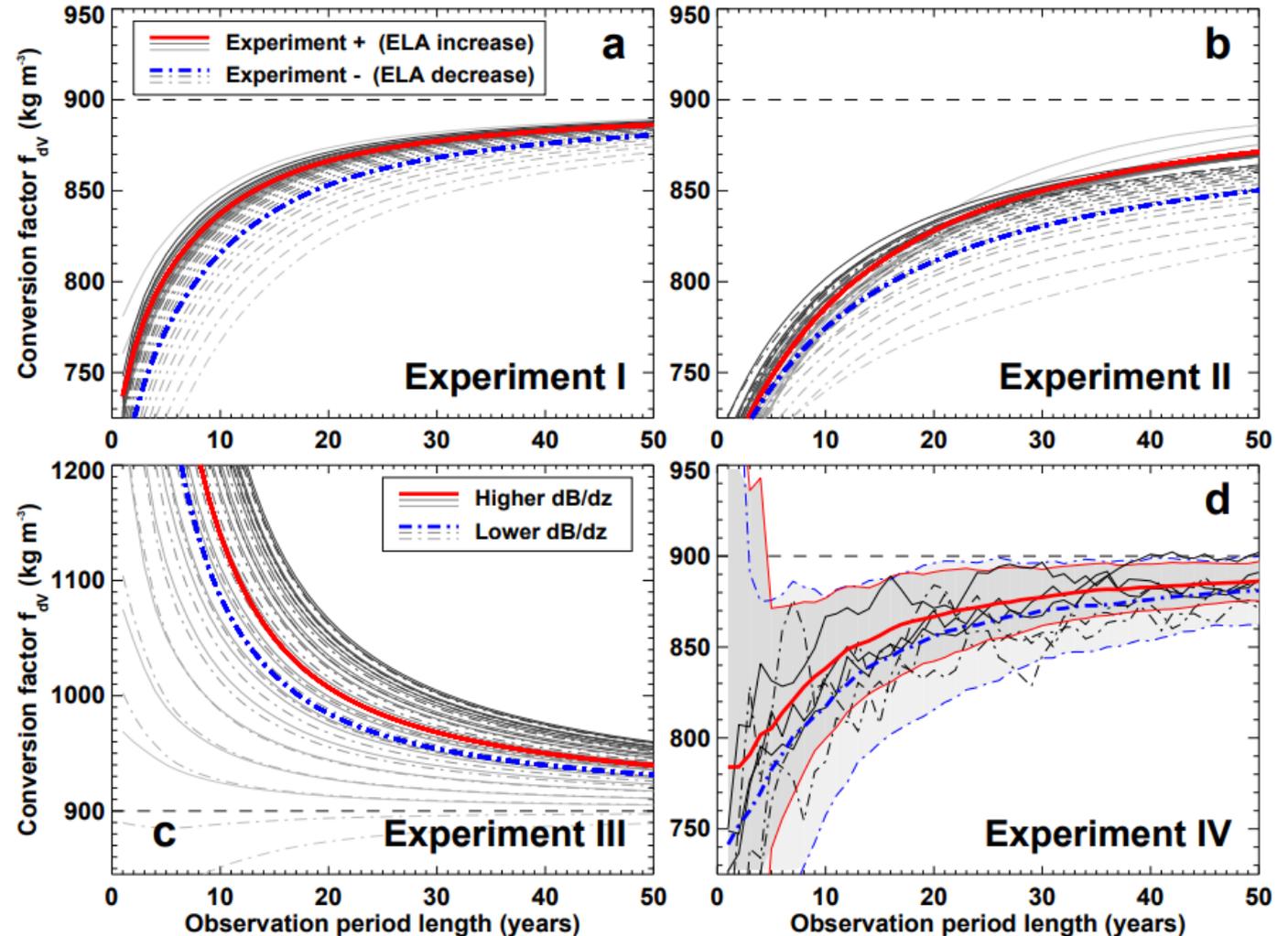


Glacier-wide  
volume & mass  
changes

# Density conversion: where does this come from?

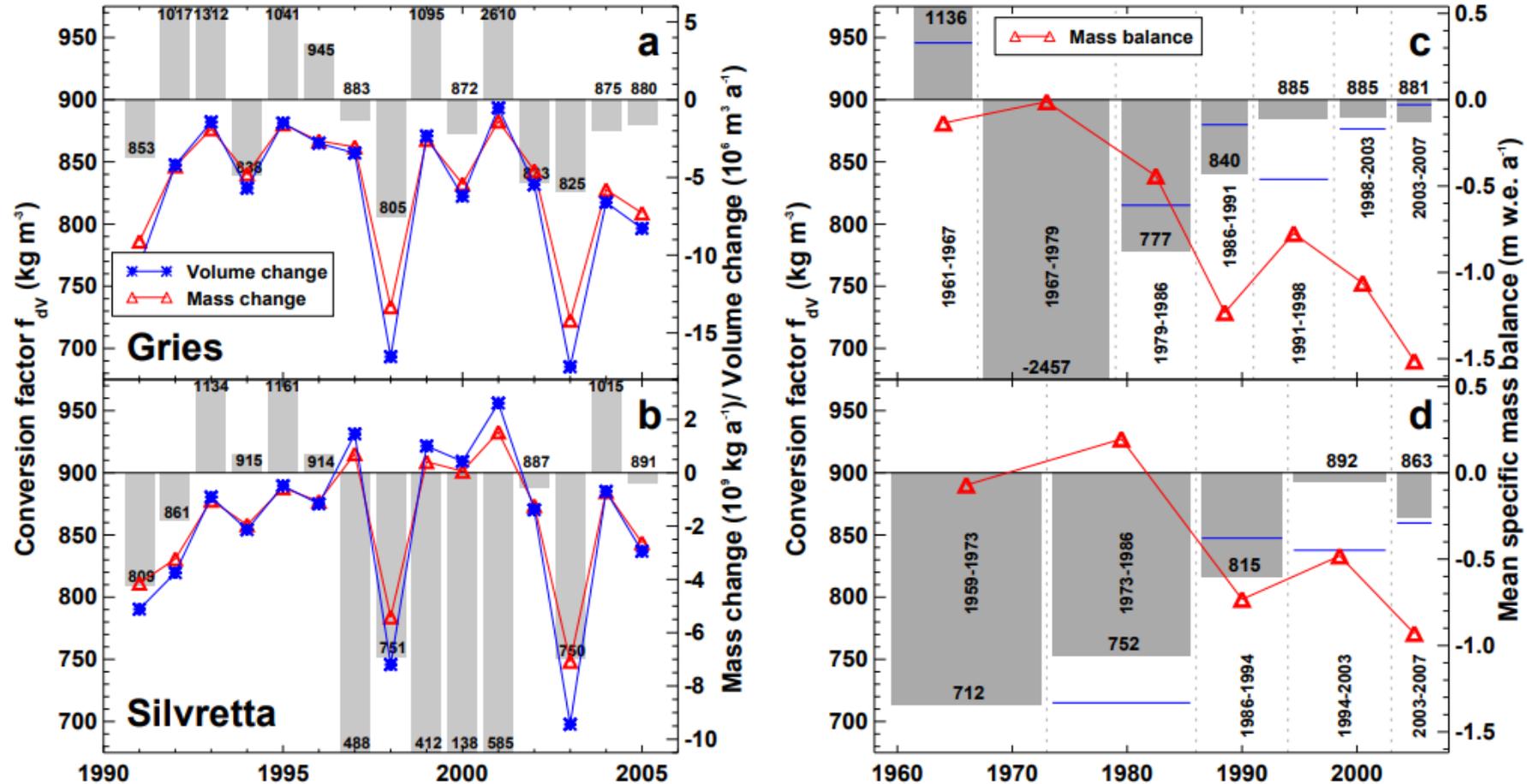
Experiments on synthetic glaciers to compute  $f_{\Delta V}$ :

$f_{\Delta V}$  can be higher than  $900 \text{ kg/m}^3$ :  
think about a glacier with constant  
mass but volume change due to  
compaction...



# Density conversion: where does this come from?

Testing on real glaciers:



# Density conversion: where does this come from?

Outcome:

- **The conversion factor  $f_{\Delta V}$  is, in most cases, below the ice density** which is used in numerous studies for calculating mass balance from geodetic volume change. The effect is relatively small, yielding a **typical overestimate of mass change of roughly 2–15 %**, but it is **systematic**, and thus needs to be accounted for.
- **$f_{\Delta V} = 850 \pm 60 \text{ kg.m}^{-3}$  is recommended in the case of periods longer than 5 yr**, with stable mass balance gradients, the presence of a firn area, and volume changes significantly different from zero. The conversion factor can however strongly diverge from this mean value for particular conditions, and the density assumption might represent a **significant component of uncertainty** in geodetically determined mass balances that is larger than previously assumed.

# What about uncertainties?



# Uncertainty estimation

Uncertainty on the mass change:

$$\sigma_{\Delta M} = \sqrt{(\sigma_{\Delta V} \times \rho)^2 + (\sigma_{\rho} \times \Delta V)^2} \times \frac{1}{A_{\text{mean}}}.$$

Uncertainty on the volume change (accounting for gaps in the data):

$$\sigma_{\Delta V} = \sqrt{(\sigma_{h_{\Delta \text{DEM}}} \times A_{\text{mean}}(p + 5 \times (1 - p)))^2 + (\sigma_A \times h_{\Delta \text{DEM}})^2}.$$

$p$  is the share of non-voided glacier area and then interpolated and assigning an uncertainty factor of 5 to this area (Berthier et al., 2014)

- Combination of uncertainties:**
- ***Volume -> elevation change***
  - ***Area***
  - ***Density***
  - ***+ some gaps needing to be filled & accounted for***

# Uncertainty estimation

Uncertainty on the elevation change:

Hypothesis: all systematic errors removed using stable terrain

The standard error  $\sigma_{\overline{dh}}$  of the mean  $\overline{dh}$  of the elevation changes samples  $dh$  can be written as:

$$\sigma_{\overline{dh}} = \frac{\sigma_{dh}}{\sqrt{N}},$$

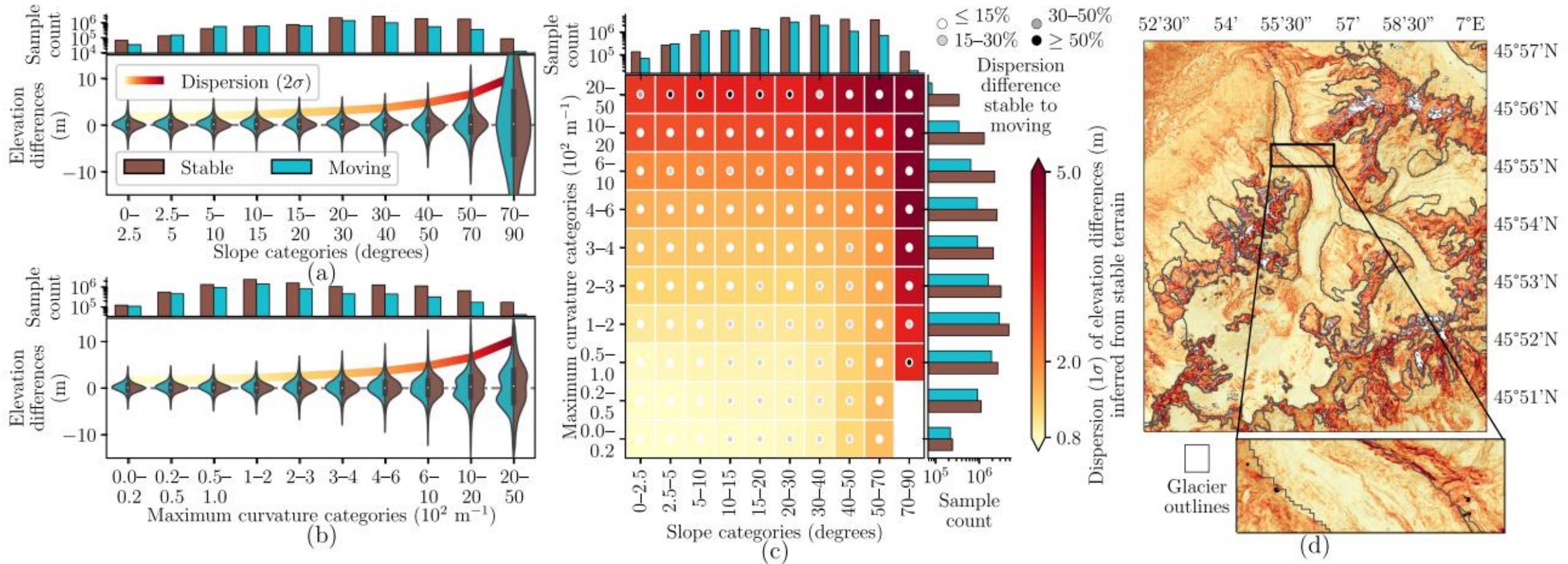
where  $\sigma_{dh}$  is the dispersion of the samples, and  $N$  is the number of **independent** observations.

-> If all the observations are considered independent,  $N$  = number of pixels over which we are averaging => uncertainties drop quickly. However, this is in general NOT the case:

Where  $n_f$  is given by the parameters of the variogram.

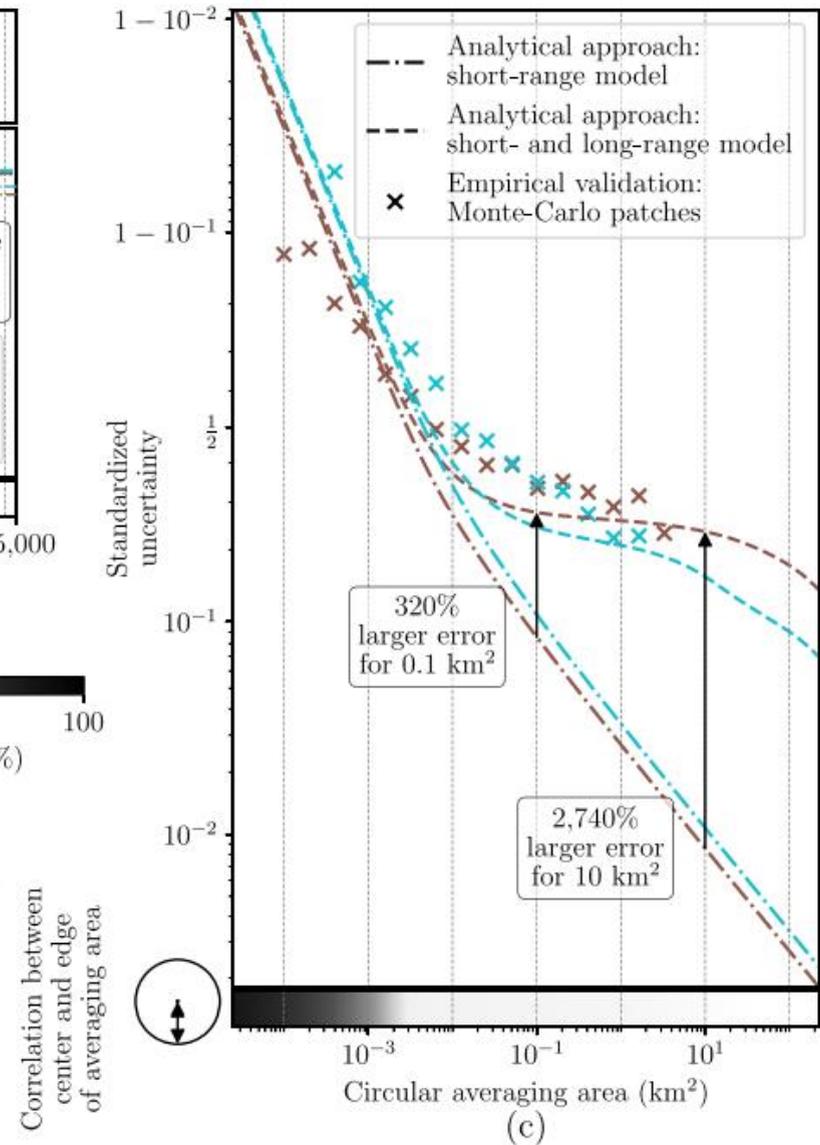
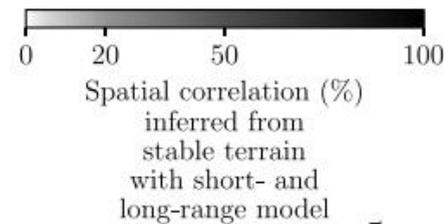
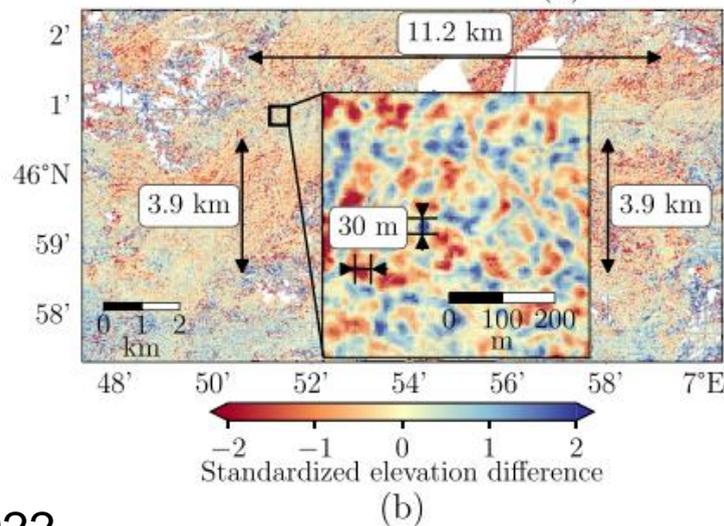
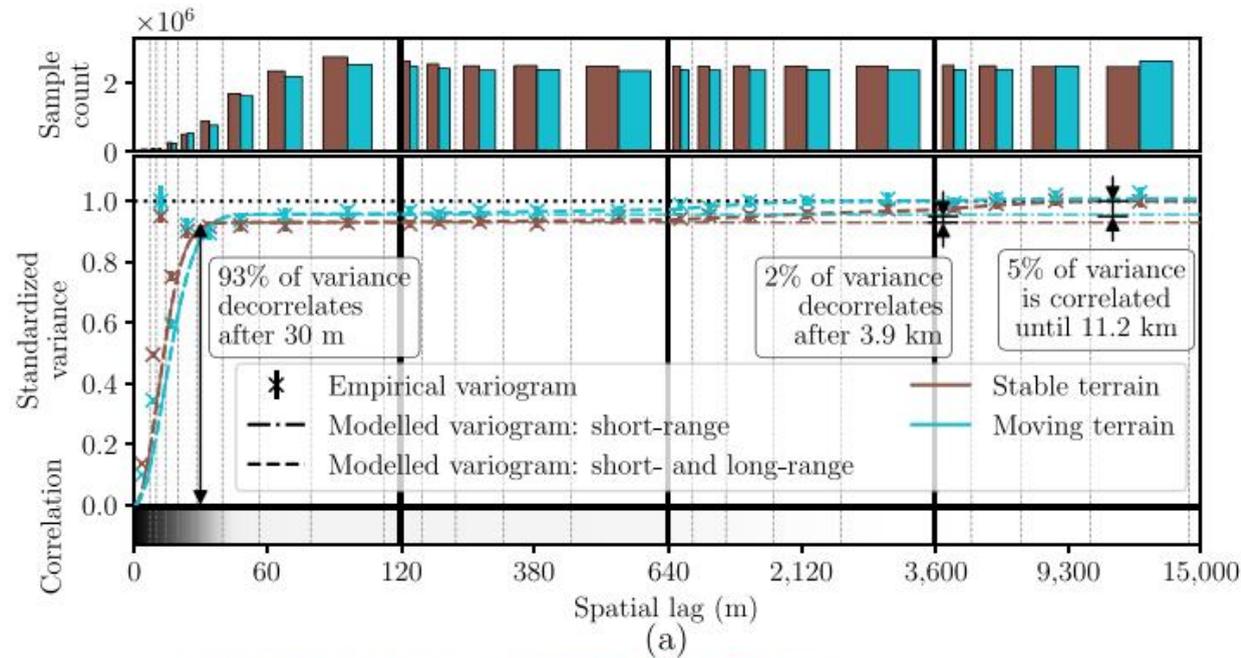
$$\frac{\sigma_{dh}}{\sqrt{N_{Pix}}} < \sigma_{\overline{dh}} = \frac{\sigma_{dh}}{\sqrt{n_f}} < \sigma_{dh}$$

# Uncertainty estimation: heteroscedasticity



-> Dependence on slope/curvature that needs to be accounted for to satisfy the 1st order stationarity hypothesis to compute the variograms ( $\sim$  distance is the only controlling variable for the variance)

# Uncertainty estimation: spatial correlation



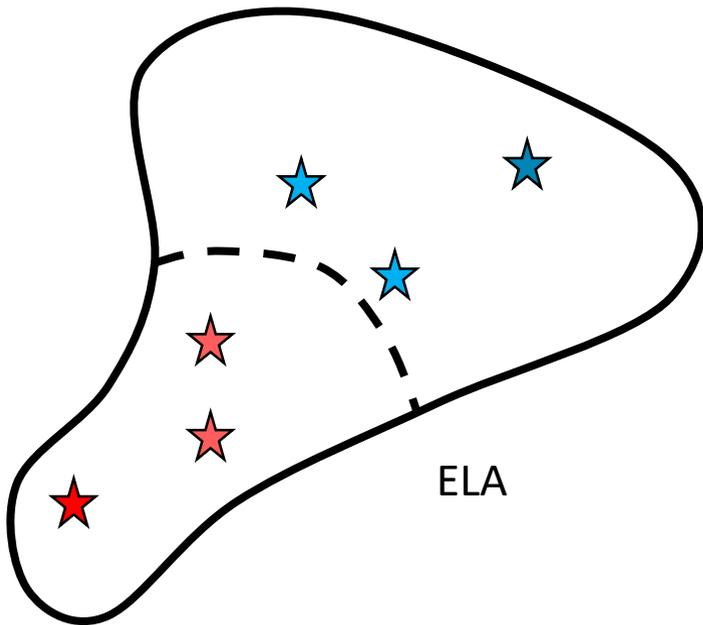
# Geodetic VS glaciological mass balance



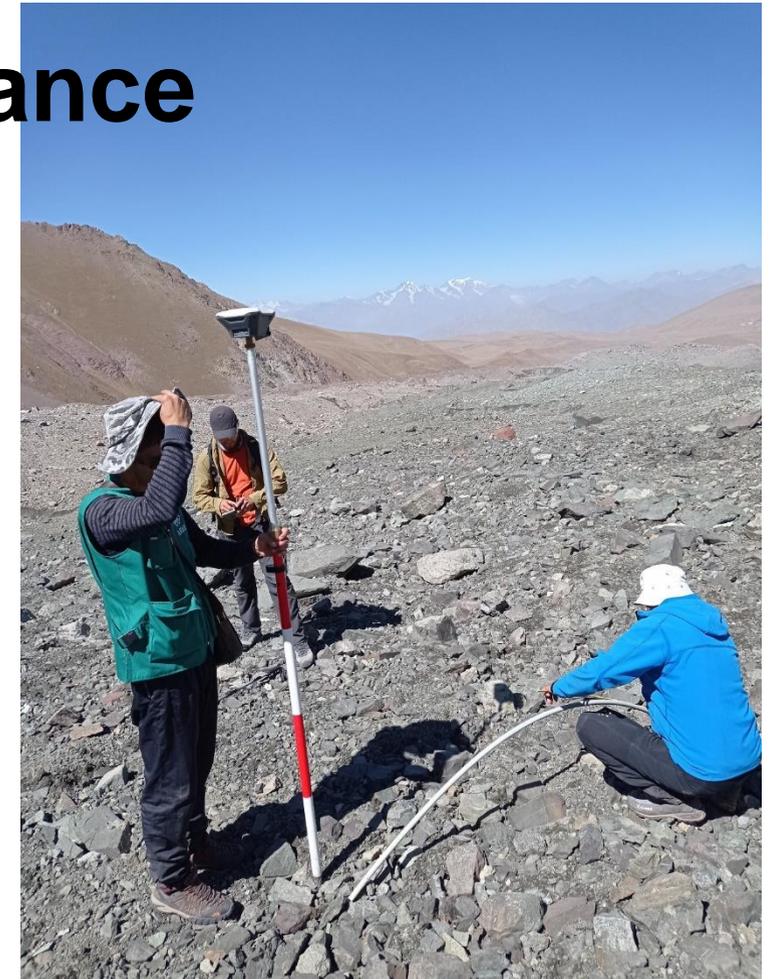
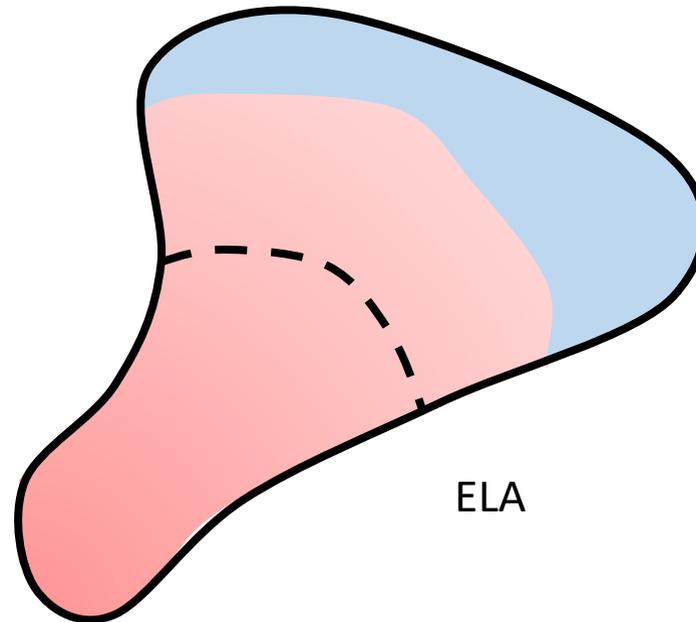
# Geodetic VS glaciological mass balance

- Stakes: **point-scale** measurements of **melt** over a given time
- GMB: **mass gain/loss** of the **whole glacier** over a given time

Stakes

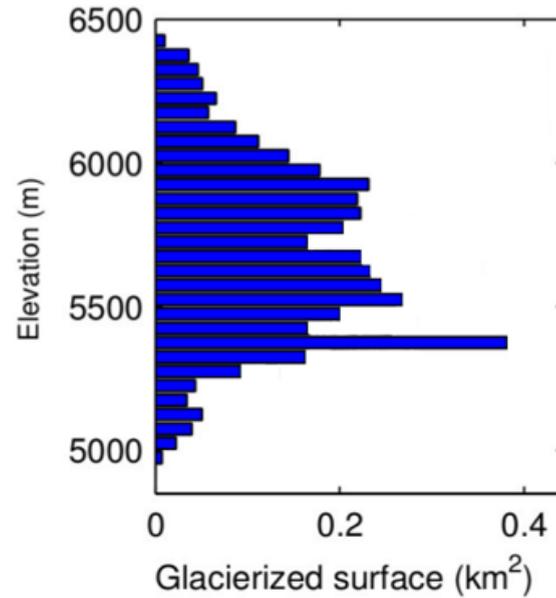
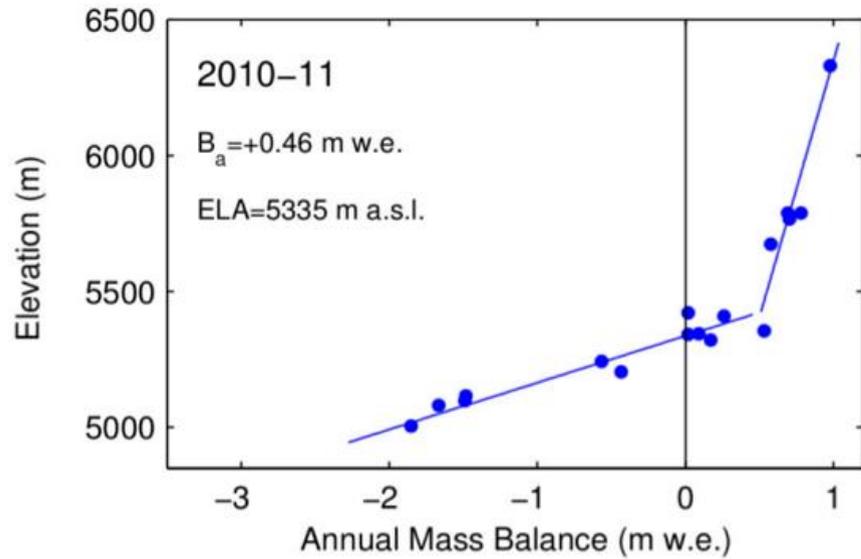


Full-glacier DH  
( $DH = DEM2 - DEM1$ )

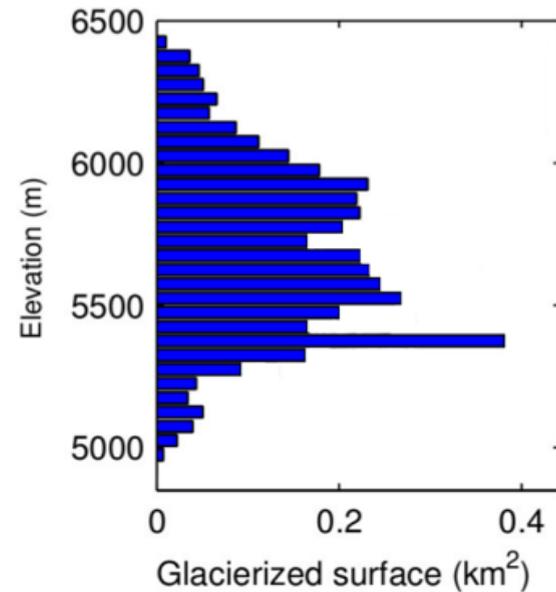
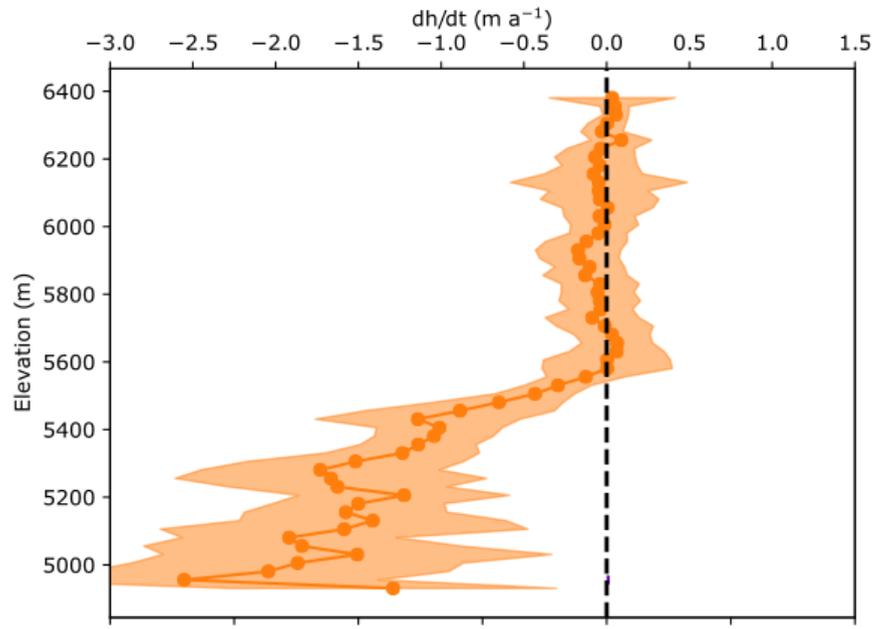


$$GMB = \text{Mass gain in accumulation zone} - \text{Mass loss in ablation zone} = \text{Sum of all } dH$$

# Geodetic VS glaciological mass balance



Glacier-wide  
mass balance



Glacier-wide  
mass balance



# Geodetic VS glaciological mass balance

Rate of mass change [kg]

Englacial mass balance [kg m<sup>-2</sup>]

$$\dot{M} = \int_A [\dot{b}_s + \dot{b}_e + \dot{b}_b] dA$$

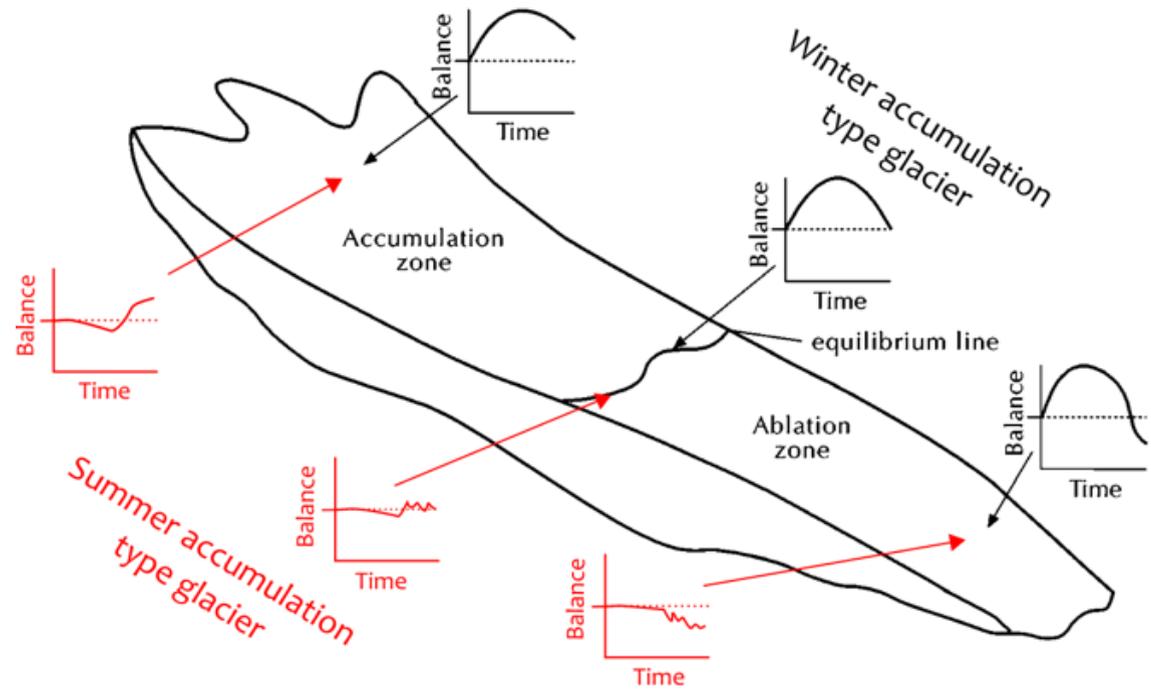
Glacier area [m<sup>2</sup>]

Surface mass balance [kg m<sup>-2</sup>]

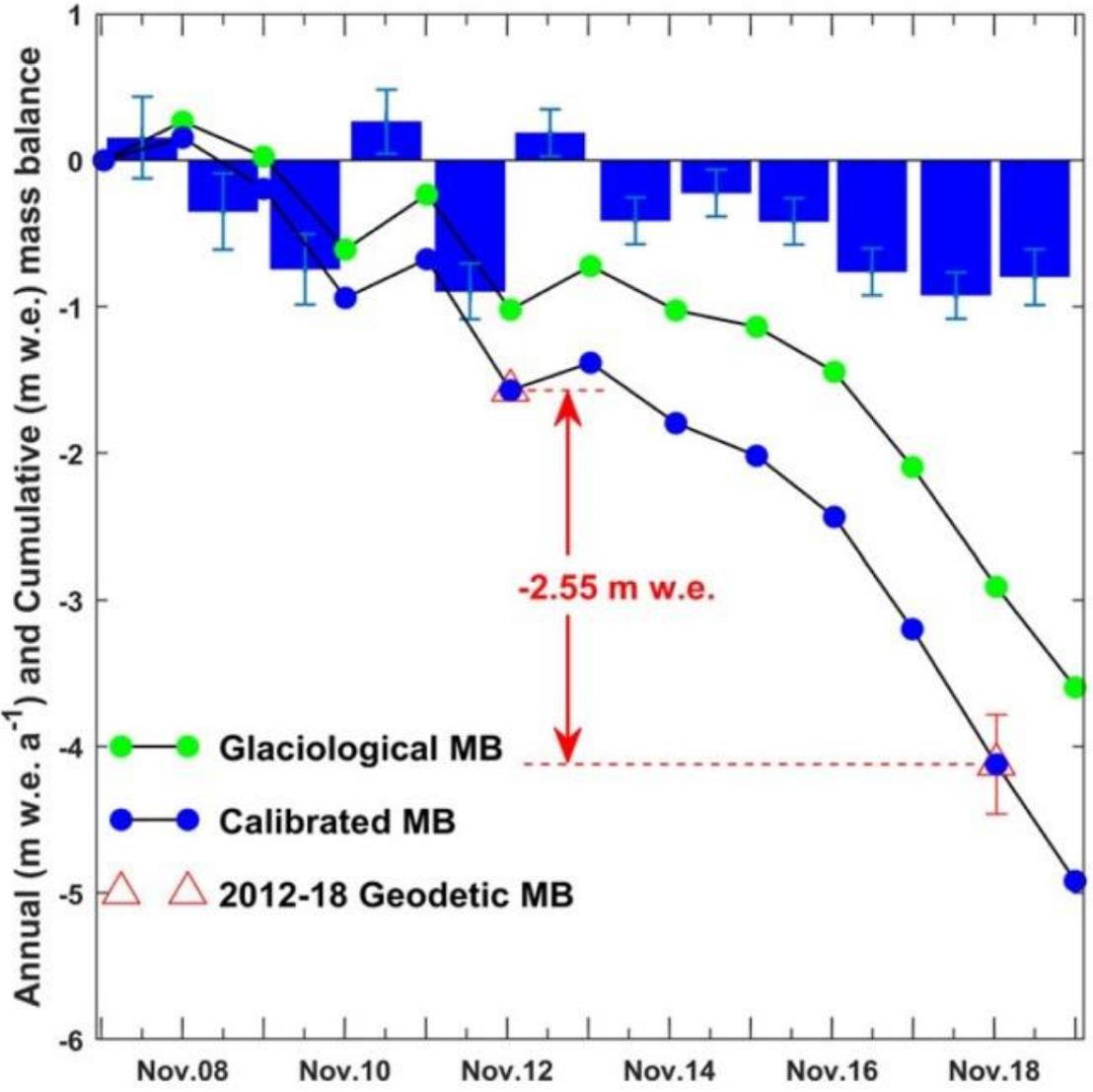
Basal mass balance [kg m<sup>-2</sup>]

For temperate glaciers:

$$\dot{M} \simeq \int_A \dot{b}_s dA$$



# Geodetic VS glaciological mass balance



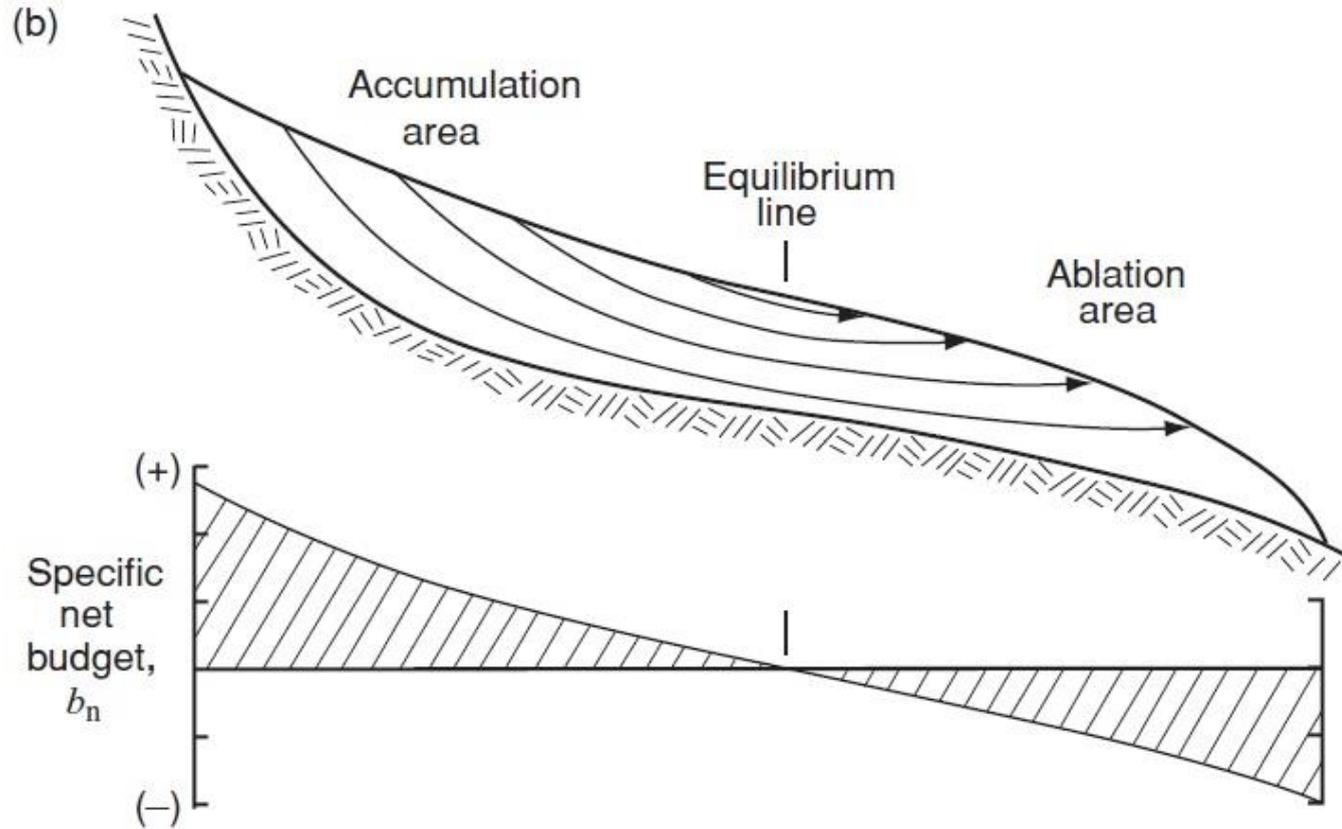
Glaciological mass balance series calibrated on geodetic series

# Geodetic VS glaciological mass balance

In the accumulation area, flowlines are pointing **downward (submergence)**.  
 In the ablation area, flowlines are pointing **upward (emergence)**.

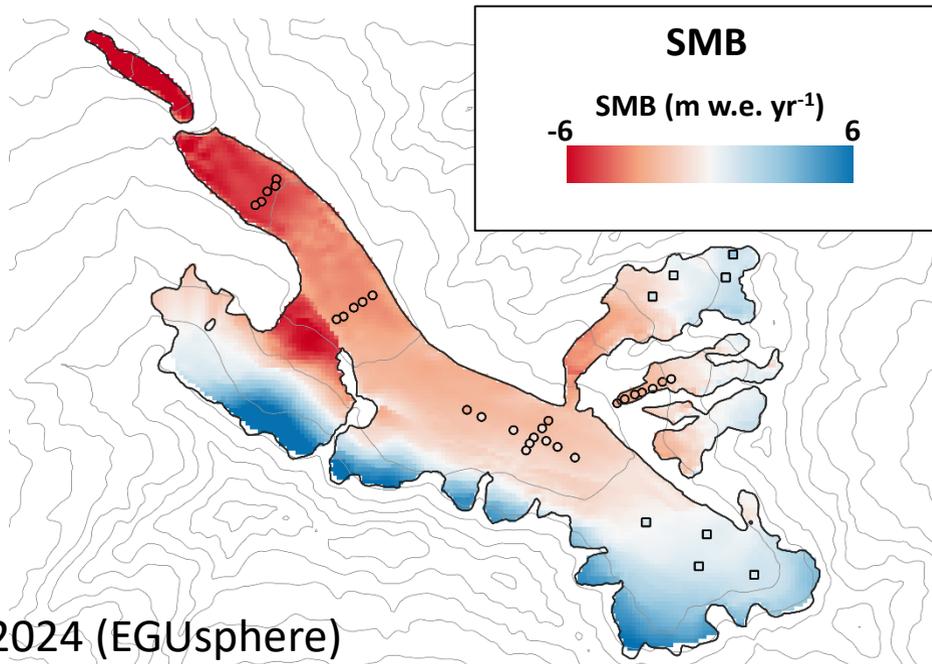
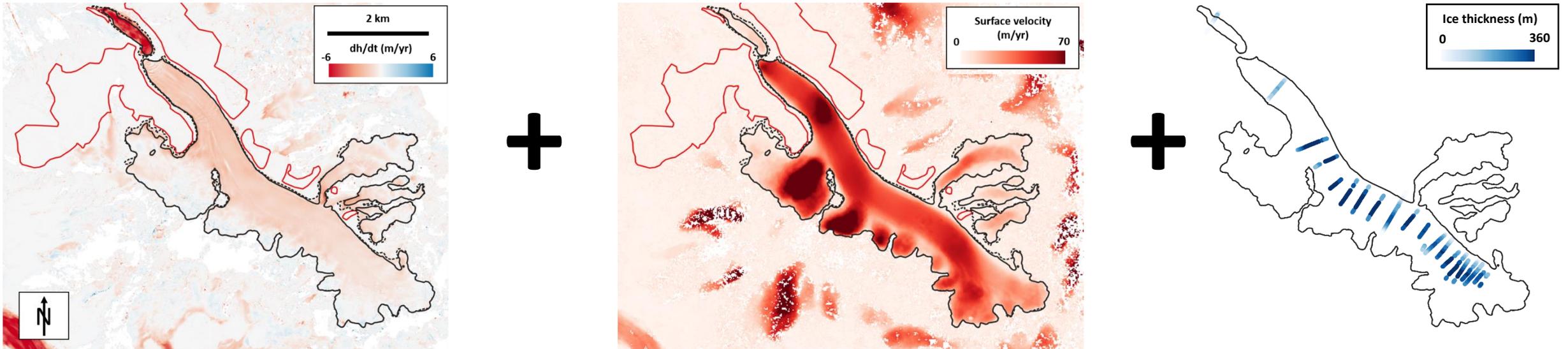
In every point of the glacier, we then have:

$$\frac{\rho_{dH}}{\rho_{H_2O}} \frac{dH}{dt} = \dot{b} - \frac{\rho_{\nabla q}}{\rho_{H_2O}} \nabla * \mathbf{q}$$



Longitudinal profile of a glacier [Hook, 2005]

# Distributed 'geodetic' mass balance?

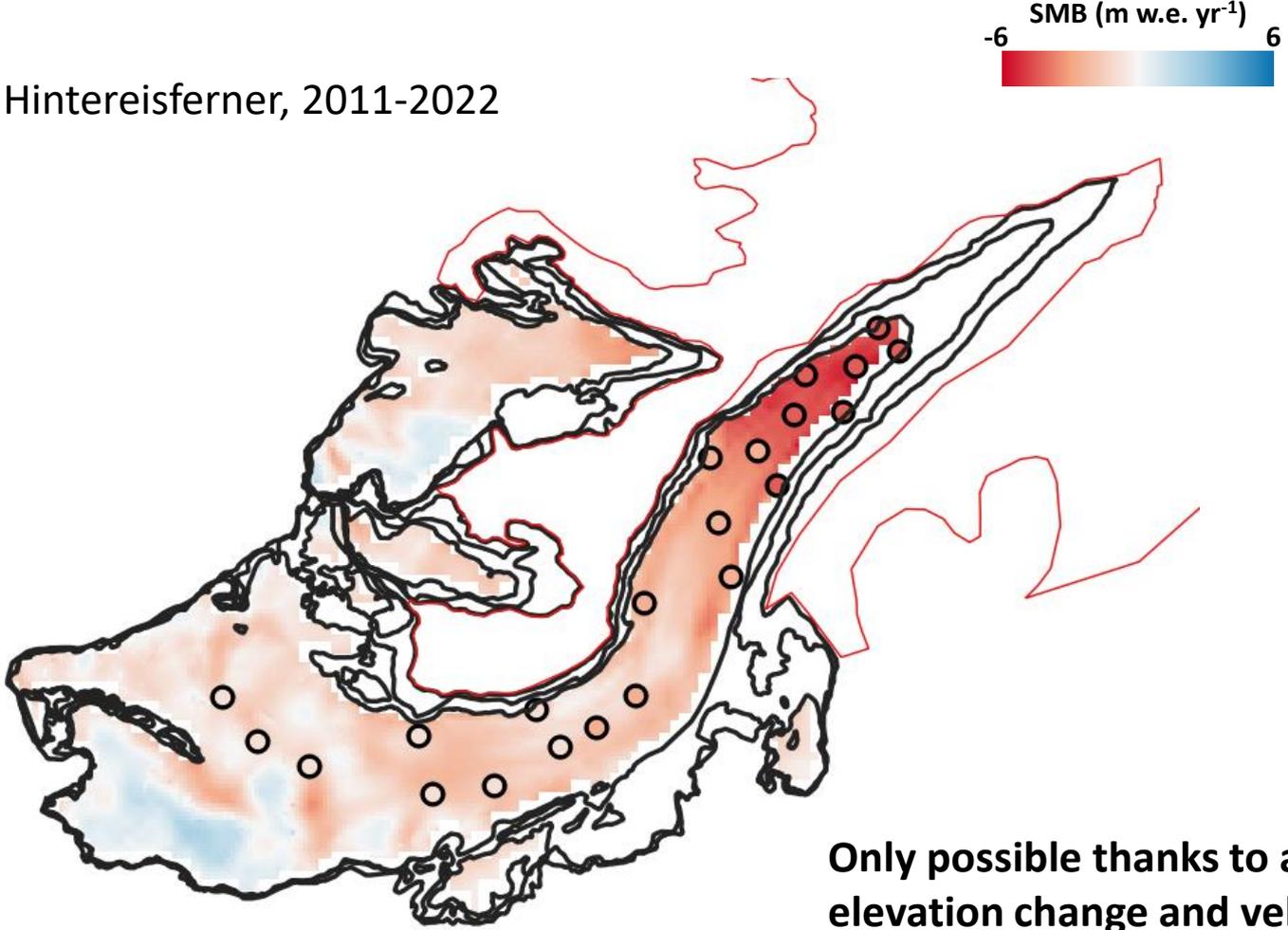


$$\frac{\rho_{dH}}{\rho_{H_2O}} \frac{dH}{dt} = \dot{b} - \frac{\rho_{\nabla q}}{\rho_{H_2O}} \nabla * \mathbf{q}$$

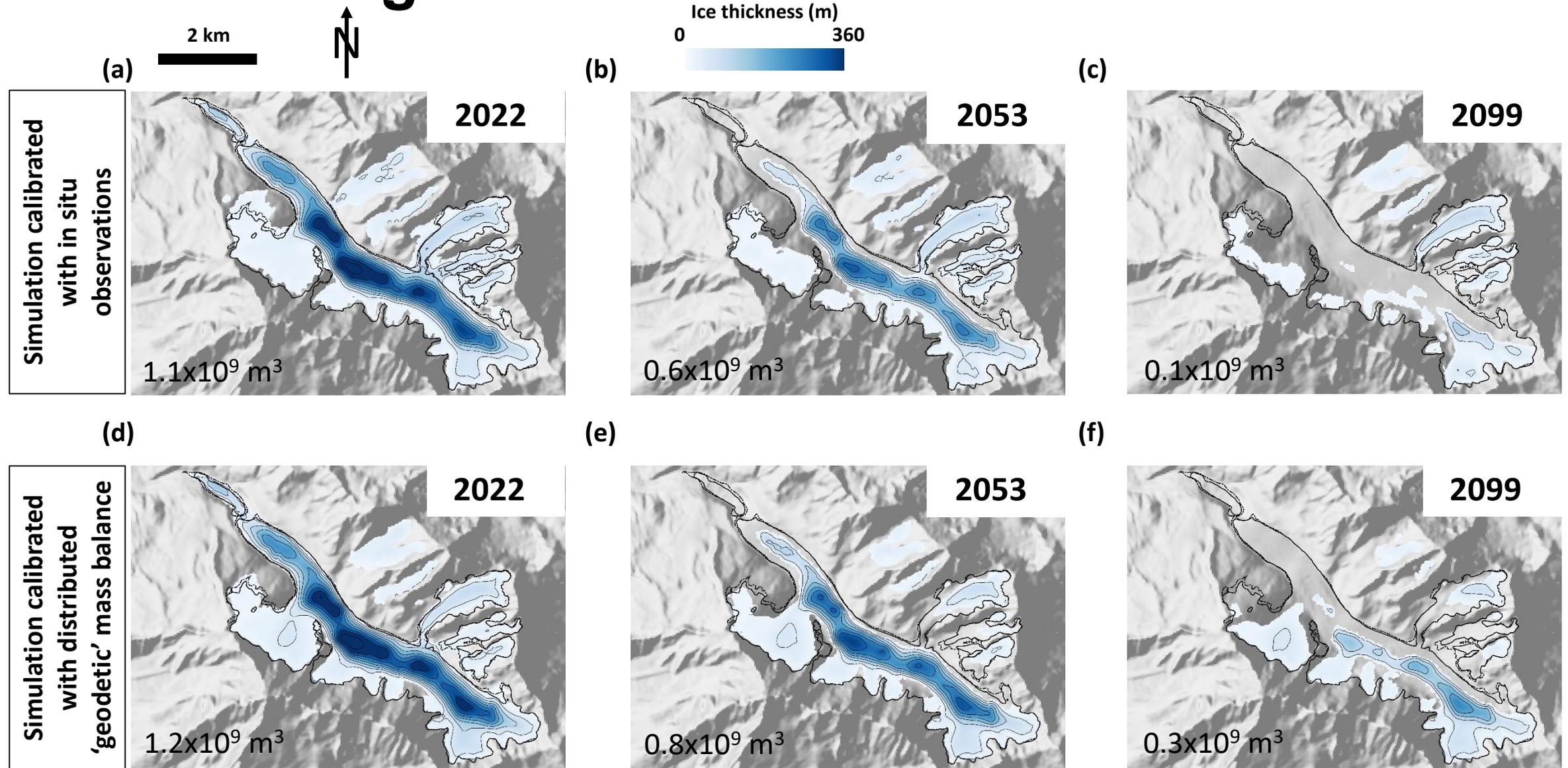
$$\mathbf{q} = h\gamma\mathbf{u}_s$$

=

# Distributed 'geodetic' mass balance?



# Distributed 'geodetic' mass balance?



Thanks for your  
attention!

